

GLOBAL PATTERN OF MESOSCALE VARIABILITY

IN SEA SURFACE HEIGHT

AND ITS DYNAMICAL CAUSES

Dake Chen Mark A. Cane, Alexey Kaplan, alexeyk@ldeo.columbia.edu mcane@ldeo.columbia.edu dchen@ldeo.columbia.edu Climate Modeling Group, Division of Ocean and Climate Physics

Lamont-Doherty Earth Observatory of Columbia University, Palisades, NY 10964, United States

1. INTRODUCTION

Using gridded satellite altimetry fields, we separate the mesoscale variability of sea surface heights into its spatial and temporal components. The ratio of these components shows a strong latitudinal dependence and to a large degree is controlled by Rossby radius of deformation for the first baroclinic mode. Further analysis results in the attribution of mesoscale variability in different areas to dynamical causes. Major portion of it can be explained as a local response to the mesoscale variability in the wind. The propagation of ocean eddies modify the pattern in such areas and nearby. Another major mechanism of generating high mesoscale variability is generation of instability waves in the areas of ocean countercurrents. Comparison with ocean models show that they mostly reproduce mesoscale variability due to current instabilities, but not the one caused by the mesoscale variability in winds. Eddypermitting ocean models reproduce temporal variability much better than the spatial variability, although the simulation of the latter is improved with the refinement of models' resolution. The pattern of mesoscale variability often appears as a pattern of model error and also as a pattern of gridding error on altimetrybased sea surface height maps.

Time-space separation of small-scale sea level height variability (a) Total SSV $\sigma_{4^o \times 1^o \times 1 \text{ month}}(s)$ (b) Short-term temporal variability $\sigma_{1 \text{ month}}([s]_{4^o \times 1^o})$



On the left: Separation of space-time sea level height SSV into temporal and spatial components for Ducet et al. [2000] 0.25 degree resolution 10 day gridded altimetry fields.

Below left: Ratios of temporal to spatial variability for ocean waves. Colors show the ratio for a monochromatic harmonic wave with a wavelength L and a period T. White lines show dispersion relations for ocean waves. Solid lines indicate Rossby (R) waves. Thin lines show off-equatorial Rossby waves for different latitudes indicated at black circles that mark points with the minimum allowable wave period for each latitude. Thick lines show the first 3 trapped equatorial Rossby modes. Dashes and dash-points show equatorial Kelvin (K) and Yanai (Y) waves respectively. White dots indicate Poisson (P) waves for the latitudes of 5 and 29 degree. Box scales of 1 month and 4 degree are used.

Below right: Zonal averages for the ratio of temporal to spatial variability estimated from the Ducet et al. [2000] analyzed altimetry fields (thick solid line) and the theoretical estimates (marked lines).





(a) Total SSV $\sigma_{4^{o} \times 1^{o} \times 1 \text{ month}}(s)$

120°E 150°E

sqrt [(ssv + { stvg10 stv }) squared] point mean: 3.3669 ± 2.4115 range [0.38037 to 34.942]

180° 150°W 120°W 90°W 60°W 30°W

POCM 4C model [*Tokmakian and Challenor* **1999**] Time-space separation of small-scale sea level height variability

30°E 60°E

stvg10 stv

(b) Short-term temporal variability $\sigma_{1 \text{ month}}([s]_{4^{o} \times 1^{o}})$

150°W 120°W 90°W 60°W 30°W 0°

120°E 150°E

point mean: 1.7174 ± 0.99461 range [0.25609 to 13.394]

Below: Contours of the surface wind SSV are shown over color



High above: Simulation of sea level height error and SSV in Monte Carlo experiments with a linear model forced by noise designed to imitate errors in the wind forcing. Shown are model responses to the noise forcings with short (1 degree) and relatively long (20 degree zonal and 10 degree meridional) spatial decorrelation scales. Above: Temporal decorrelation scale is 0.25 month. SSV in pseudostress and sea level height response: Contours of SSV RMS in zonal wind pseudostress, (m/s)**2, are shown over the color pattern of SSV, and same but for contours of the SSV RMS in the sea level height response of a linear model to the random wind with 1 degree spatial and 0.25 month temporal decorrelation scales. Below: A contrast between ROMS NPac SSH RMS with satellite altimetry (Curchitser et al, 2005) and SSH response to AMIP-based surface flux perturbations (Borovikov et al. 2005).



plots of the total SSH SSV and their model-to-altimetry ratios.

DLTR TOPEX+ERS analysis: small-scale



180° 150°W 120°W 90°W Longitude

POCM/DLTR Ratio: small-scale



Longitude

ROMS NPac run, forced by NCEP-NCAR Reanalysis fluxes Small-scale spatial variability Short-term temporal variability Temporal-to-spatial variability ratio



ROMS NPac run, forced by QuikSCAT winds



Satellite altimetry fields

deviations of these box means computed within 1 month intervals, and their ratios. Shown are ROMS NPac run forced by Reanalysis fluxes and by QuikSCAT winds (Curchitser et al. 2005) and Ducet et al. (2000) blended analyses of Topex/Poseydon and ERS-1,2 satellite altimetry. All statistics are averaged for for the QuikSCAT run period of 2000-2002.

On the left: Small-scale spatial

variability inside 4 degree by 1 degree

boxes, temporal variability (standard

On the right: Same as on the left but

ROMS NPac run, forced by NCEP-NCAR Reanalysis fluxes

ROMS NPac run, forced by QuikSCAT winds



ROMS CCS run, 3 km resolution



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