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LINKING THE SUBTROPICAL AND SUPOLAR OCEANS: EXCHANGE IN THE NORTH ATLANTIC

(ATMOSPHERIC FORCING, SSH STRUCTURE, CARBON,
AMOC, AMV)

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THANKS TO JIM CARTON FOR SODA ASSIMILATION DATA AND JPL FOR ECCO DATA

Why are we interested in the subpolar zone, especially in the N Atlantic?

people and ecosystems and natural resources: all are on the move.

concentrated sector of the global climate system's meridional heat flux/fresh water flux in both ocean and atmosphere 'headwaters' of the AMOC

center of variability: decadal, multidecadal Arctic rim is where cryosphere change may have greatest impact

leading edge of global warming yet with dynamical cycles producing extended cold periods

strongest carbon uptake of any oceanic region

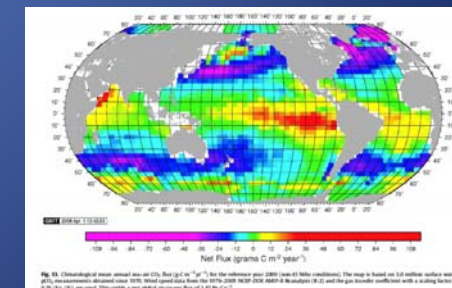
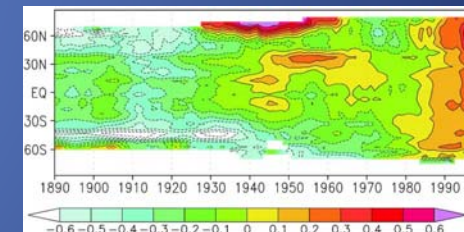
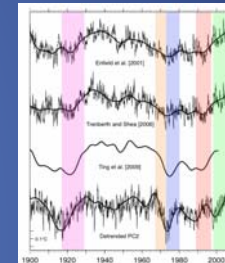
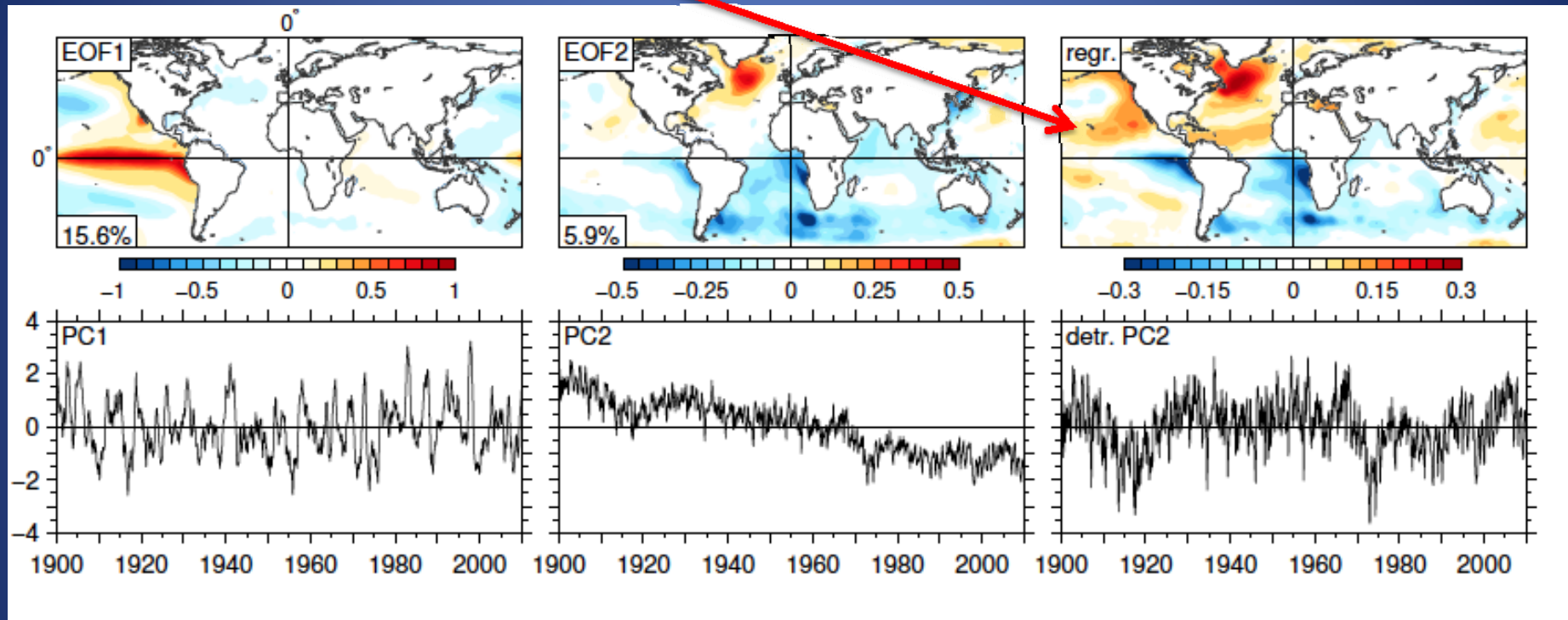
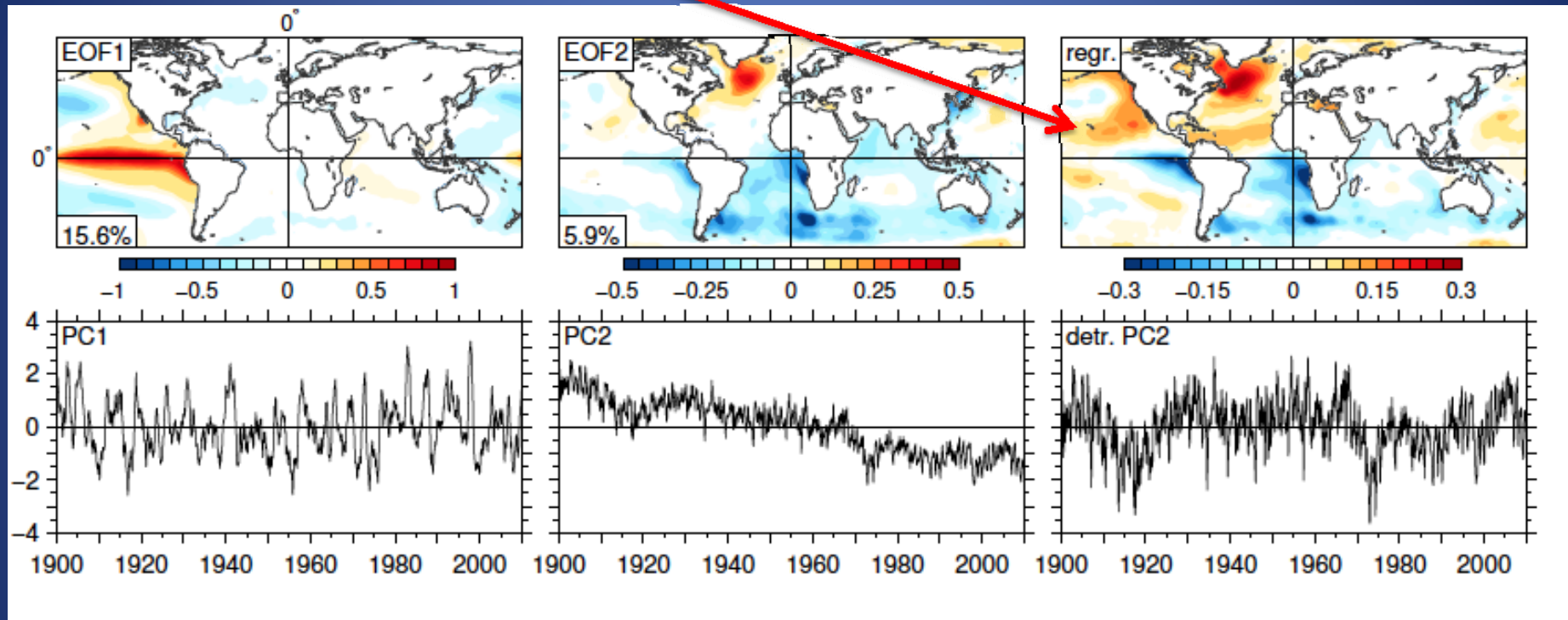


Fig. 18. Climatological mean annual mean (a) $\overline{Q_{\text{net}}}$ (W m⁻²) for the reference year 2000 (see-63 Niles conditions). The map is based on 2.0 million surface water cells, the contours are shown every 10%. Wind speed data from the 1970s, 2000 WRF 2000 AMIP-9 Wavelength (10, 2) and the gas transfer coefficient with a scaling factor of 0.20 (10, 10) are used. This panel is a global overview from 180° to 180°.

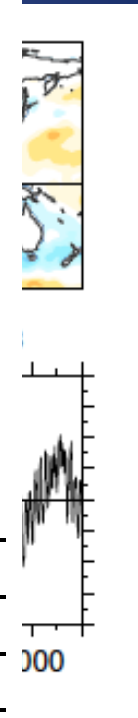
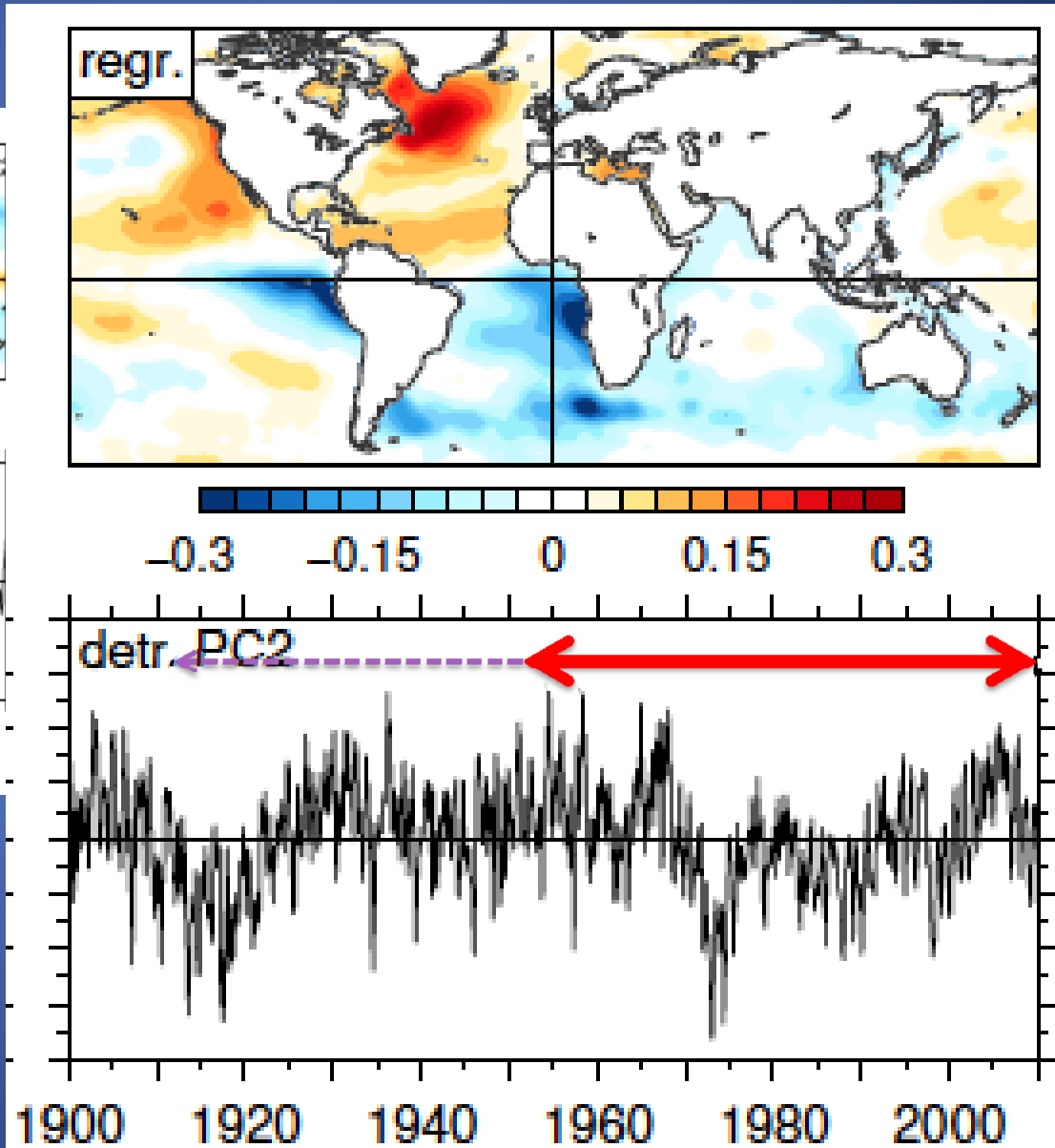
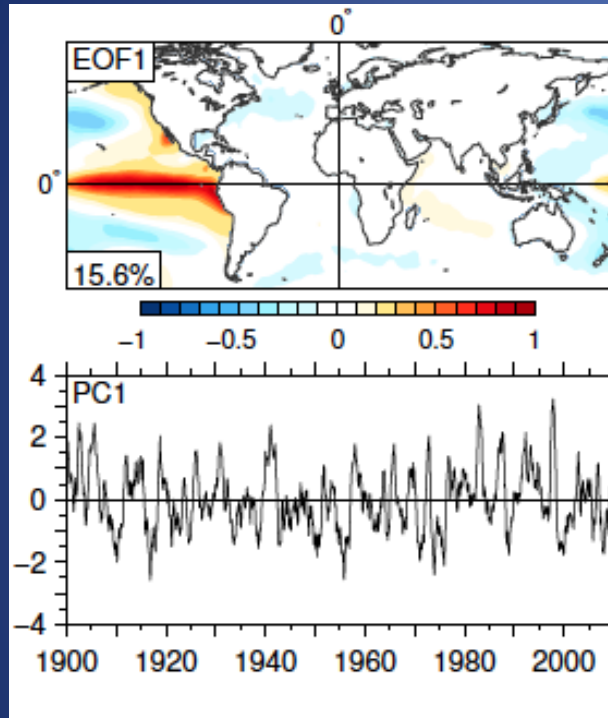
Wallace & Smoliak 2010 AMV: COWL or MOC dynamics?
detrended SST EOF-2 is a typical AMO/AMV timeseries (right)



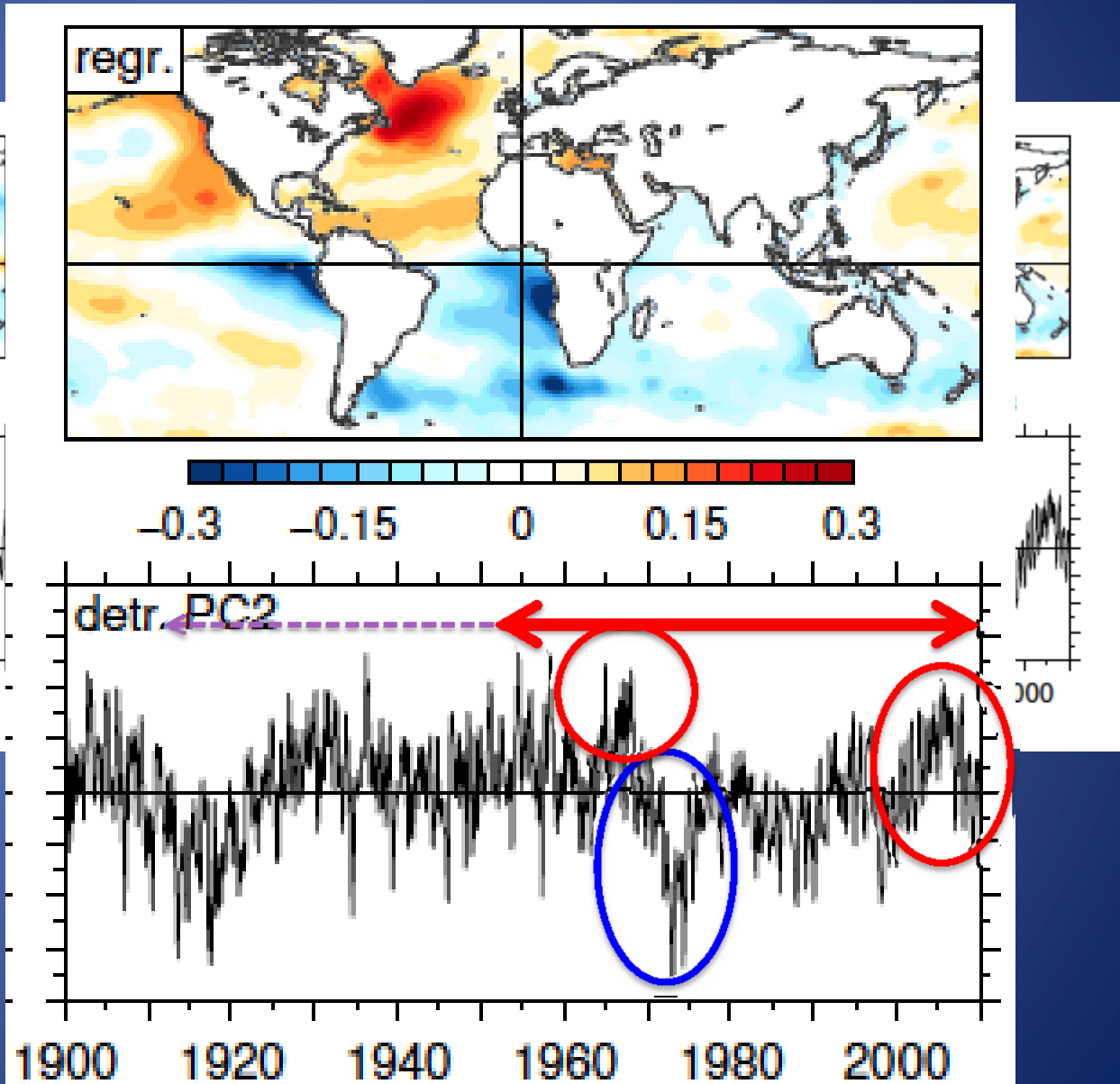
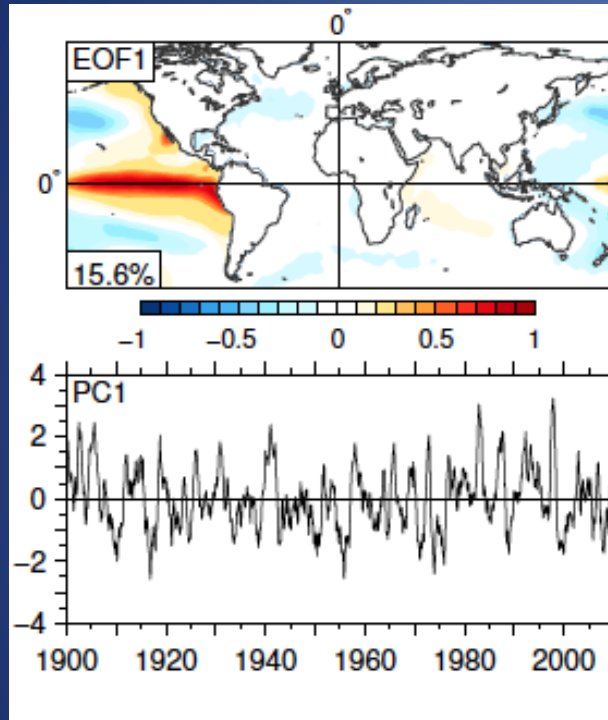
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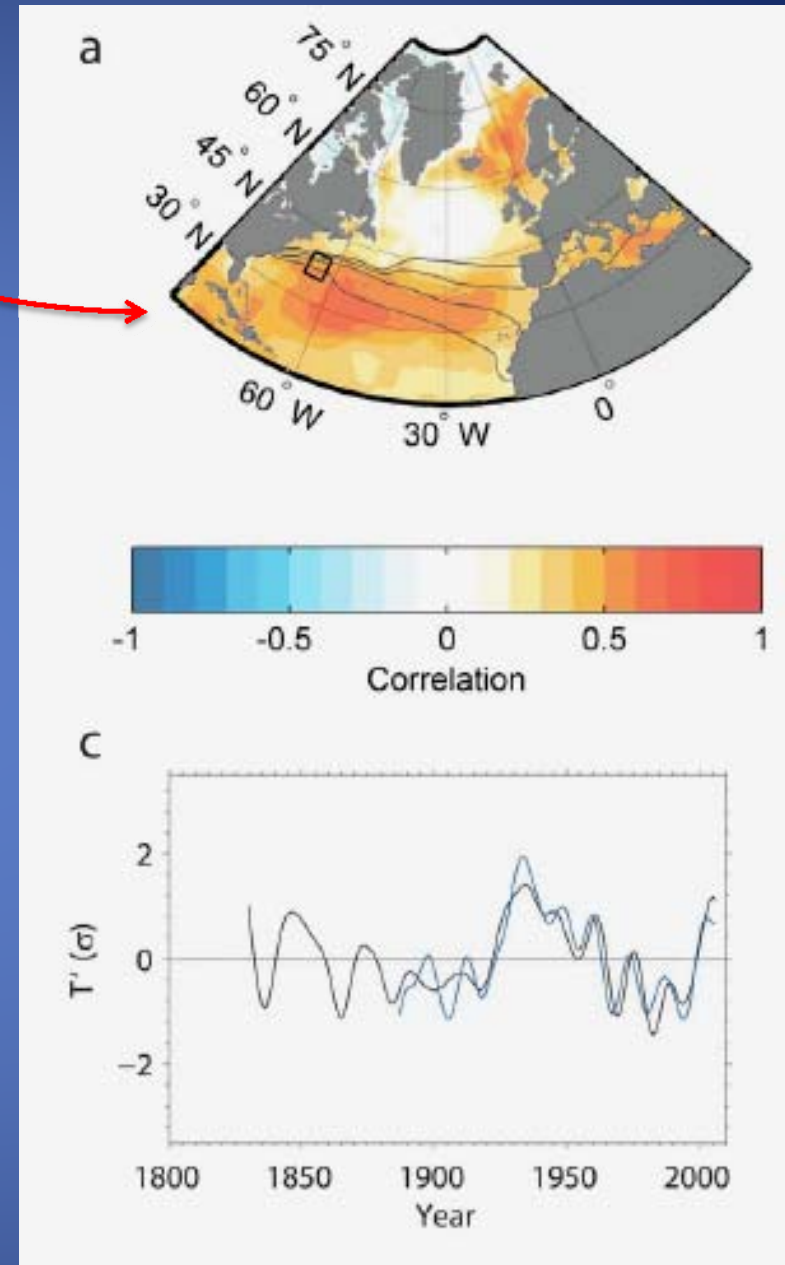
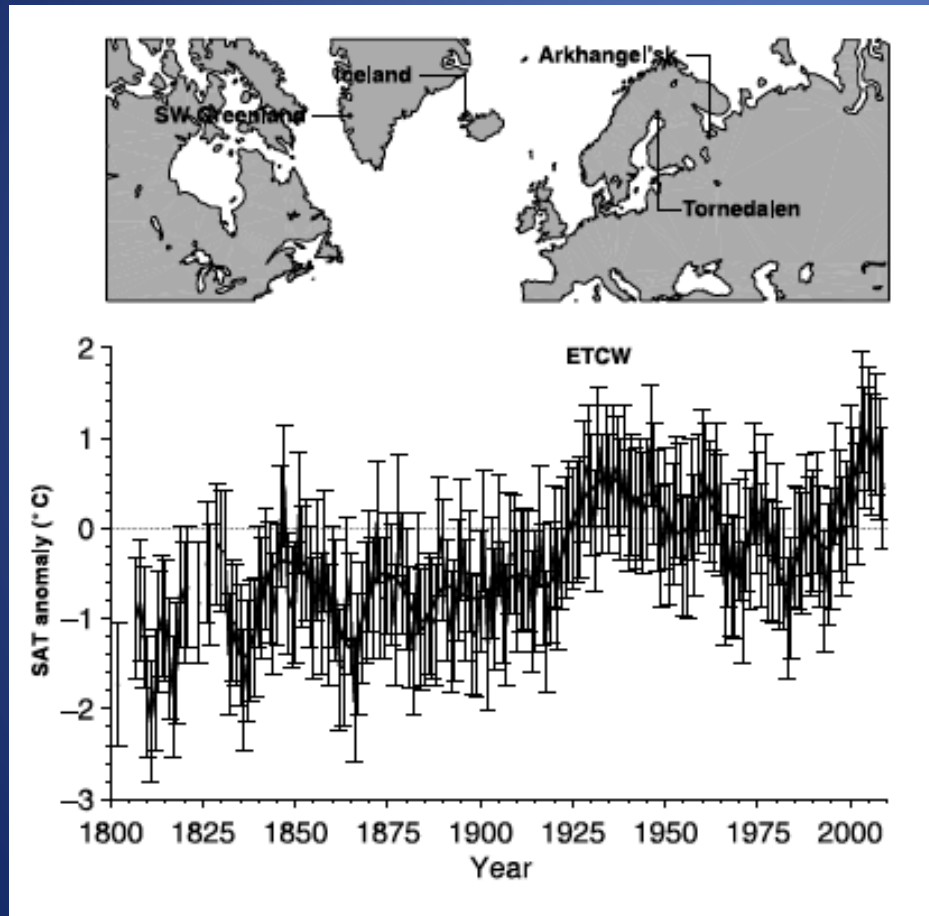
Wallace & Smoliak 2010 AMV: COWL or MOC dynamics?

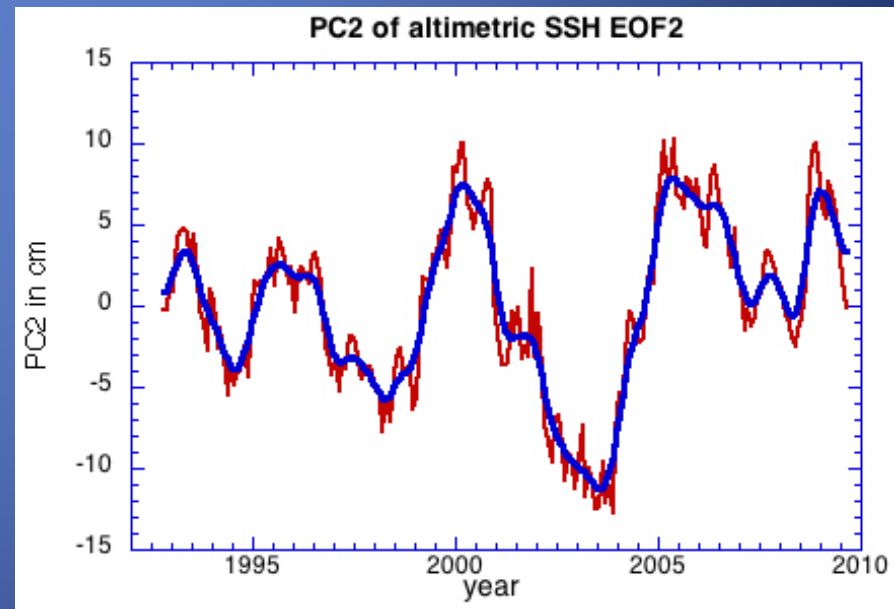
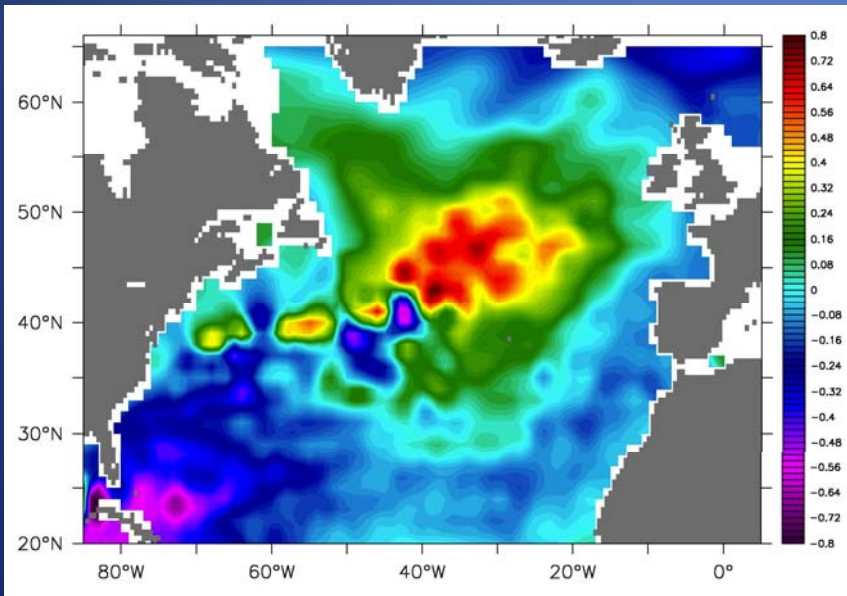
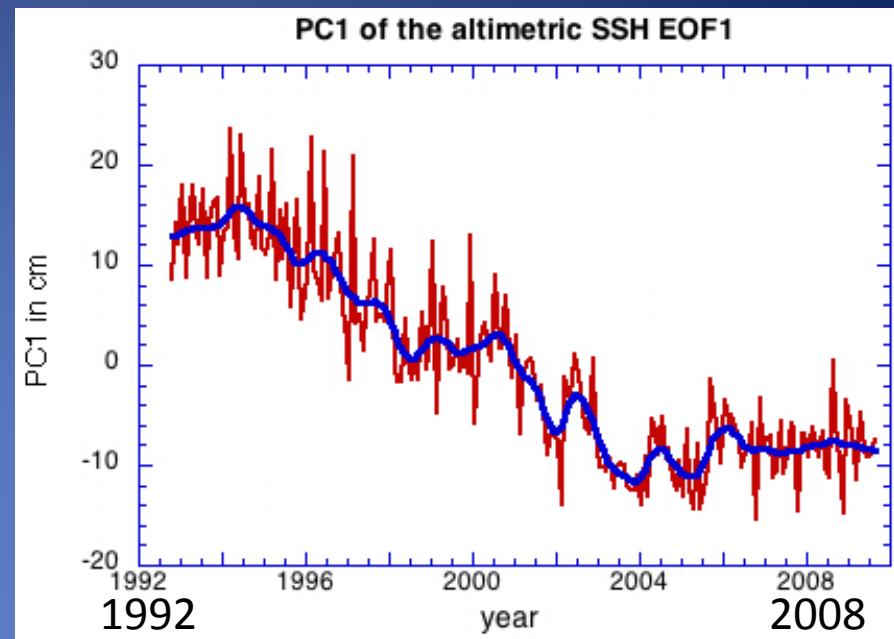
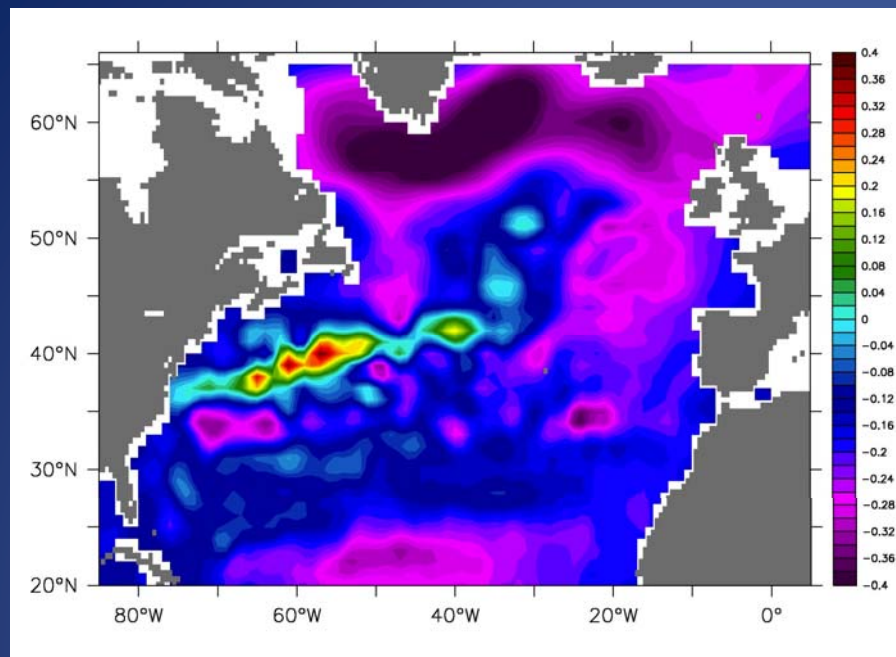


Wallace & Smoliak 2010 AMV: COWL or MOC dynamics?



Wood, Overland, Jonnson & Smoliak
GRL 2010: subarctic air temperature
index regressed on SST





■ **WEAKENING and SHRINKING SUBPOLAR GYRE - FROM ALTIMETRY** (*Hakkinen & Rhines Science 2004*) -> **WESTWARD SHIFT OF THE SUBPOLAR FRONT** (*Bersch 2002; Hatun et al., Science 2005*)

■ **INCREASING SALINITY AND TEMPERATURE OF THE ATLANTIC FLOW** in the channels leading to the Nordic Seas since 1996 with a steep increase since 2002 (Holliday et al 2008)

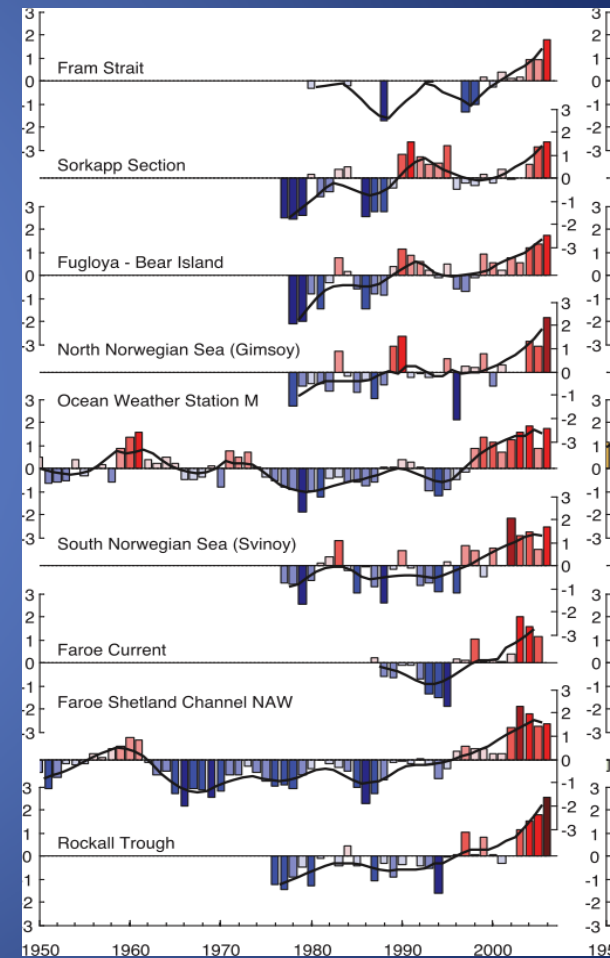
Cannot be explained by changes in surface heat and P-E fluxes

■ **EXPANSION AND WEAKENING OF THE SUBTROPICAL GYRE** FROM SATELLITE ALTIMETRY AND SEAWIFS (*McClain et al 2004, Polvina et al. 2008*)

■ **NO SUBTROPICAL SURFACE DRIFTERS ENTERED SUBPOLAR GYRE BEFORE 2002** (*Brambilla & Talley, JPO 2006*)
BUT THEY DID SO AFTER 2001 (*Hakkinen & Rhines, JGR 2009*)

Holliday et al. GRL 2008

TEMPERATURE

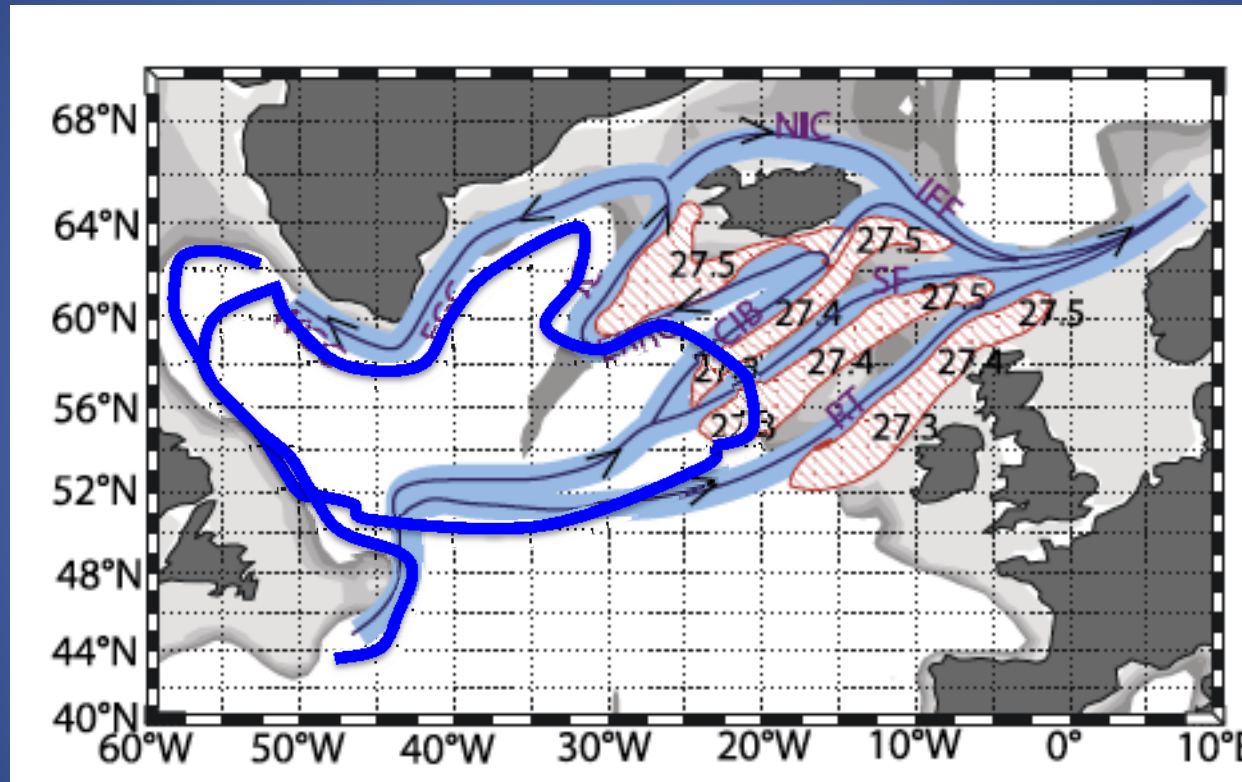


1950

2000

The cyclonic subpolar gyre competes with the overlying warm, saline northeastward transport to the Nordic Seas

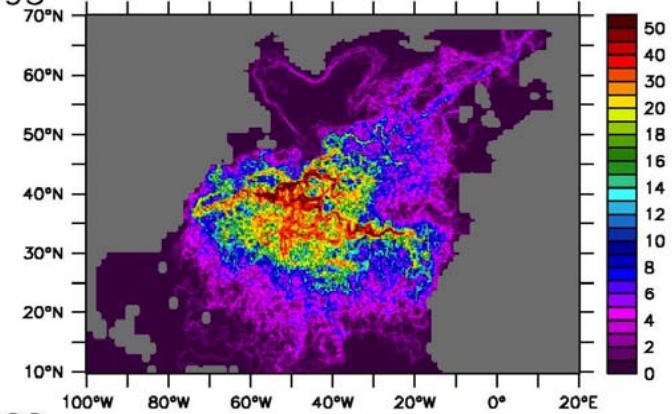
Brambilla & Talley, JGR 2008



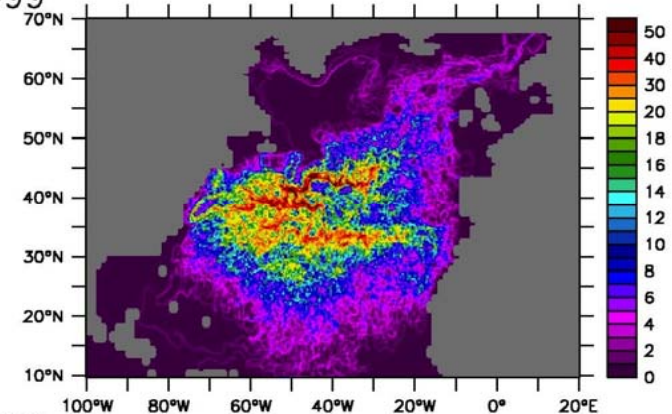
Along these paths air/sea interaction and lateral/diapycnal stirring creates Deep cooling (subpolar mode waters, SPMWs) and also mixing with warm eastern Atlantic waters

Synthetic surface drifters launched in AVISO/OSCAR altimetric surface currents the Gulf Stream region show episodes of warm, saline penetration northward, notable during the 2000s.

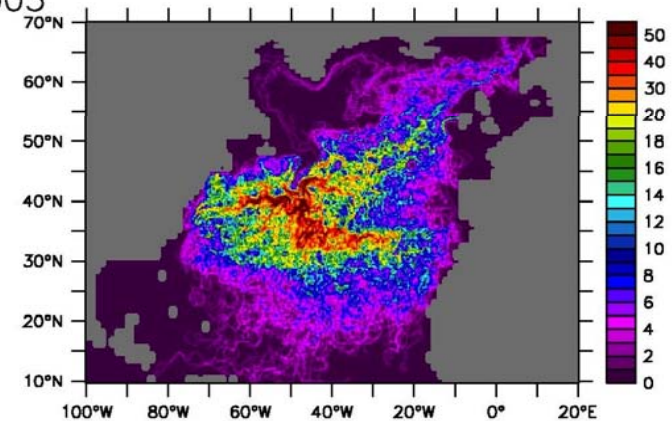
oct 1992 – dec 1995



oct 1996 – dec 1999

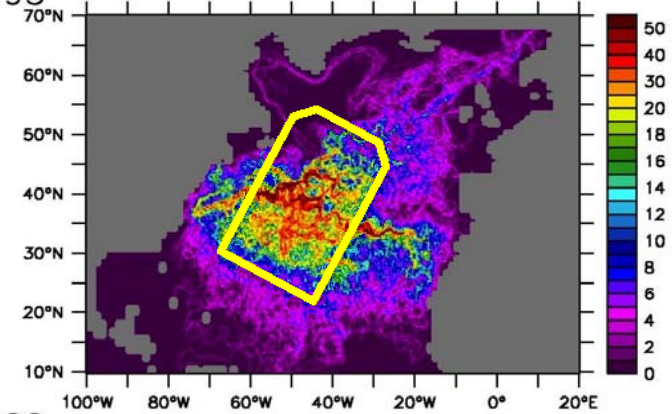


oct 2000 – dec 2003

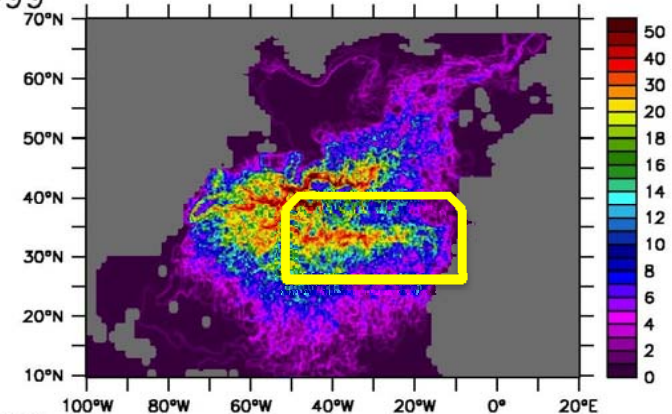


Synthetic surface drifters launched in AVISO/OSCAR altimetric surface currents the Gulf Stream region show episodes of warm, saline penetration northward, notable during the 2000s.

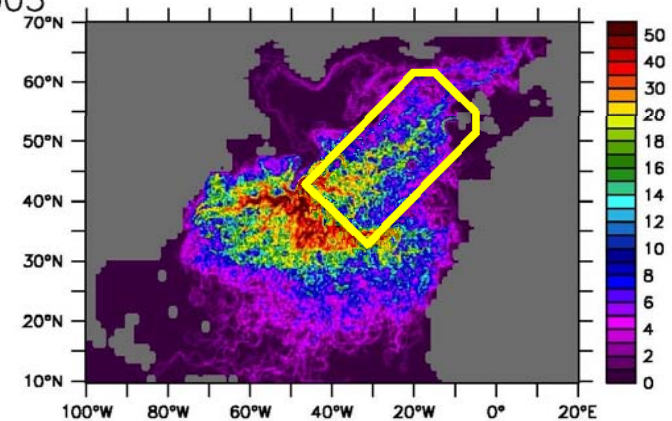
oct 1992 – dec 1995



oct 1996 – dec 1999



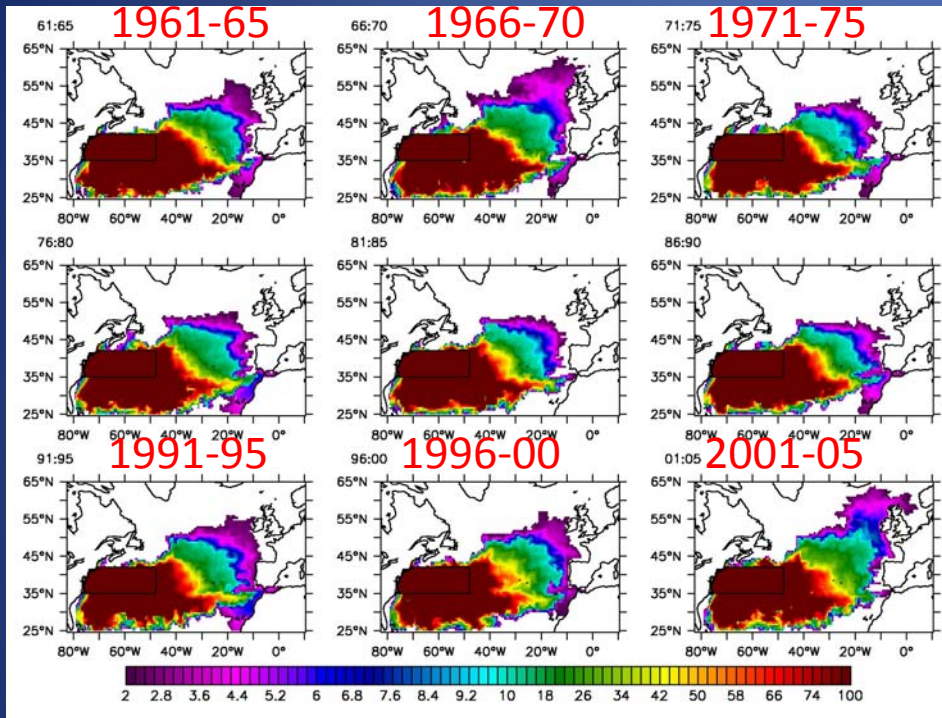
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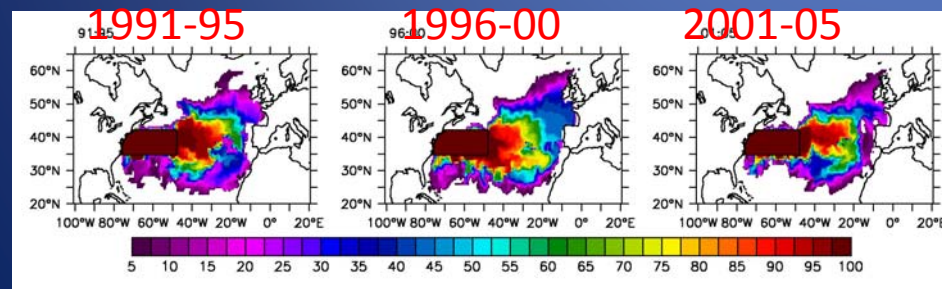
TRACER ADVECTION IN THE 15M 5-YEAR AVERAGE VELOCITY FIELD

SODA VELOCITY DATA:

TRACER KEPT CONSTANT IN THE GS BOX (left),



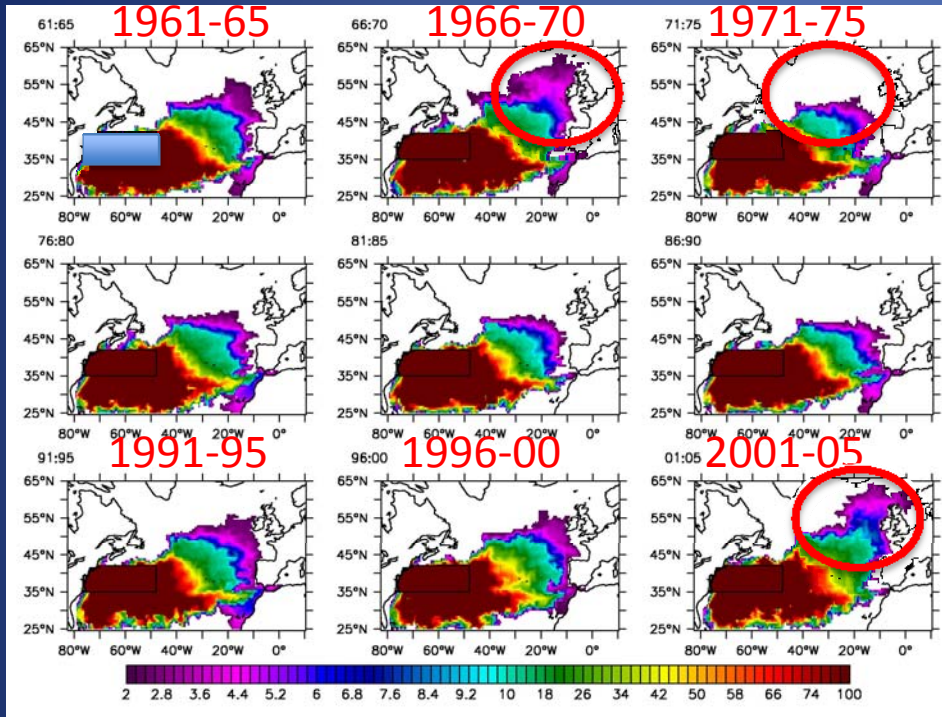
TRACER IN THE SURFACE DRIFTER VELOCITY FIELD



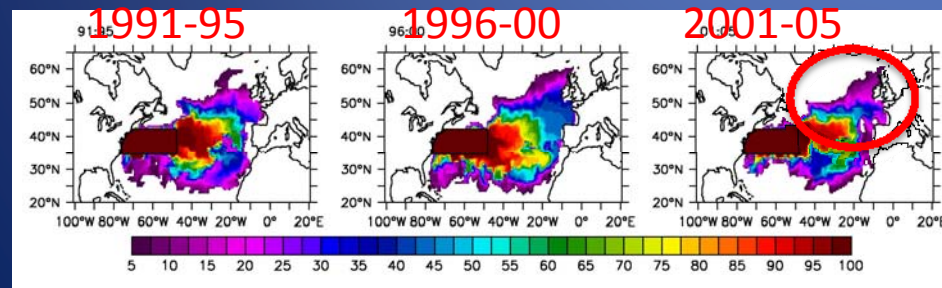
TRACER ADVECTION IN THE 15M 5-YEAR AVERAGE VELOCITY FIELD

SODA VELOCITY DATA:

TRACER KEPT CONSTANT IN THE GS BOX (left),



TRACER IN THE SURFACE DRIFTER VELOCITY FIELD

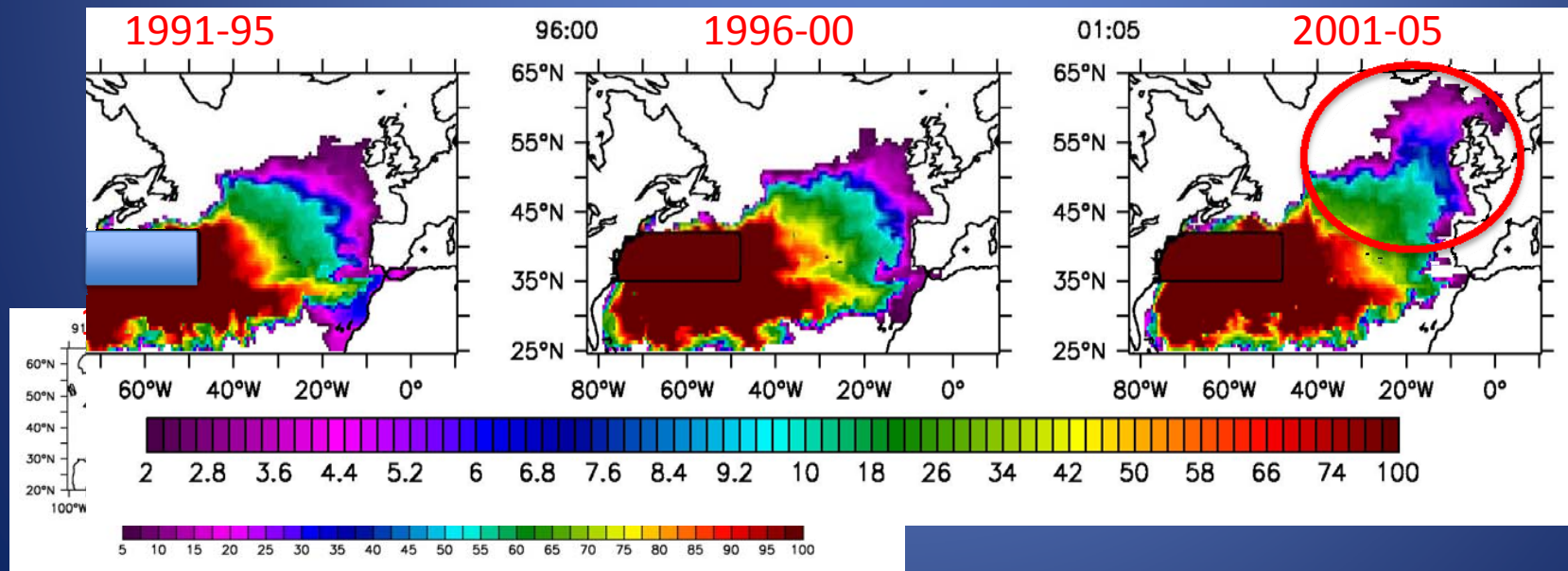


TRACER ADVECTION IN THE 15M 5-YEAR AVERAGE VELOCITY FIELD

SODA VELOCITY DATA:

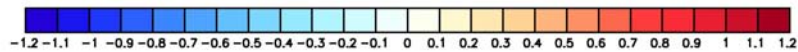
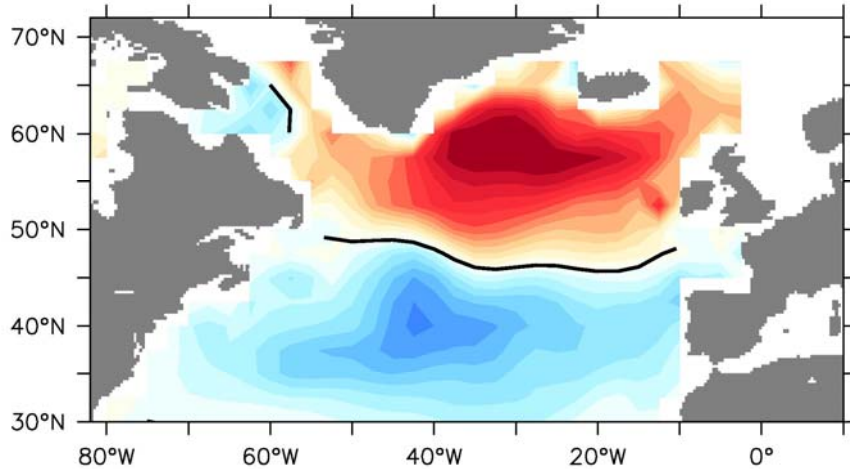
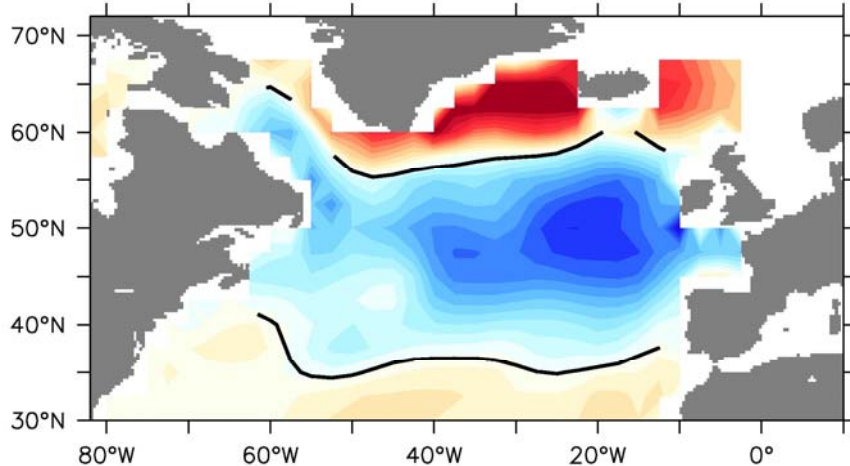
TRACER KEPT CONSTANT IN THE GS BOX (left),

TRACER IN THE SURFACE DRIFTER VELOCITY FIELD



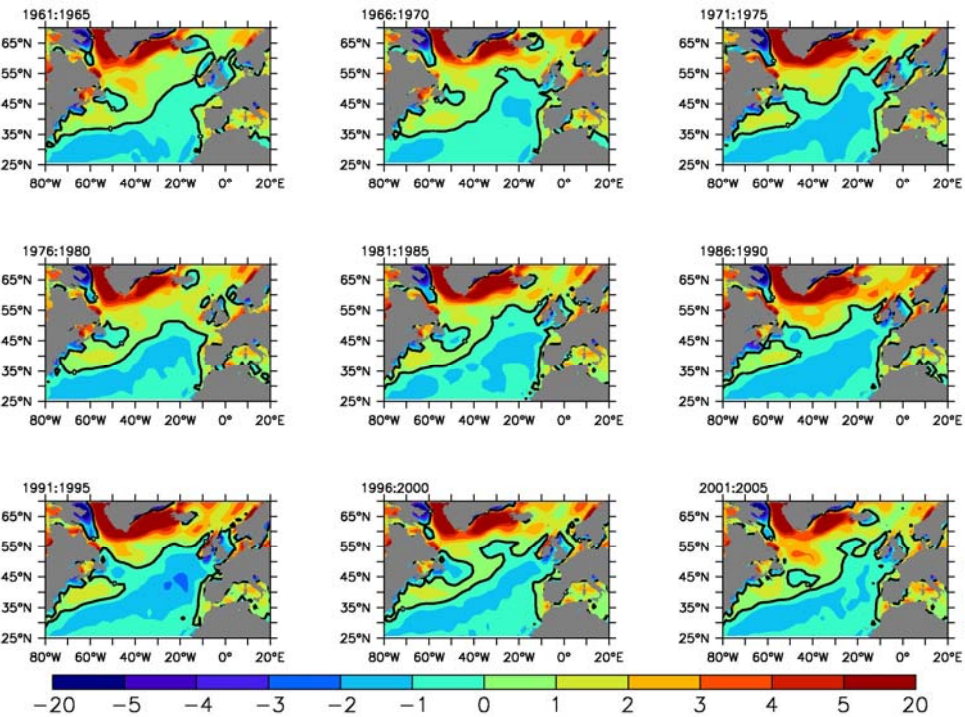
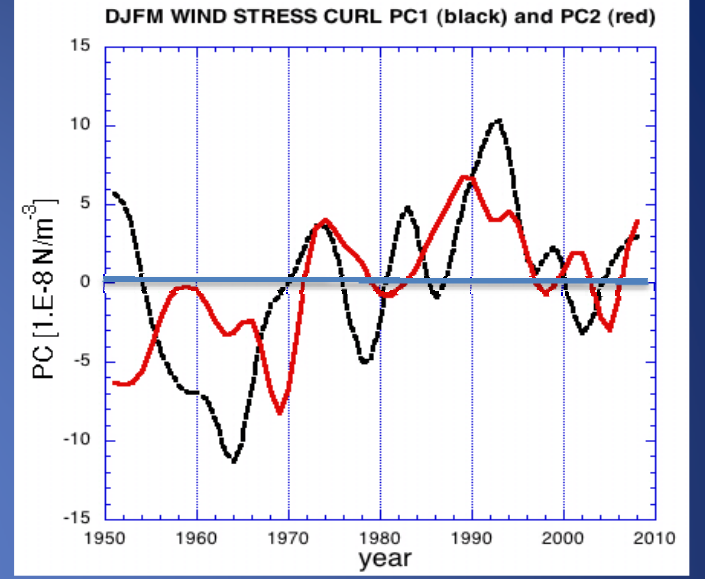
DJFM WIND STRESS CURL VARIABILITY

EOF1 (24%) and EOF2 (16%)

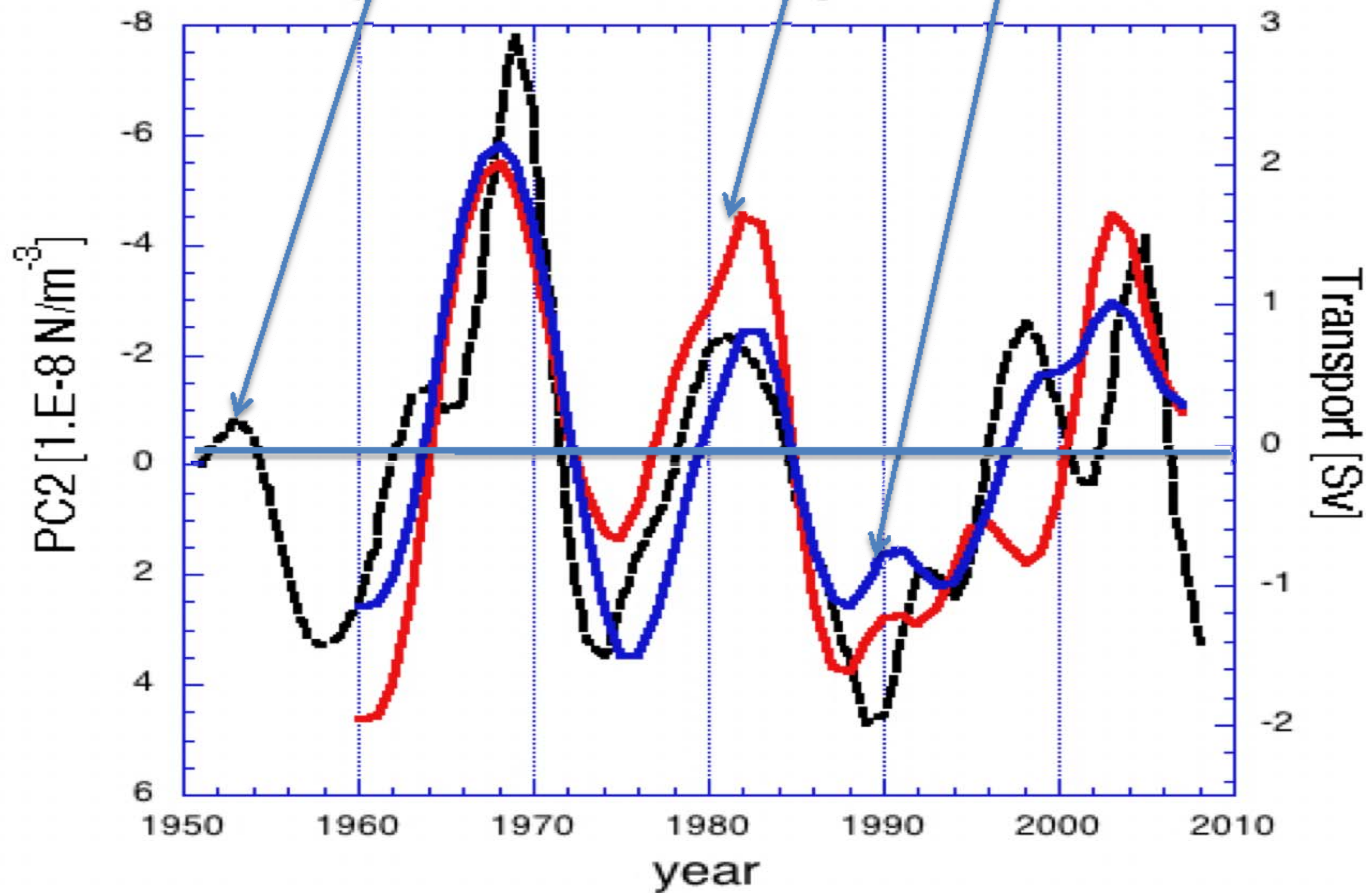


PCs (right)

DJFM CURL
FIELDS
FROM NCEP
(below)

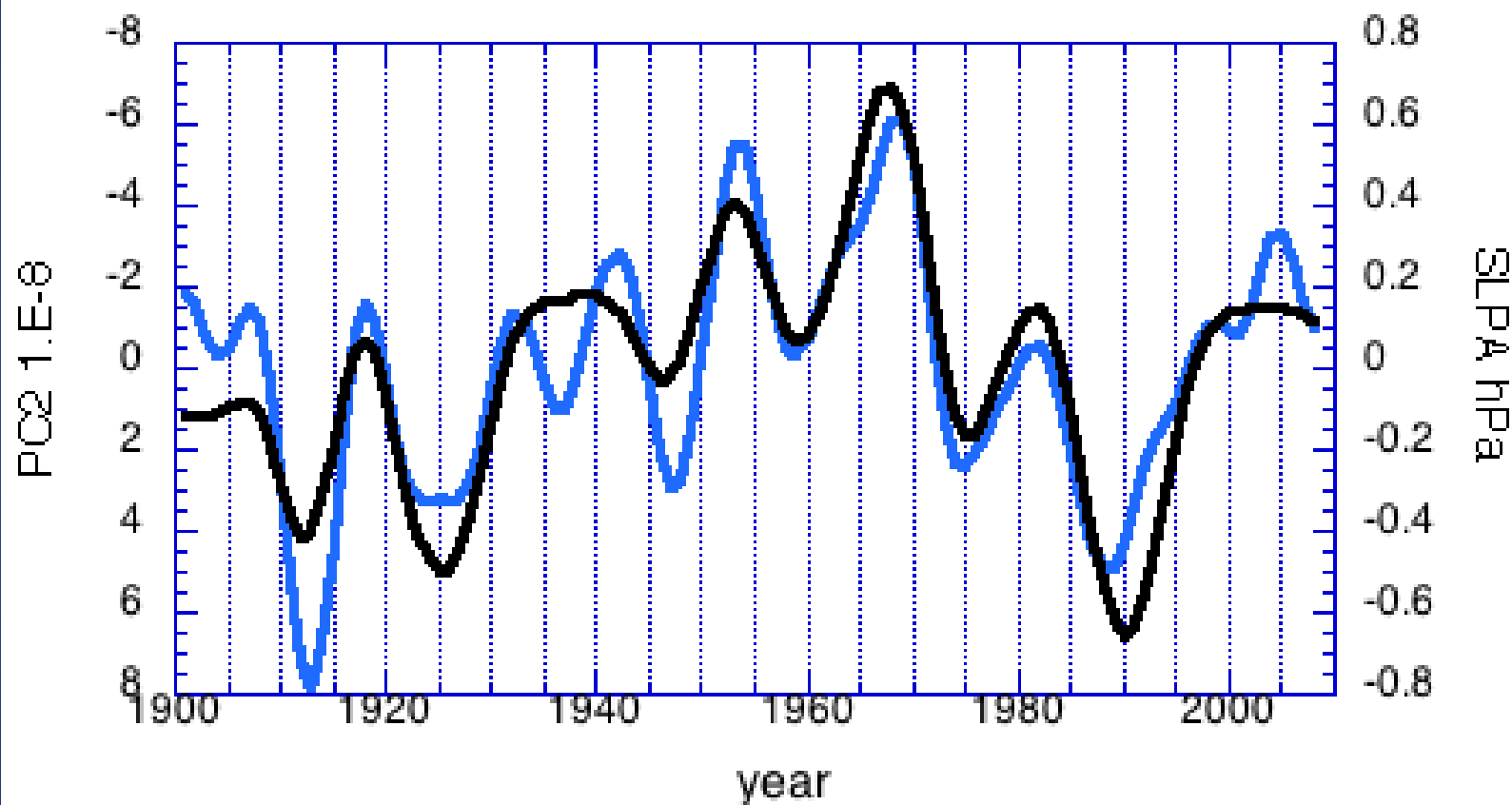


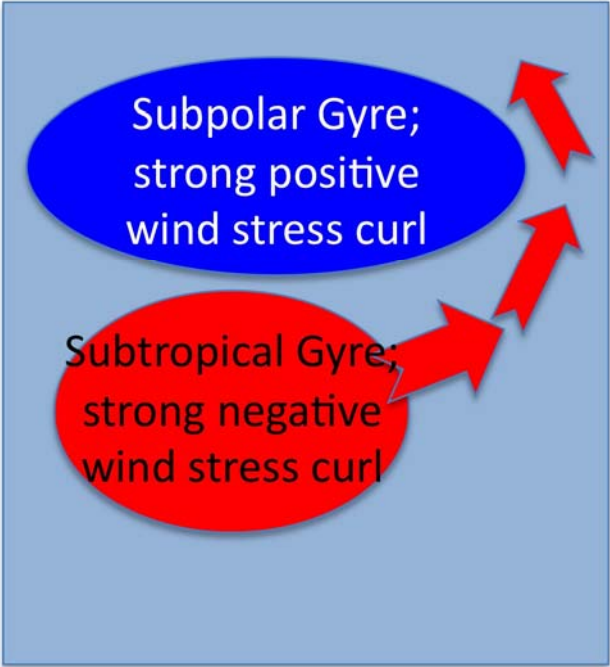
PC2 of DJFM WIND STRESS CURL (quadratic trend removed)(black)
TRANSPORT OF WATERS S>35.3 (red), S>35.4(blue) (detrended)
All quantities smoothed by 5 binomial filters



MAXIMUM TRANSPORT WHEN SUBTROPICAL AND SUBPOLAR
CURL ANOMALIES ARE ASSOCIATED WITH THE WEAKENED GYRES

Subpolar SLPA (black) and PC2 of curl (blue)
(from Compo 100yr analysis)



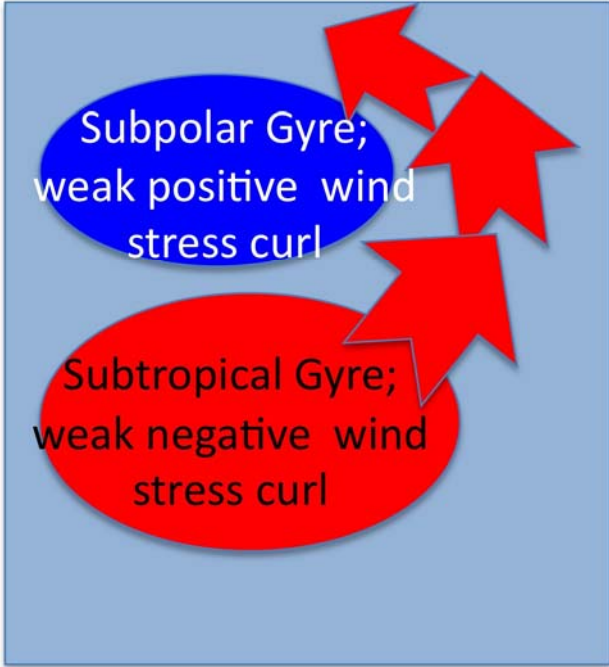


Subpolar Gyre;
strong positive
wind stress curl

Subtropical Gyre;
strong negative
wind stress curl

This diagram shows two gyres on a light blue background. The top gyre is a blue oval with white text, and the bottom gyre is a red oval with black text. Red arrows indicate a clockwise flow from the top gyre to the right, then down, then left, and finally up back to the top gyre. The arrows are larger and more numerous, indicating a strong circulation.

Subtropical Gyre;
strong negative
wind stress curl



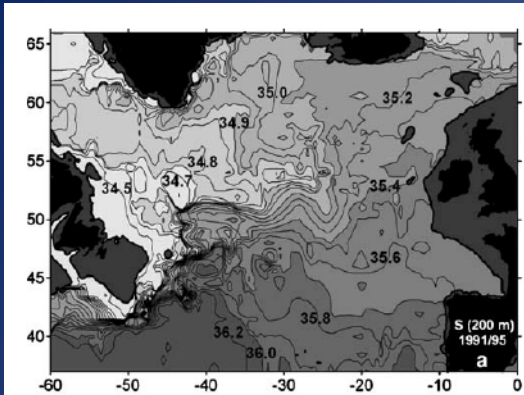
Subpolar Gyre;
weak positive wind
stress curl

Subtropical Gyre;
weak negative wind
stress curl

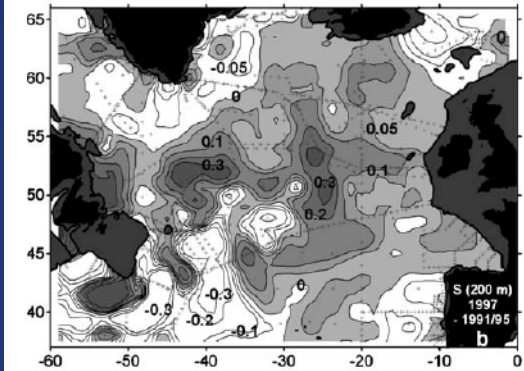
This diagram shows two gyres on a light blue background. The top gyre is a blue oval with white text, and the bottom gyre is a red oval with black text. Red arrows indicate a clockwise flow from the top gyre to the right, then down, then left, and finally up back to the top gyre. The arrows are smaller and fewer in number compared to the left diagram, indicating a weak circulation.

Subtropical Gyre;
weak negative wind
stress curl

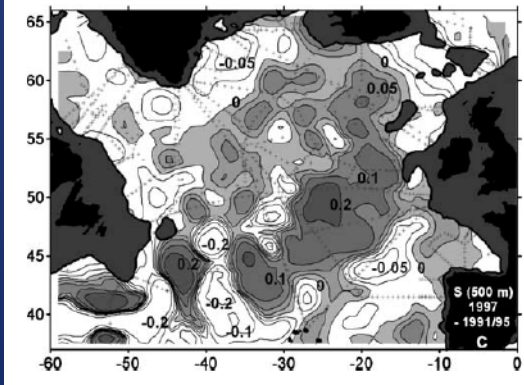
The warm, saline signal extends to great depth (>500m)



S 200m
1991-95



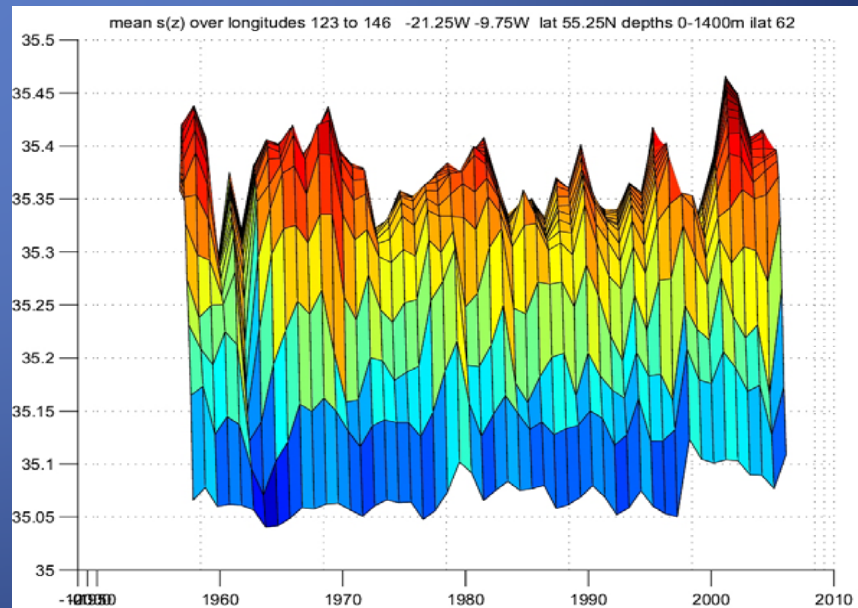
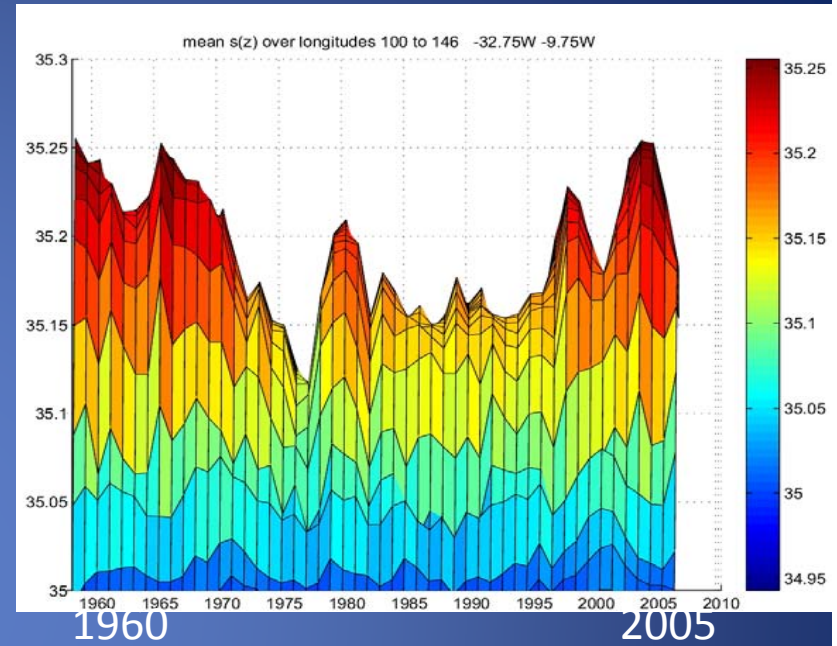
S 200m
1997-(1991-95)



S 500m
1997-
(1991-95)

60N

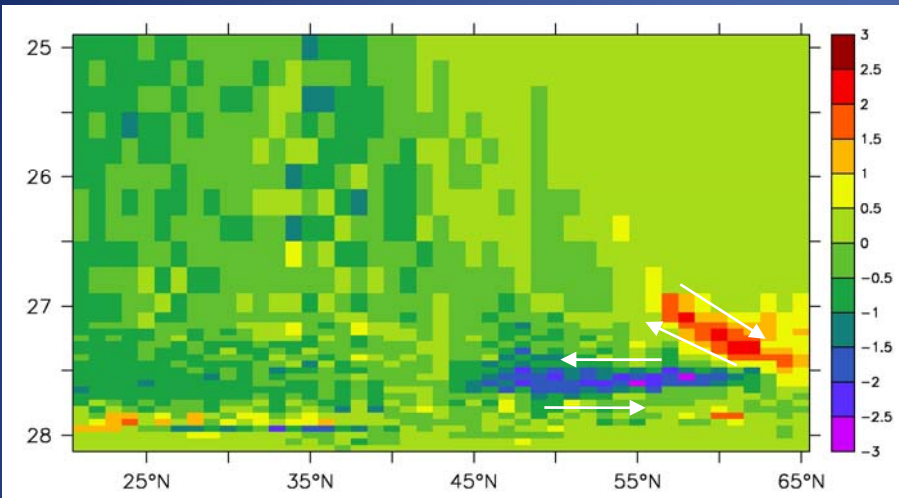
55N



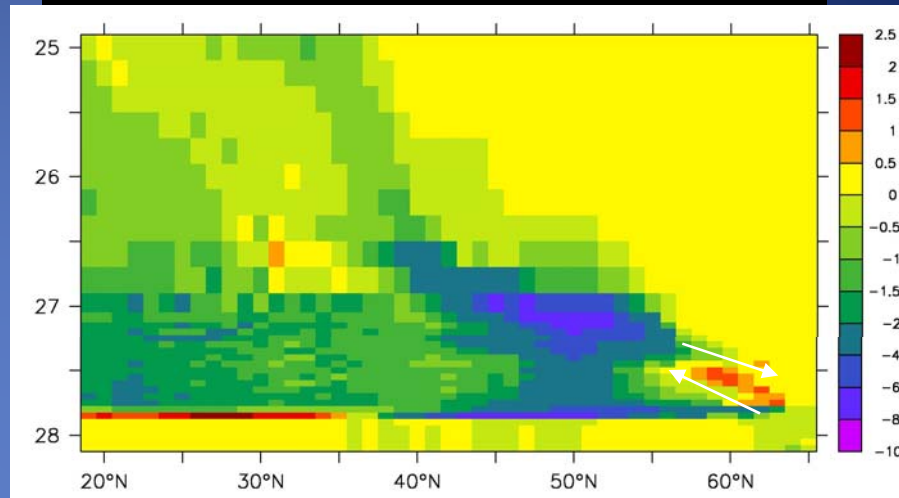
Bersch, Yashayaev,
Koltermann OD2007

AMOC CHANGE IN DENSITY SPACE [2001-2007] - [1994-2000]

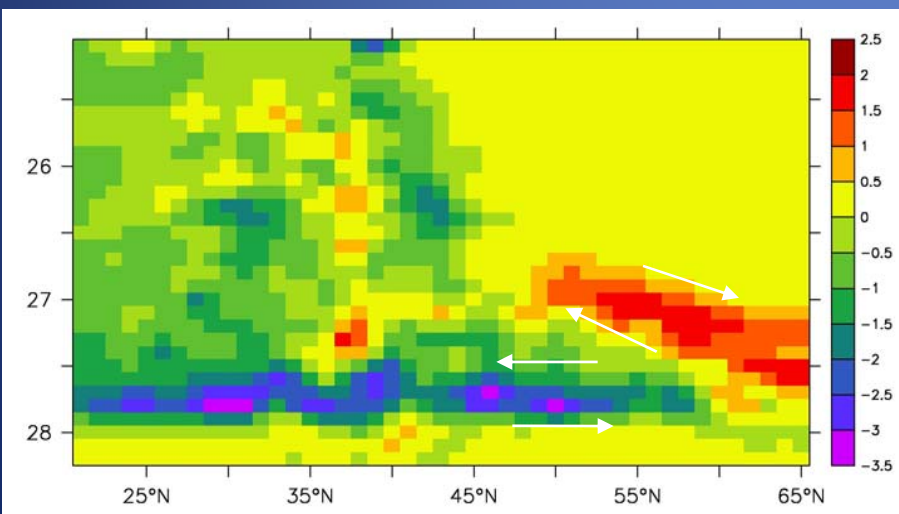
PROGNOSTIC POM



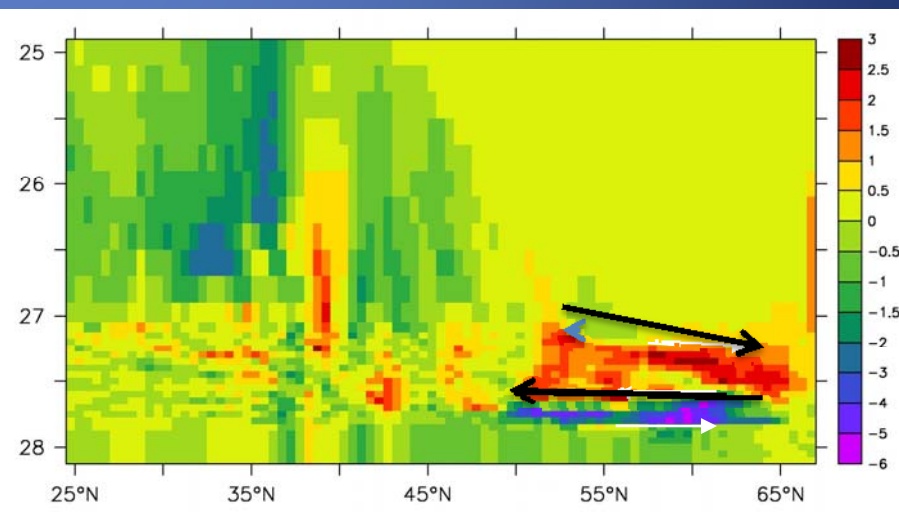
STATE ESTIMATION- ECCO-KF



PROGNOSTIC POM-Z



STATE ESTIMATION- SODA



Units are Sv

the basic dynamics of intergyre exchange appears in a 2-layer isopycnal model: cold water outcrops in the north, forcing warm water to flow northward and join the subpolar gyre HIM model...

see *Huang & Flierl JPO 1987*

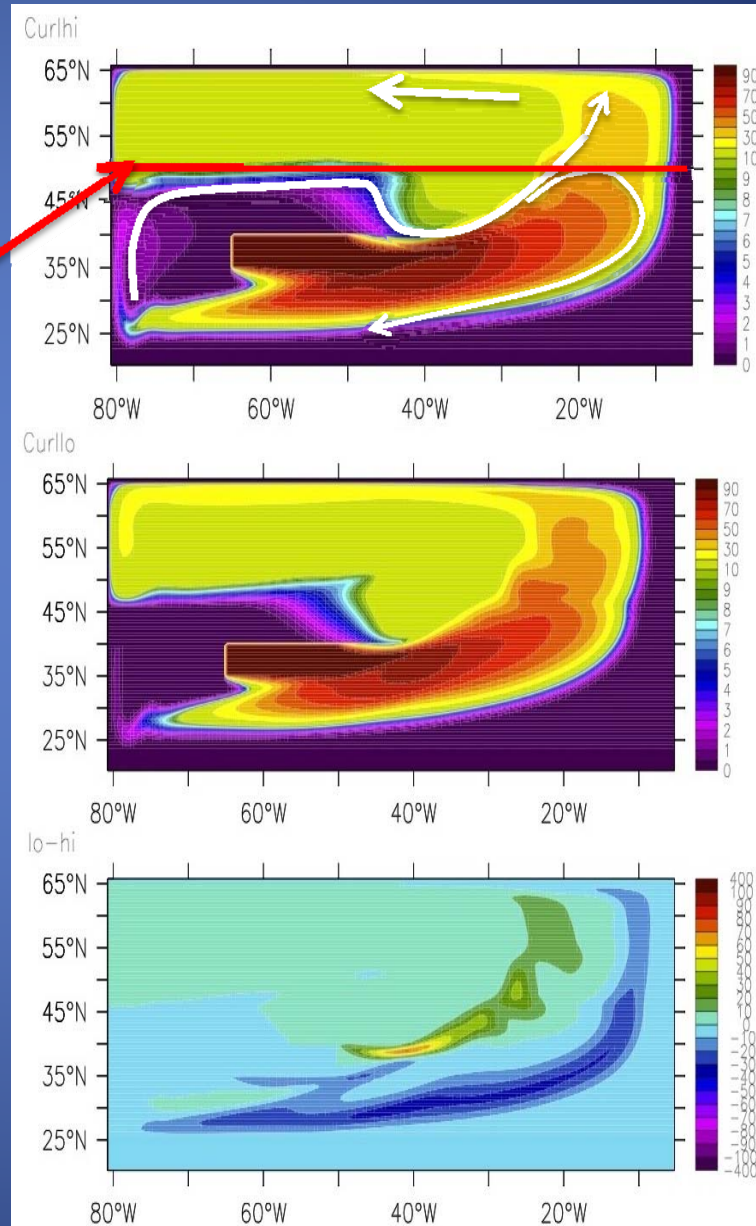
tracer released in subtropical gyre penetrates northward into the subpolar gyre even with strictly zonal wind-stress; outcropping of cold layer forces transport from subtropics to subpolar gyre

strong wind-curl

zero-curl line

weak wind-curl

difference



This is just the beginning of the story...of the web of interactions between global climate and ocean circulation in the high latitudes.

While the mean-state ocean has always been acknowledged as an active partner with the atmosphere, its role as active through *extratropical variability* is relatively new.

AMO, AMV, AMOC are 3 dimensional circulations, with complex subsurface structure. Example:

velocity transect across Irminger gyre: *Vage, Pickart, et al. DSR 2010;*
Knutsen, Rossby et al GRL 2005

AVISO altim

Nuka Arctica
adcp

GRACE DOT

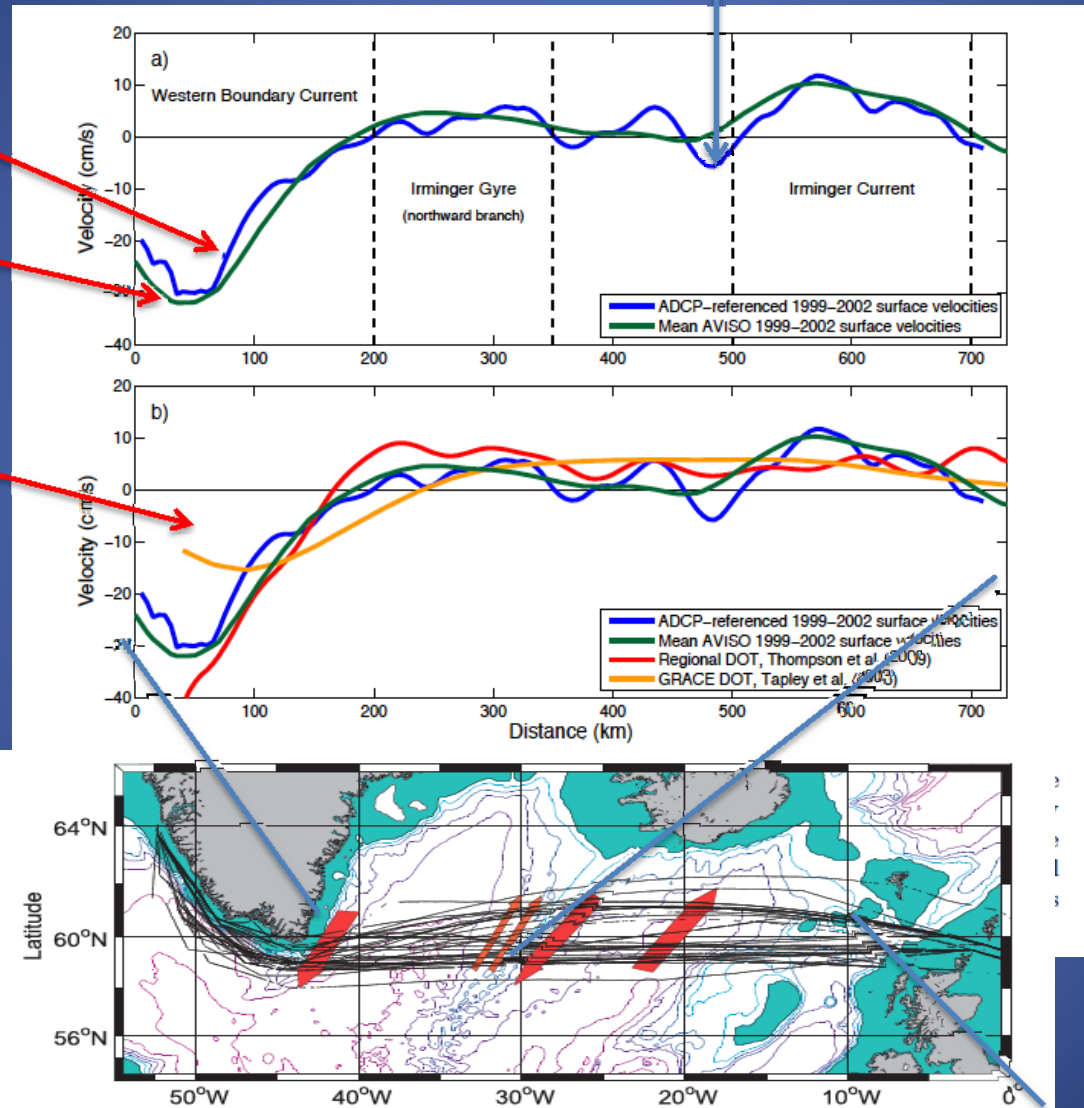


Figure 5: Nuka Arctica transects between 1999 and 2002, from Knutsen et al. (2005). The gre

In this example, the **dense overflows of the AMOC interact with** the warm, saline **surface circulation** through *barotropization* and turbulent entrainment. **The 1960s and 2000s invasion of warm subtropical water** to the subpolar gyre shows some continuity with the longer period Atlantic Multidecadal Variability (**AMV**)

Topographic control and pseudo-Taylor-columns are everywhere at high latitude (e.g., in the two-way circulation through the Charlie-Gibbs Fracture zone (*Bower & Van Appen, JPO 2008*)).

At high latitude the **altimetric sea-surface height** differs from the **baroclinic SSH derived from hydrography**: the ‘thermometer’ versus the ‘barotropic circulation gauge’. (And both differ from GRACE ocean mass variability (*Willis, Chambers & Nerem JGR 2008*))

In situ observations (e.g. *Vage et al. DSR 2010*) are beginning to show the mean and variable relationships between the **subpolar gyre strength** seen at the surface **and** key **meridional overturning** indices.

Meanwhile, new analyses of climate and circulation models are suggesting **covariance structure of subtropical/subpolar communication, structure of AMOC** and the construction of the shallow and deep AMOC branches (e.g. *Langehaug, Medhaug & Eldevik, 2011 preprint from Bergen Climate model*).

Summary

Penetration of subtropical waters northward to the subpolar gyre 'appears' as a **strengthened upper limb of the AMOC** after 2000 in the SODA model assimilation (has not happened since 1960s)

- unclear if the net upper layer transports have changed, clearly the sources of the North Atlantic Current have shifted towards more subtropical origin as confirmed by tracer simulations, satellite altimetry, sea-surface drifters, OSCAR surface currents (*Hakkinen & Rhines, JGR 2009*)

The forcing: The first two EOFs of wind stress curl correlated in broad terms with the **three recent warming-high salinity periods in the NE Atlantic (1960s, 1980, 2000s)**

-The first mode PC follows closely NAO-index, however the spatial pattern of the second mode resembles climatologic North-Atlantic SLP pattern

- The second mode appears as the modulation of the climatologic forcing pattern of the two gyres and hence dominates the gyre response such that weak gyres => strong warming of subpolar gyre.

Changes in the NAC source waters may not be of immediate significance to AMOC but can have a delayed effect in increasing water mass modification in the Nordic Seas as happened in the 1960s and early 1970s. These cycles appear to be part of the longer period AMO/AMV variability seen in recent SST historical analyses over 1900-2010 and longer.



