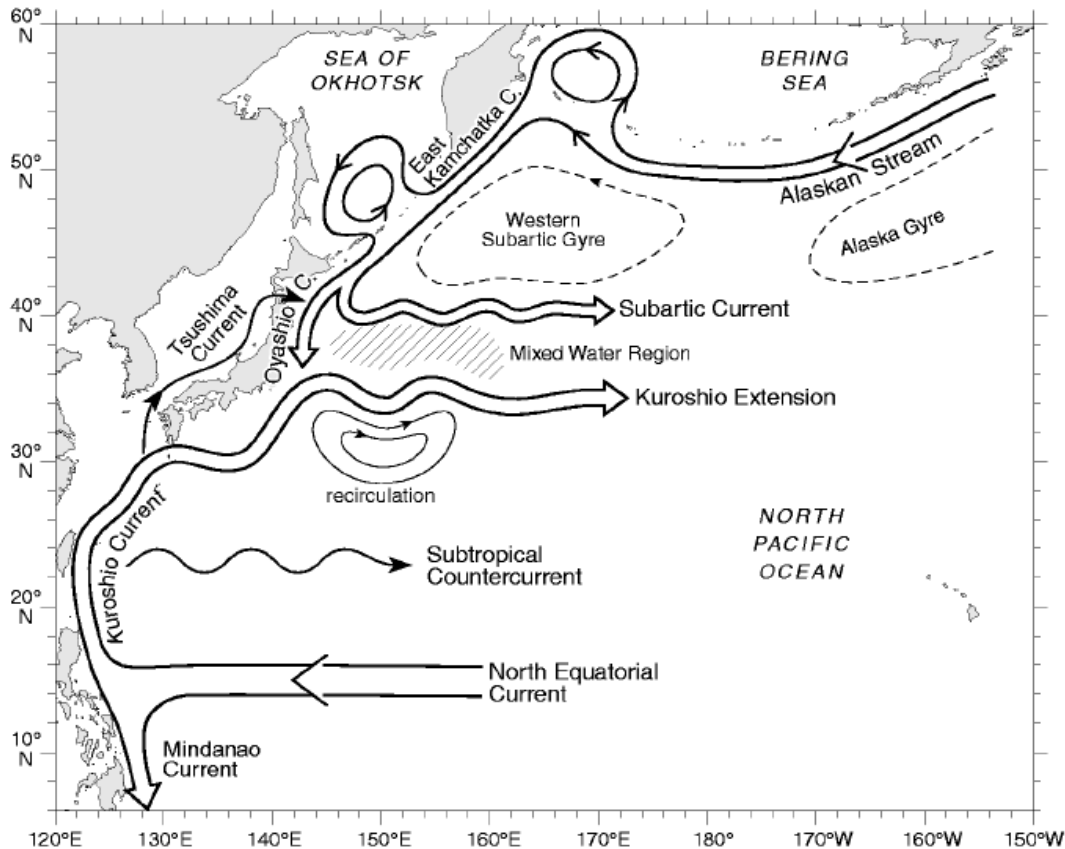


# Use of Satellite Altimeter Data for Decadal Predictions of the Kuroshio Extension Dynamic State

**B. Qiu<sup>1</sup>, S. Chen<sup>1</sup>, N. Schneider<sup>1</sup> and B. Taguchi<sup>2</sup>**

**1. Dept of Oceanography, University of Hawaii, USA**

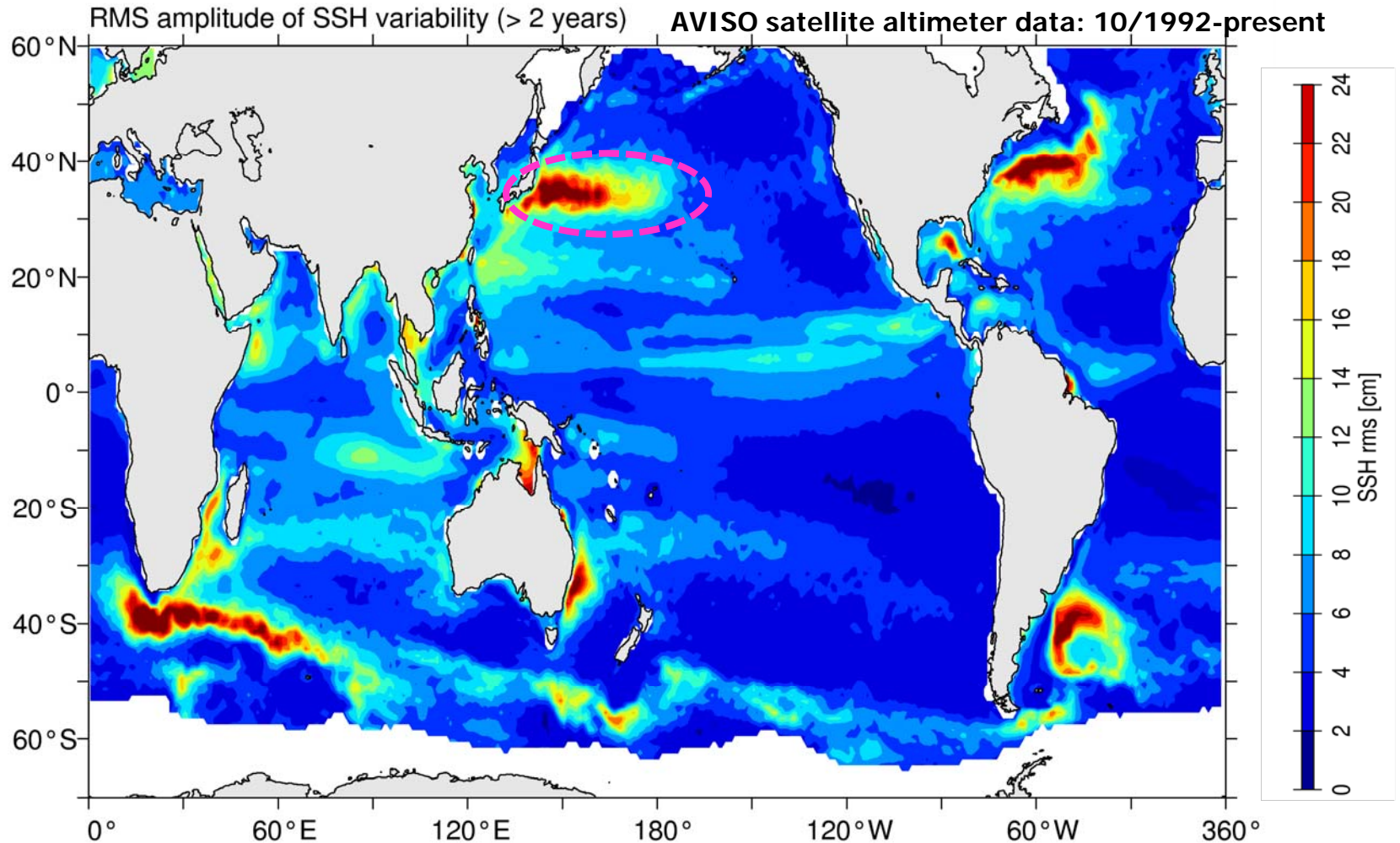
**2. Earth Simulator Center, JAMSTEC, Japan**



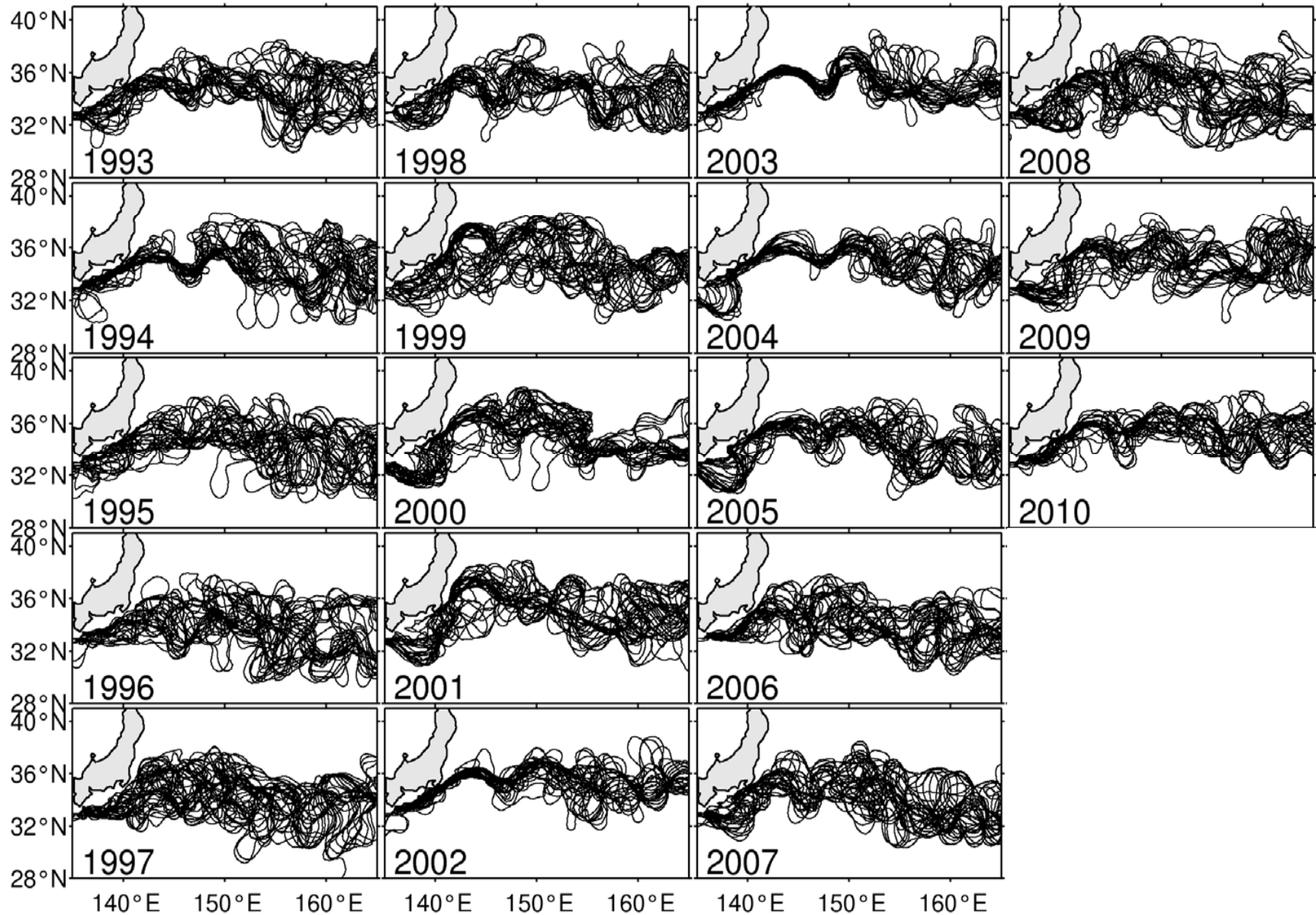
**Western North Pacific circulation schematic**

# Outlines

- Observed changes in the Kuroshio Extension system
- Decadal KE variability as a coupled ocean-atmos. mode
- Impact of forced vs coupled mode on prediction



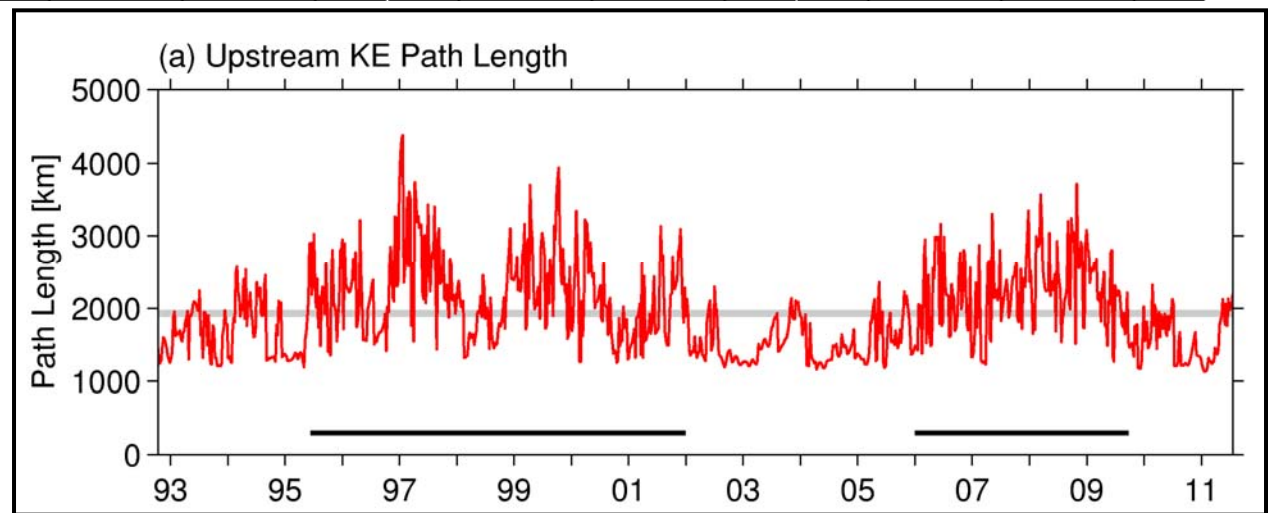
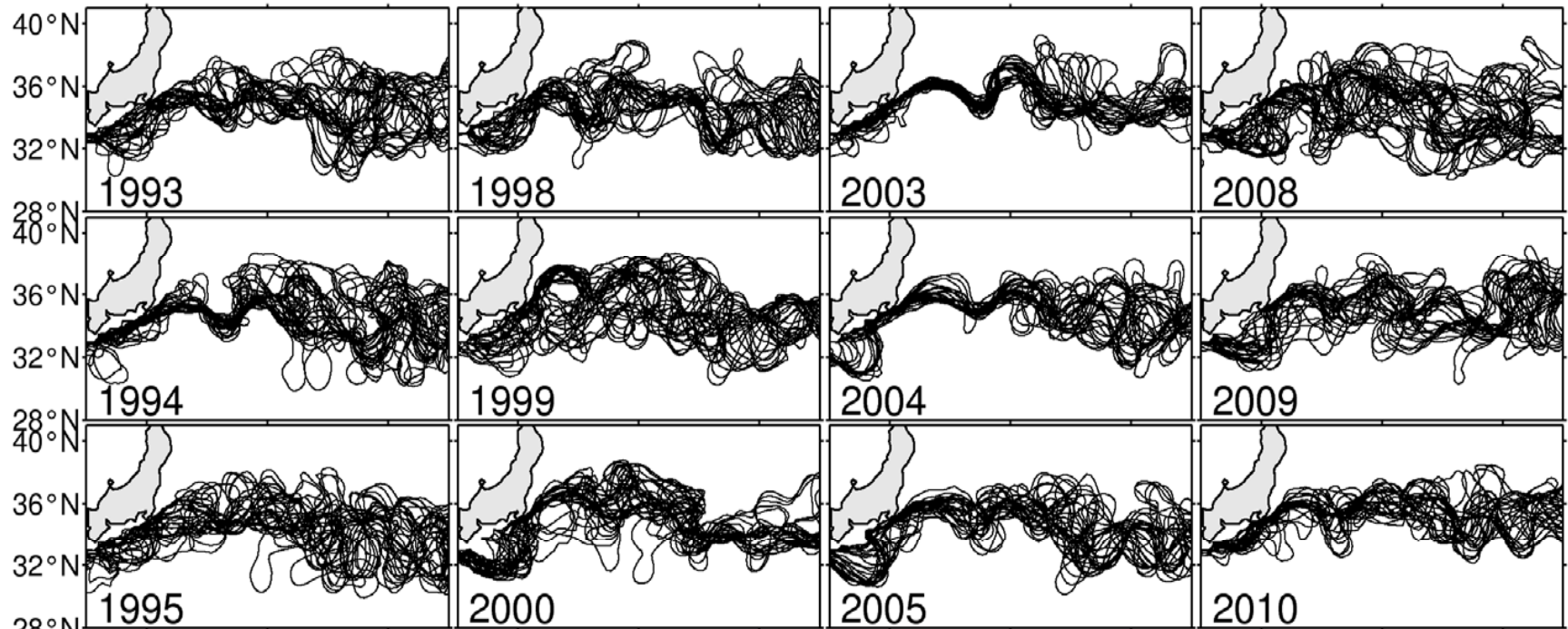
# Semi-monthly KE paths from altimeter-derived 1.7m SSH contours



Stable yrs: 1993-94, 2002-04, 2010-

Unstable yrs: 1996-2001, 2006-08

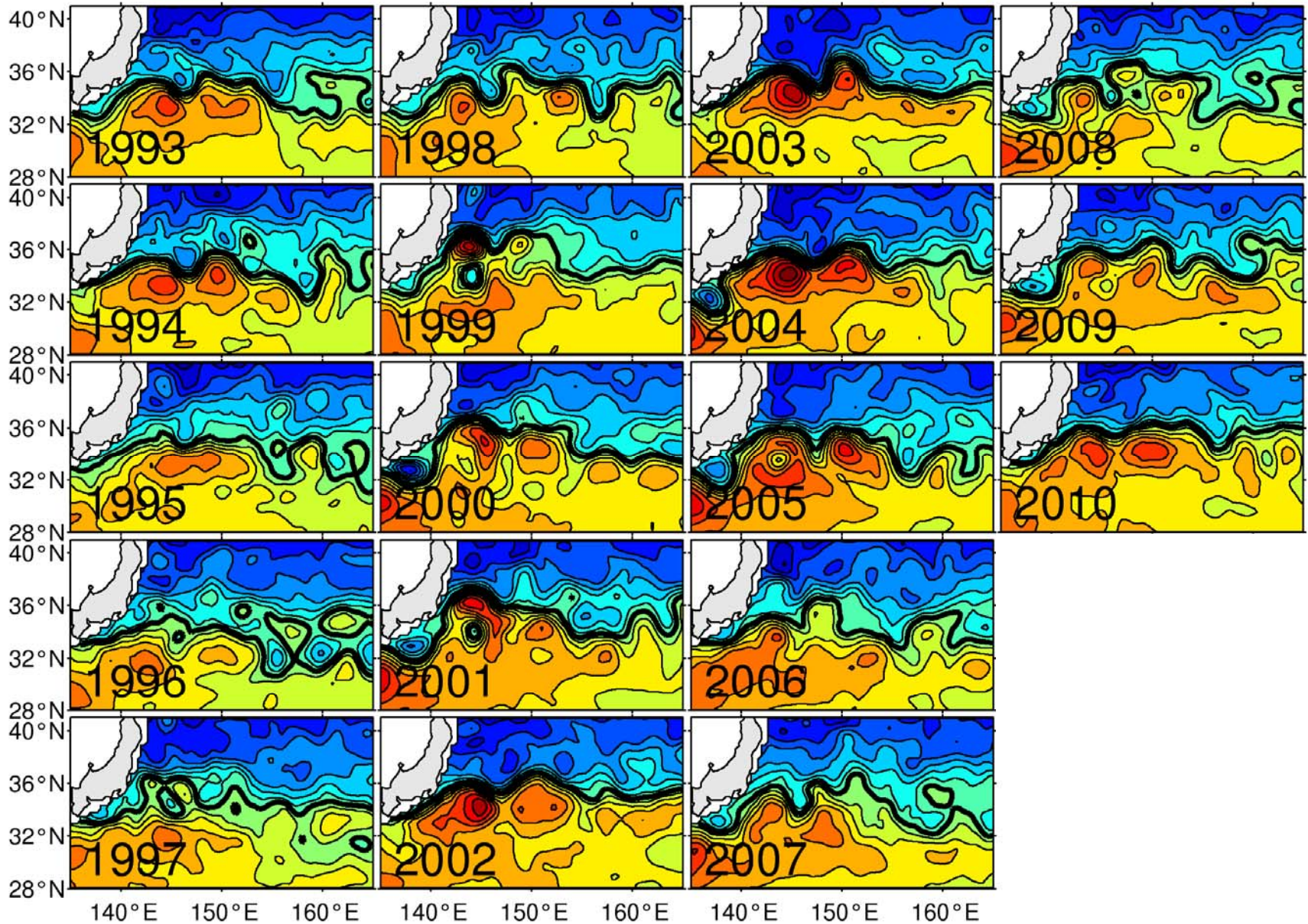
# Semi-monthly KE paths from altimeter-derived 1.7m SSH contours



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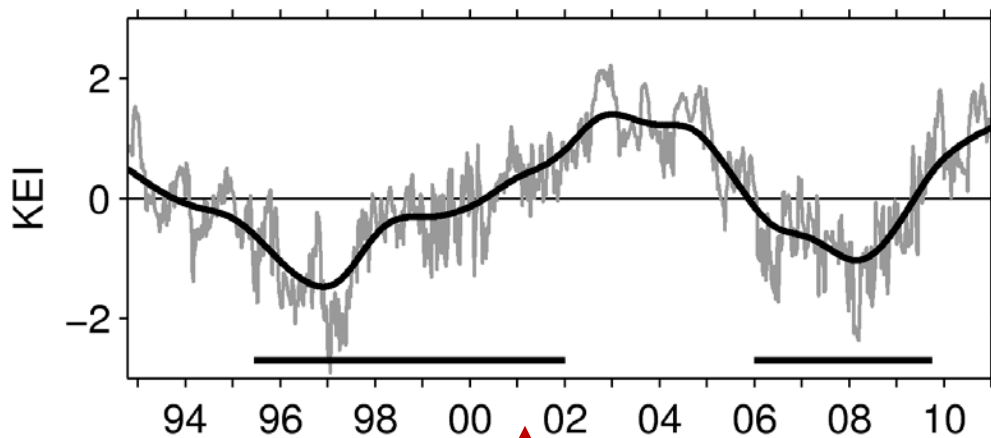
# Yearly-averaged SSH field in the region surrounding the KE system



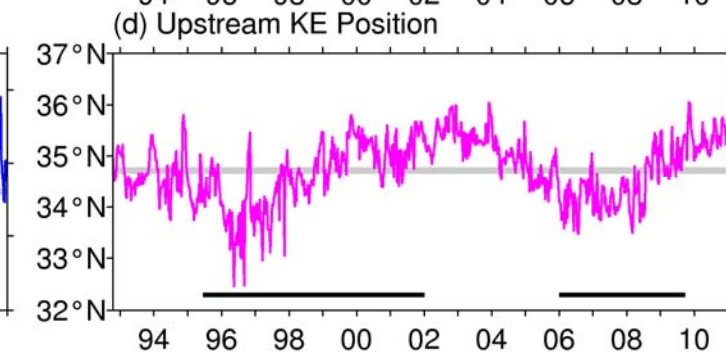
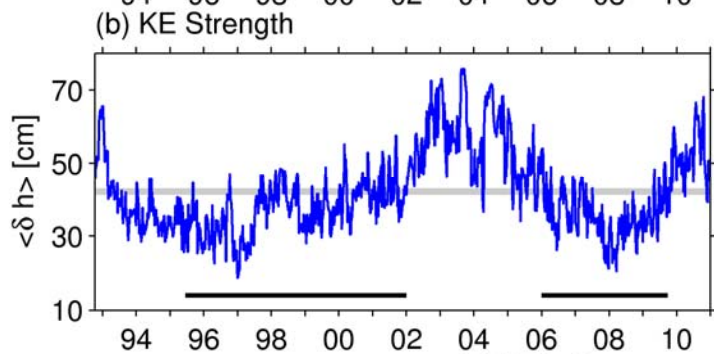
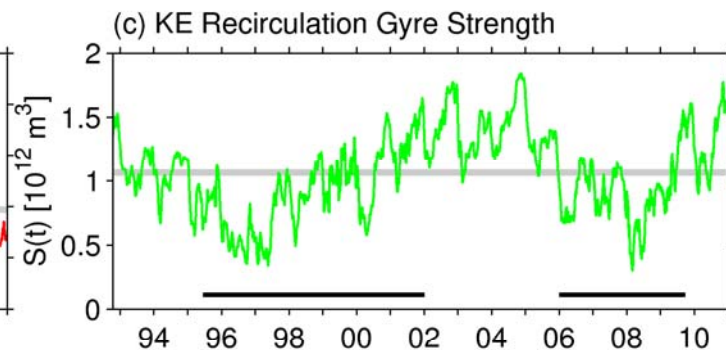
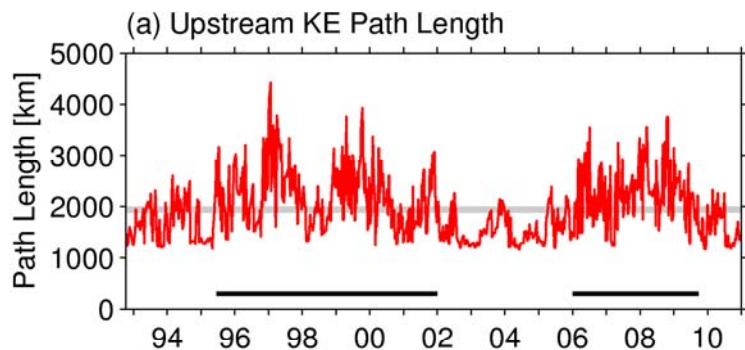
Stable yrs: 1993-94, 2002-04, 2010-

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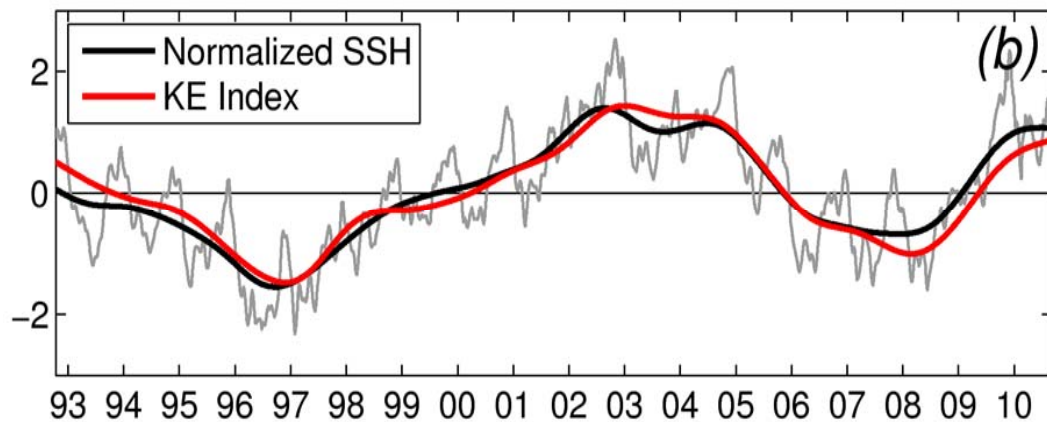
# A composite index quantifying the KE dynamic state: **the KE index**



**KE index:** average of the 4 dynamic properties (normalized)



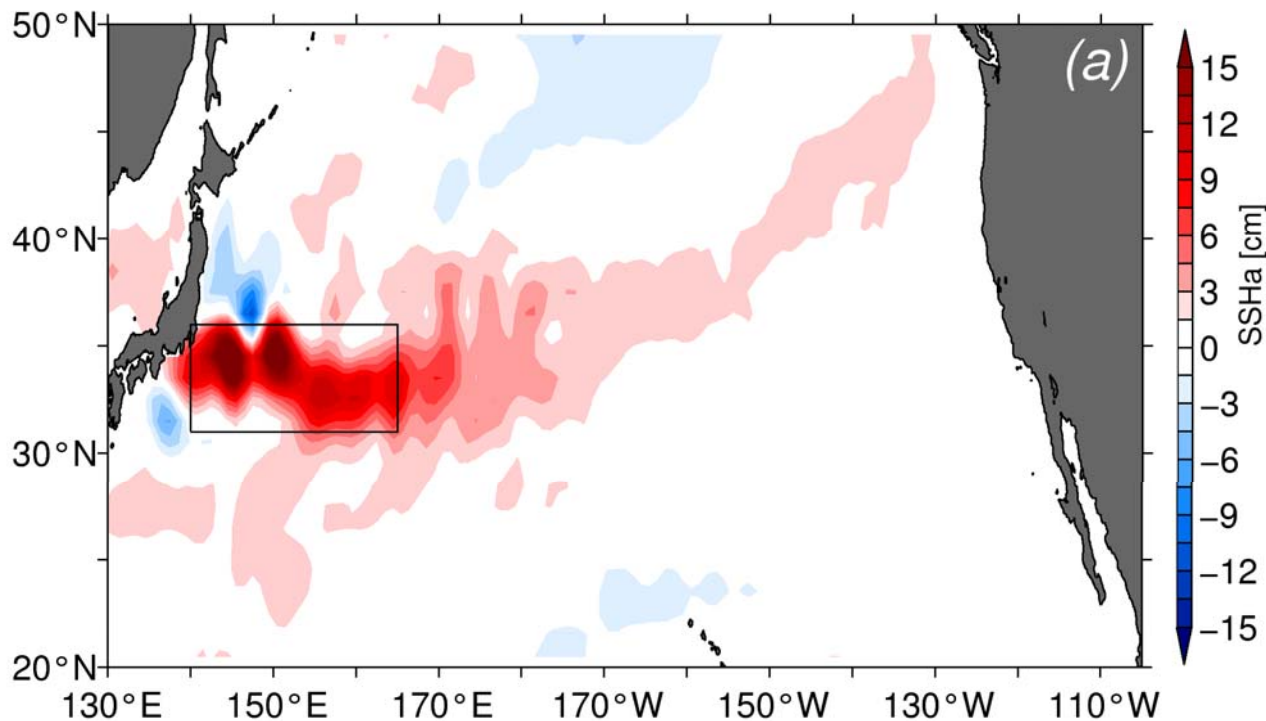
## Regression between the KE index and the AVISO SSH anomaly field



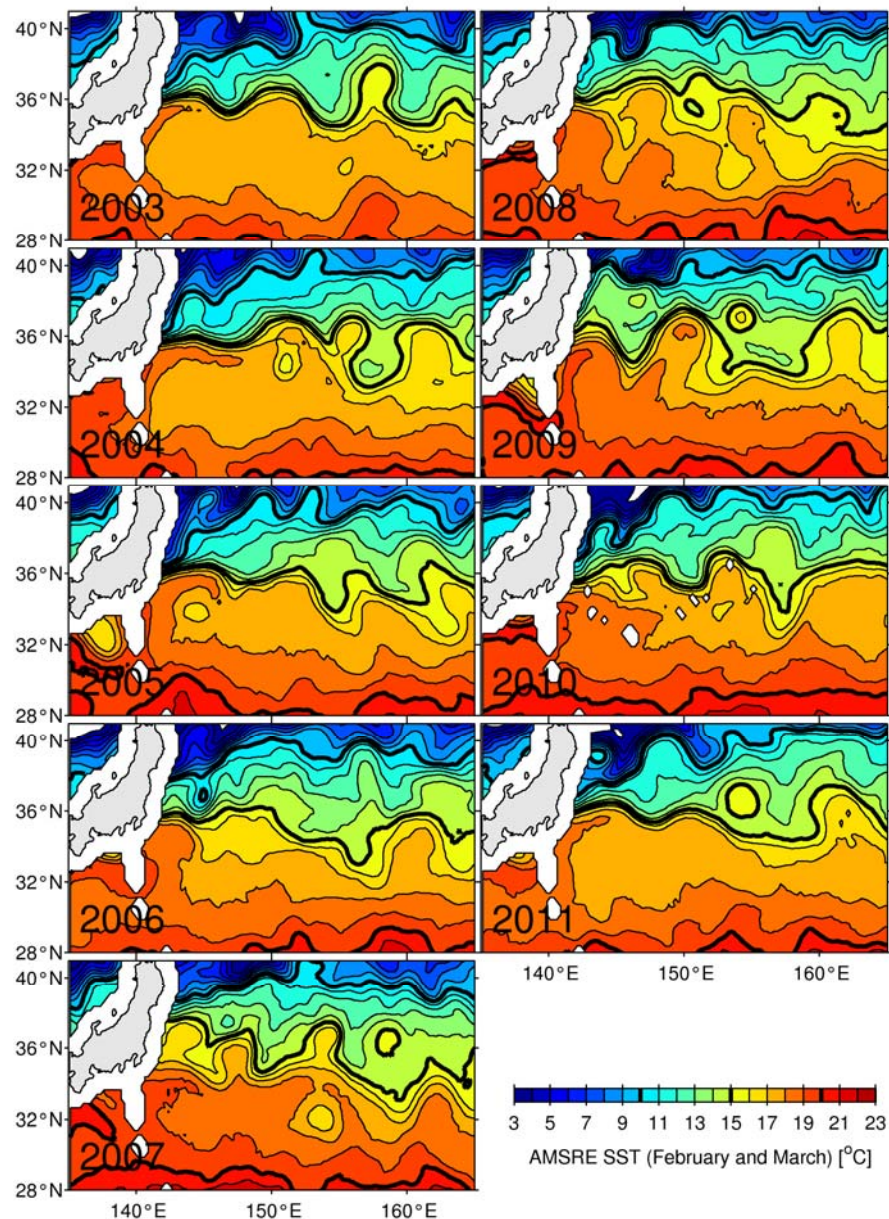
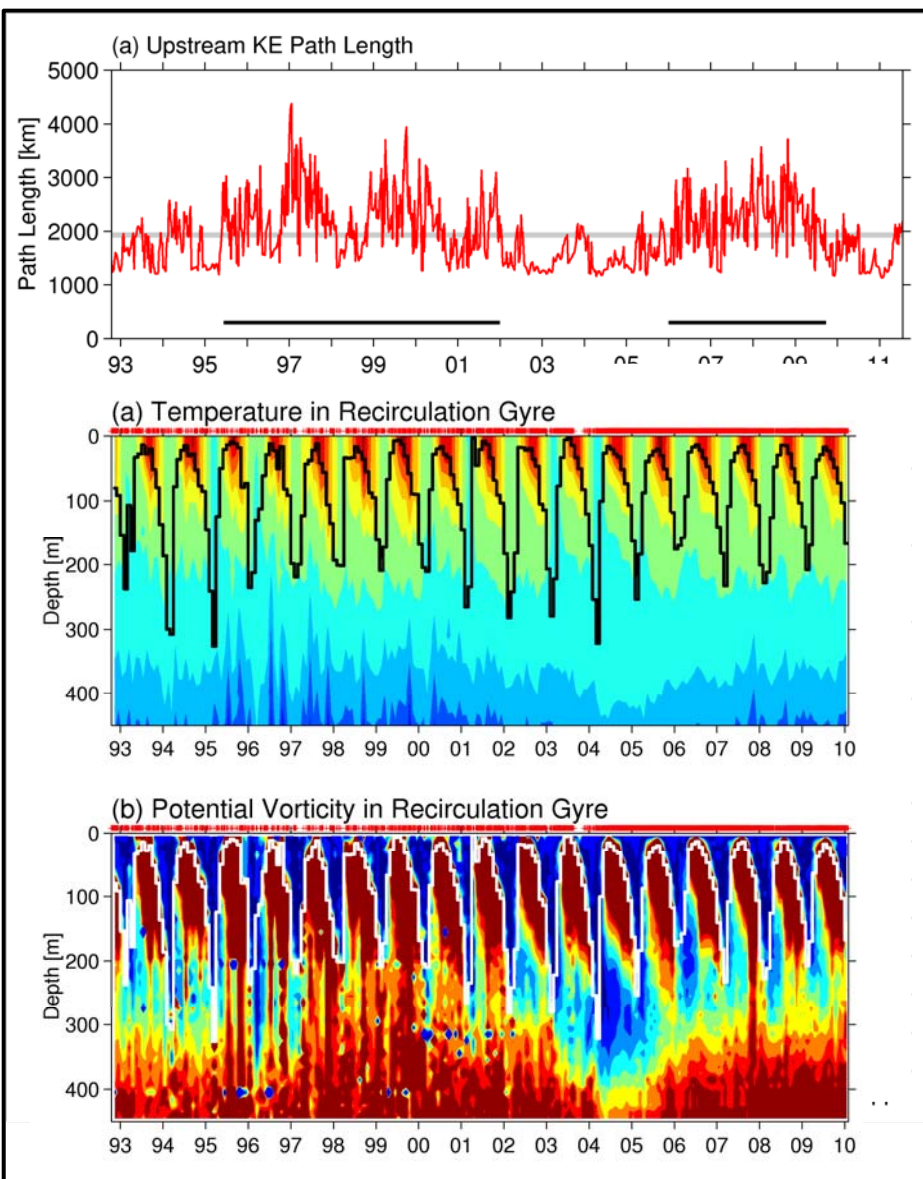
**KE index:** represented well by SSH anomalies in 31-36°N, 140-165°E

### Implications:

Predicting **KE index** becomes equivalent to predicting SSH anomalies in this key box



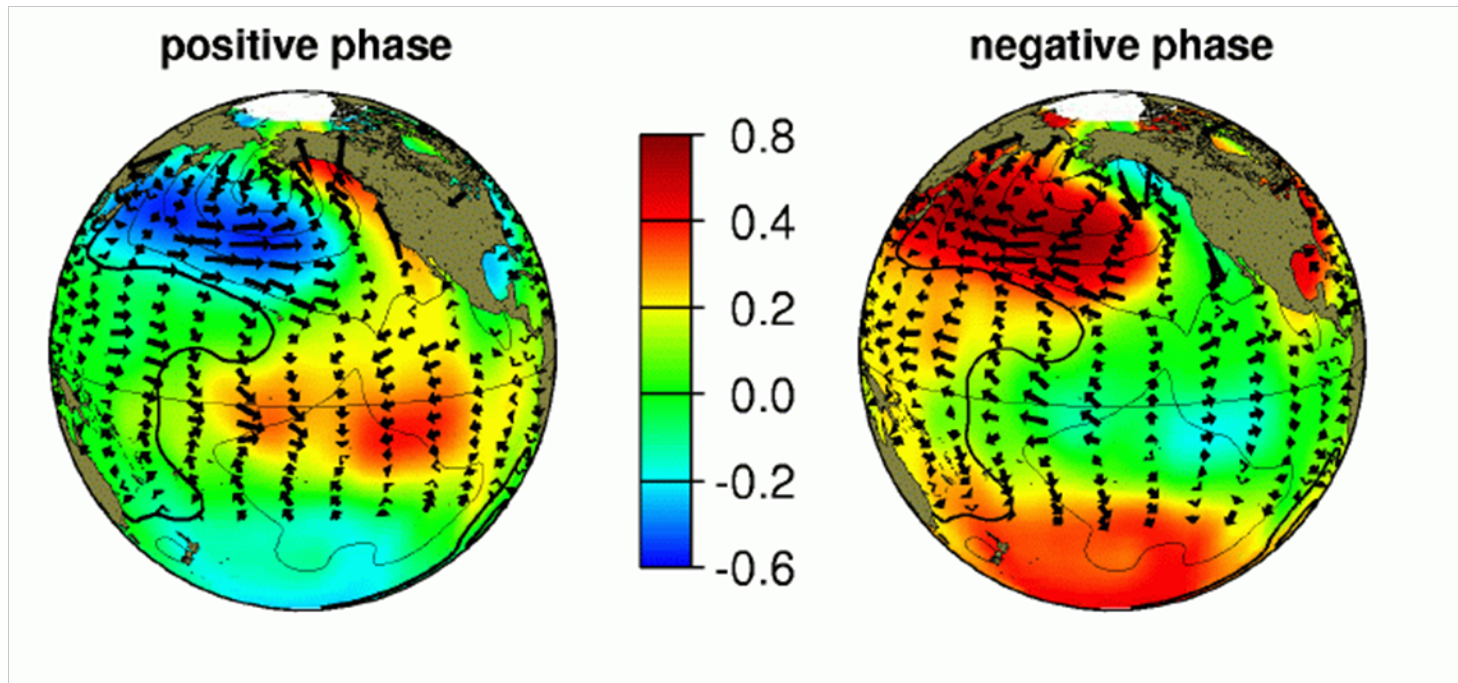
# KE dynamic state affects regional SST, MLD and mode water formation



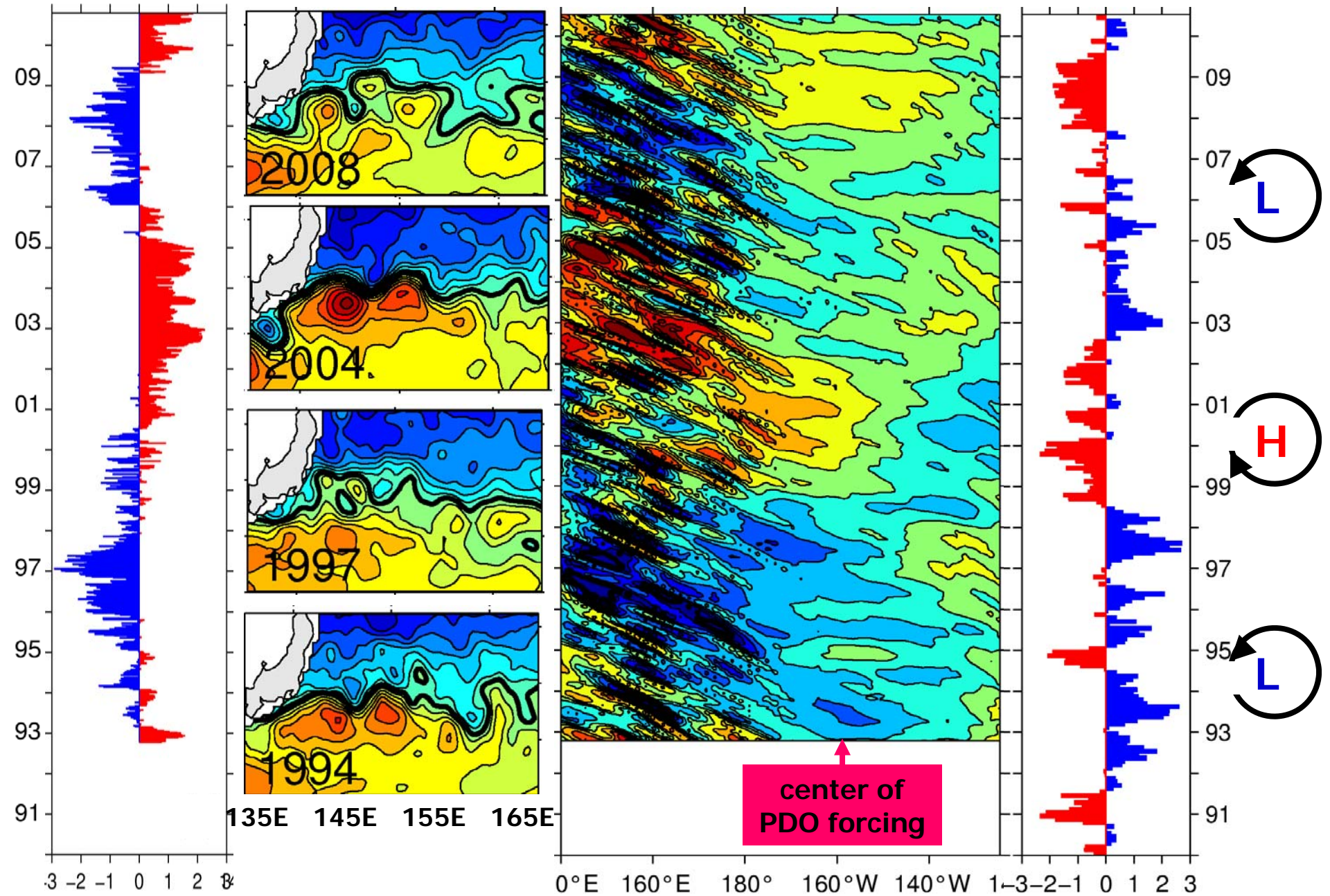


## The basis for long-term KE index prediction rests on 2 processes:

1. Oceanic adjustment in mid-latitude N Pacific is via slow, baroclinic Rossby waves + PDO wind forcing is located in the eastern basin



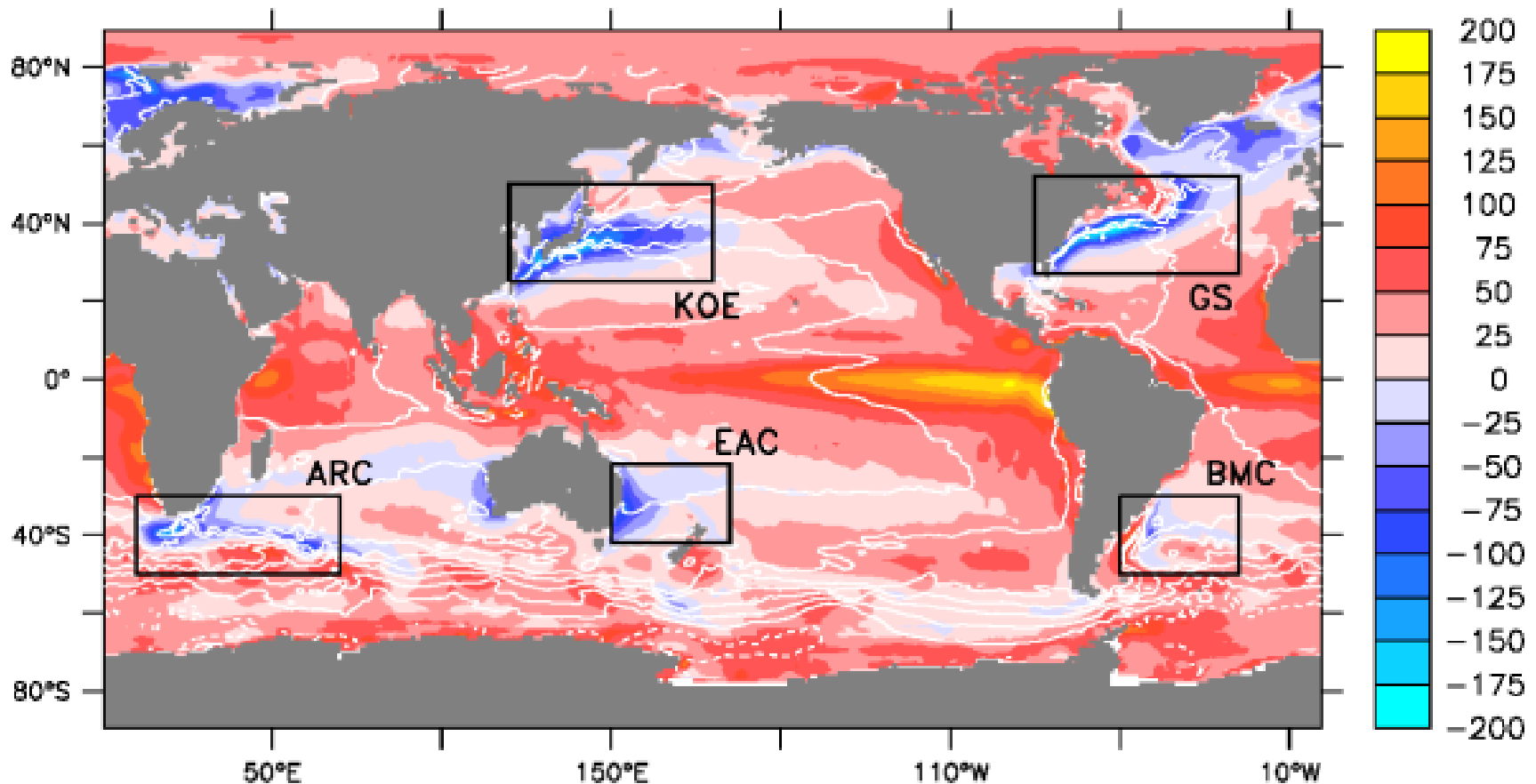
Mantua et al. (2007, BAMS)

**KE index****SSH field****SSHA along 34°N****PDO index/AL pressure**

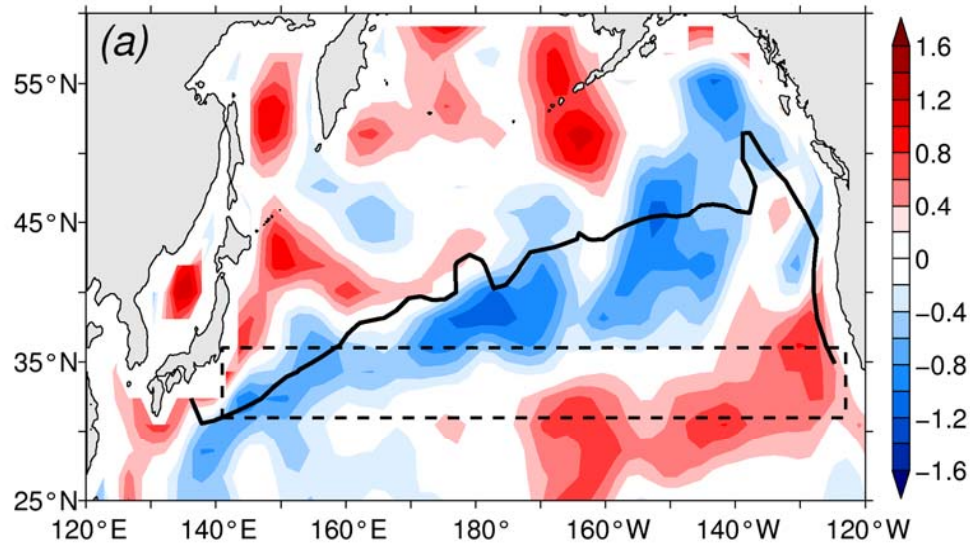
**The basis for long-term KE index prediction rests on 2 processes:**

**2. The KE jet exists in the region of intense air-sea interaction; its decadal variability affects the basin-scale wind stress curl field**

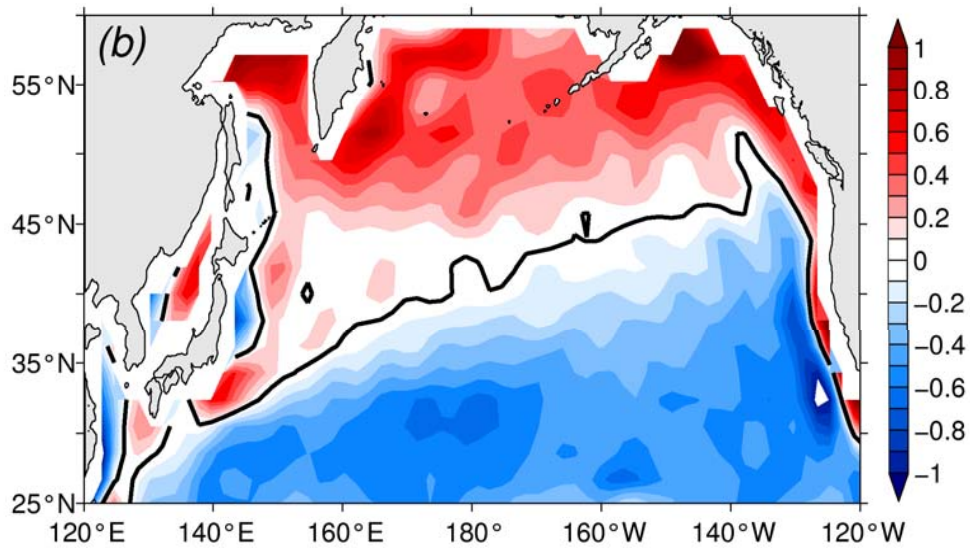
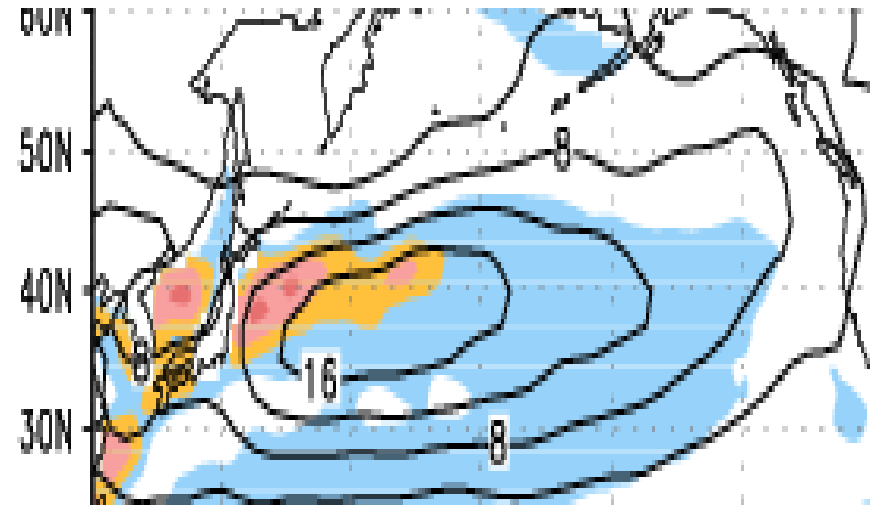
**Mean net surface heat flux ( $\text{W}/\text{m}^2$ ; Cronin et al. 2010)**



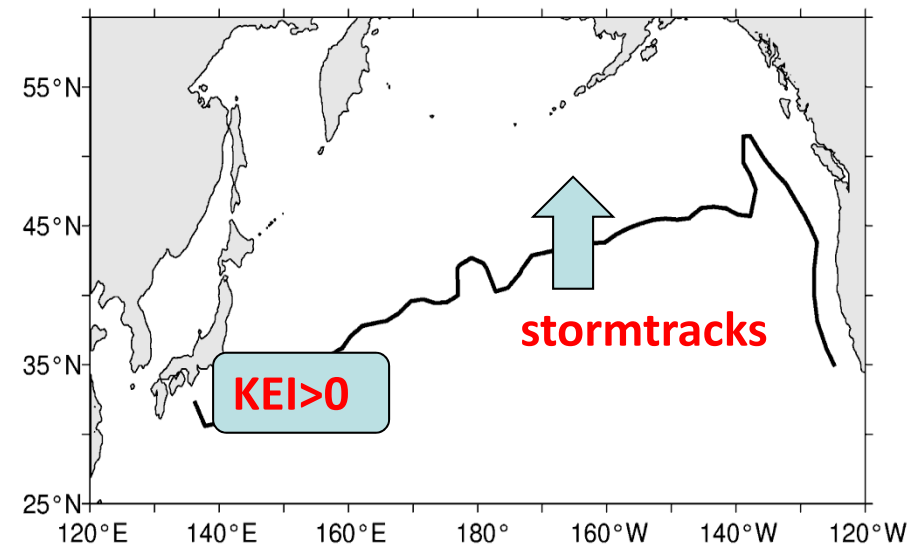
**KE index-regressed curl field**



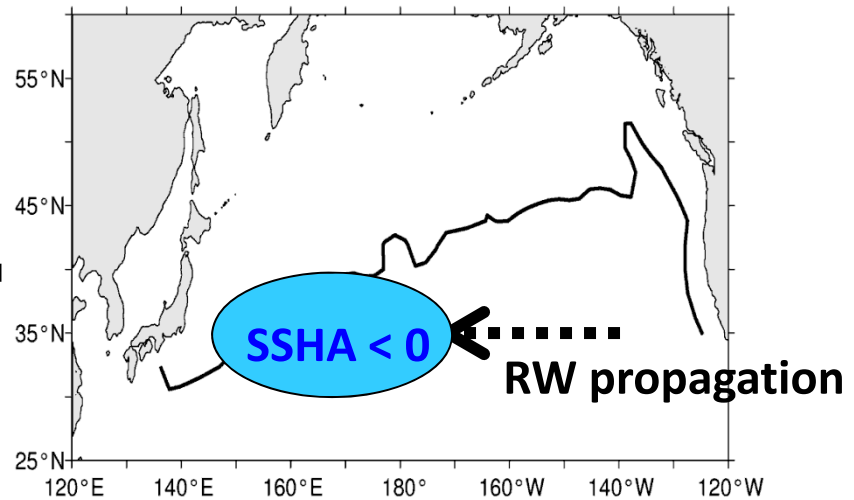
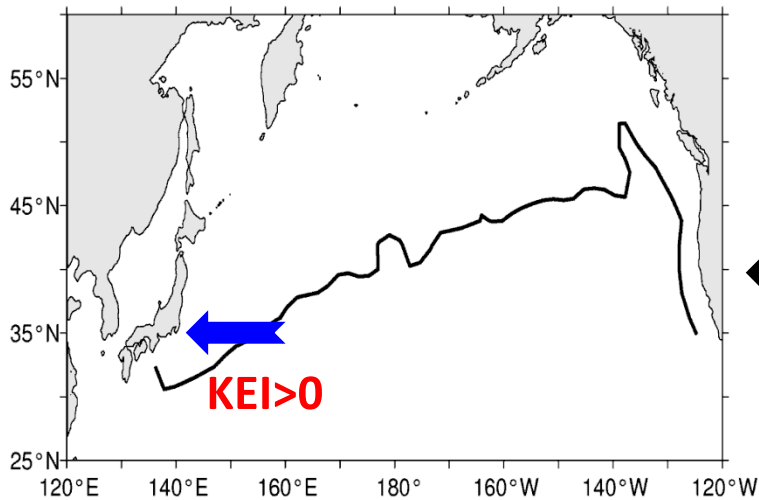
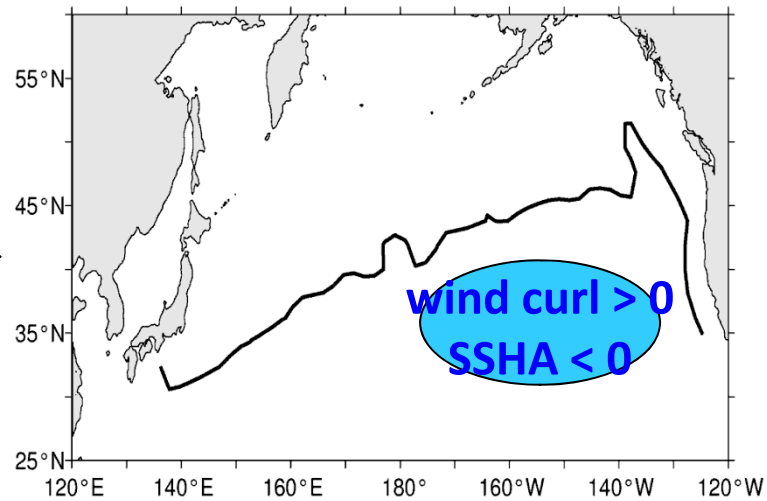
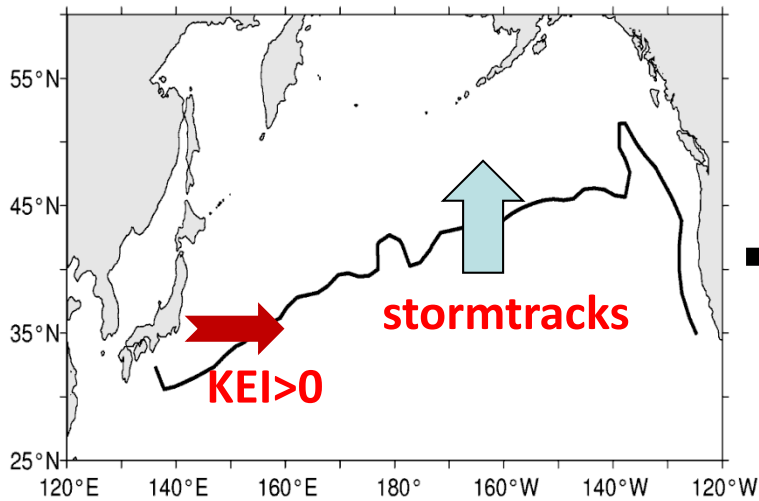
**mean wintertime stormtracks  
850mb  $v'T'$  (Nakamura et al. 2004)**



**mean wind stress curl field**

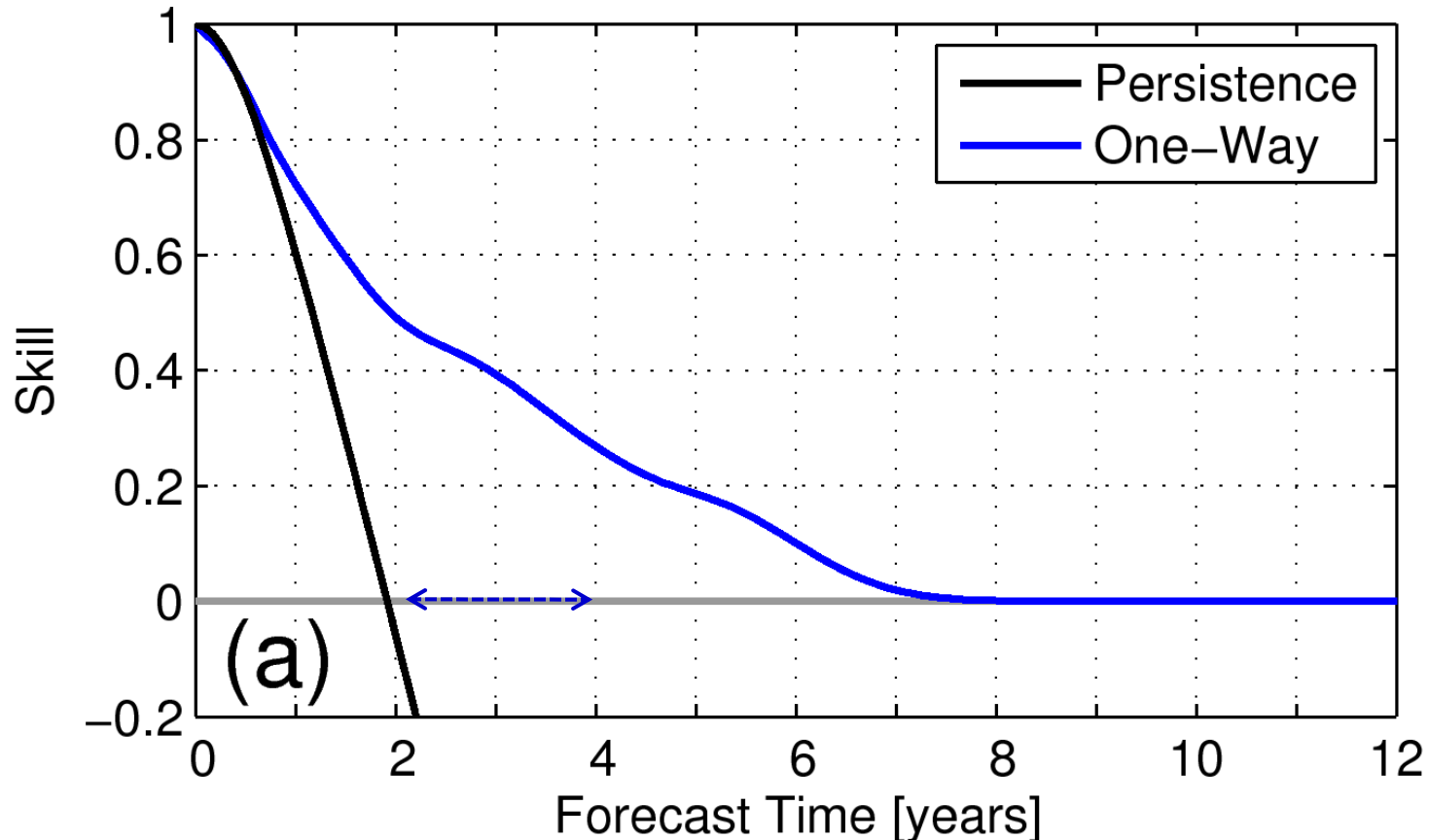


# Schematic for a delayed negative feedback decadal oscillation



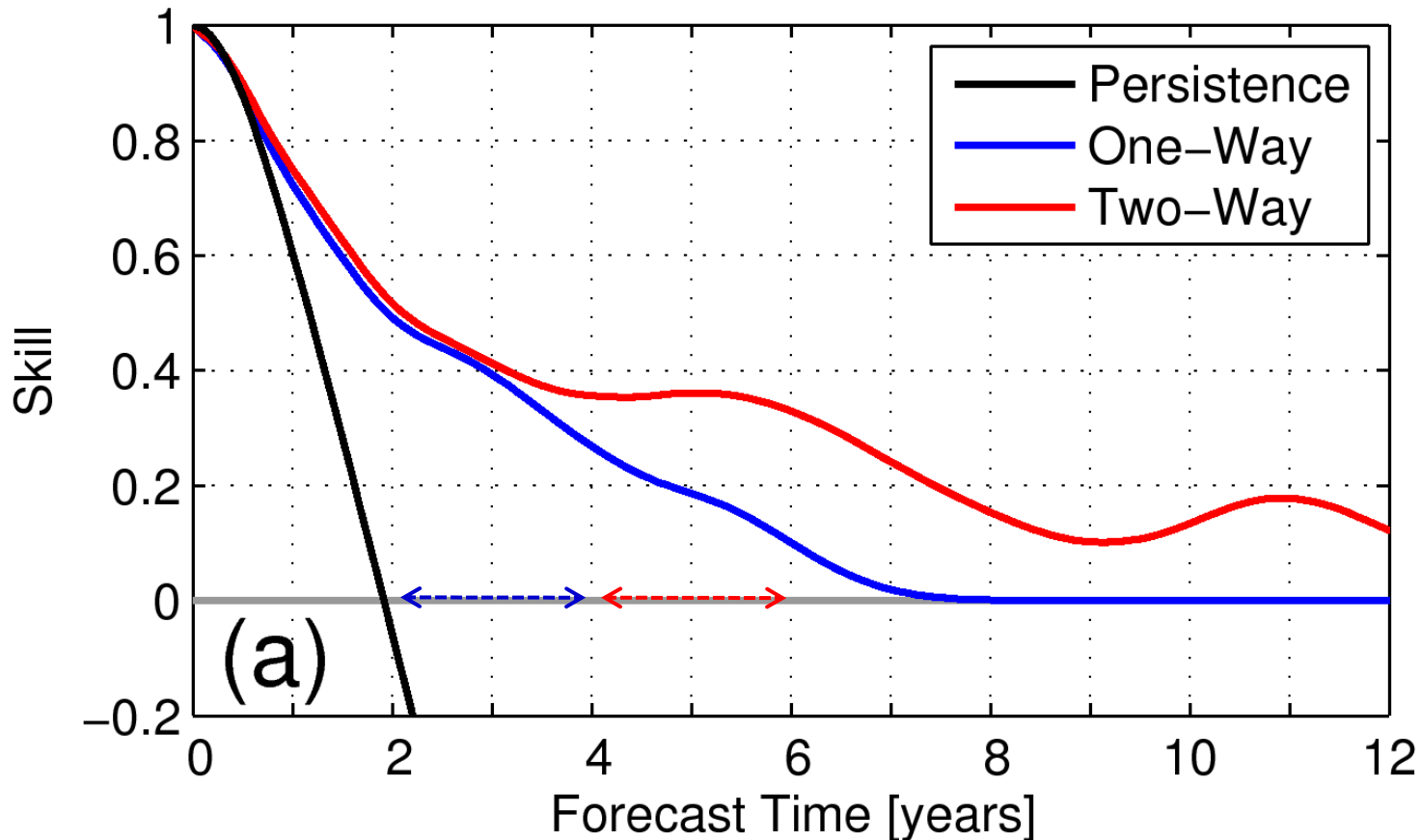
half of the oscillation cycle: ~5 yrs in the N Pacific basin

## Mean square skill of the predicted KE index



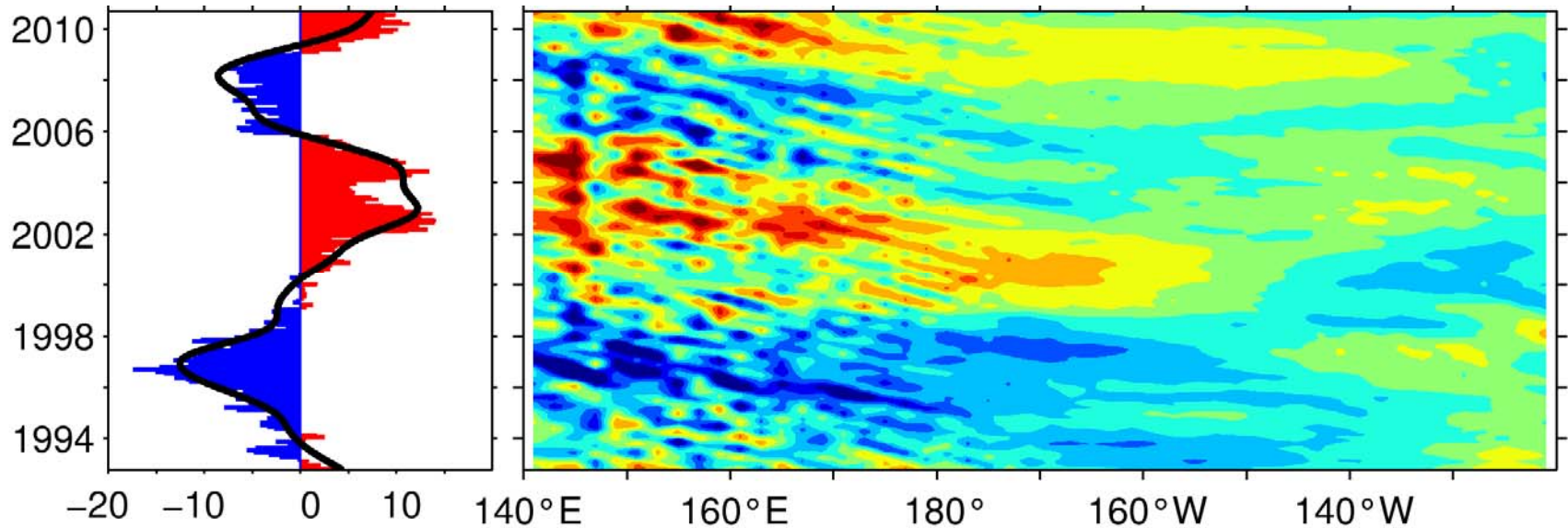
- Compared to persistence, wave-carried SSHAs increase predictive skill

## Mean square skill of the predicted KE index

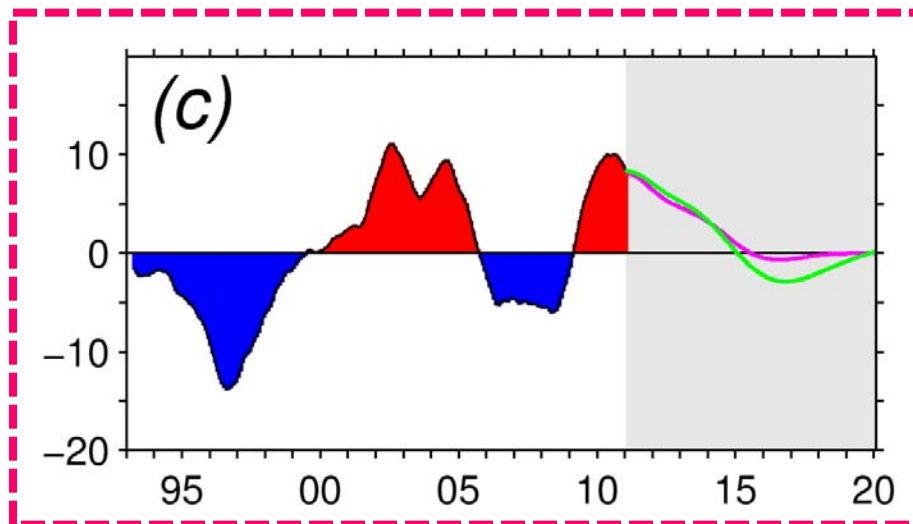


- Compared to persistence, wave-carried SSHAs increase predictive skill
- **Additional skill is gained by considering the wind forcing due to the KE feedback**

## AVISO KE index and x-t plot of SSHAs along 31°-36°N



## 2011-2020 KE index forecast based on 2010 AVISO SSH data



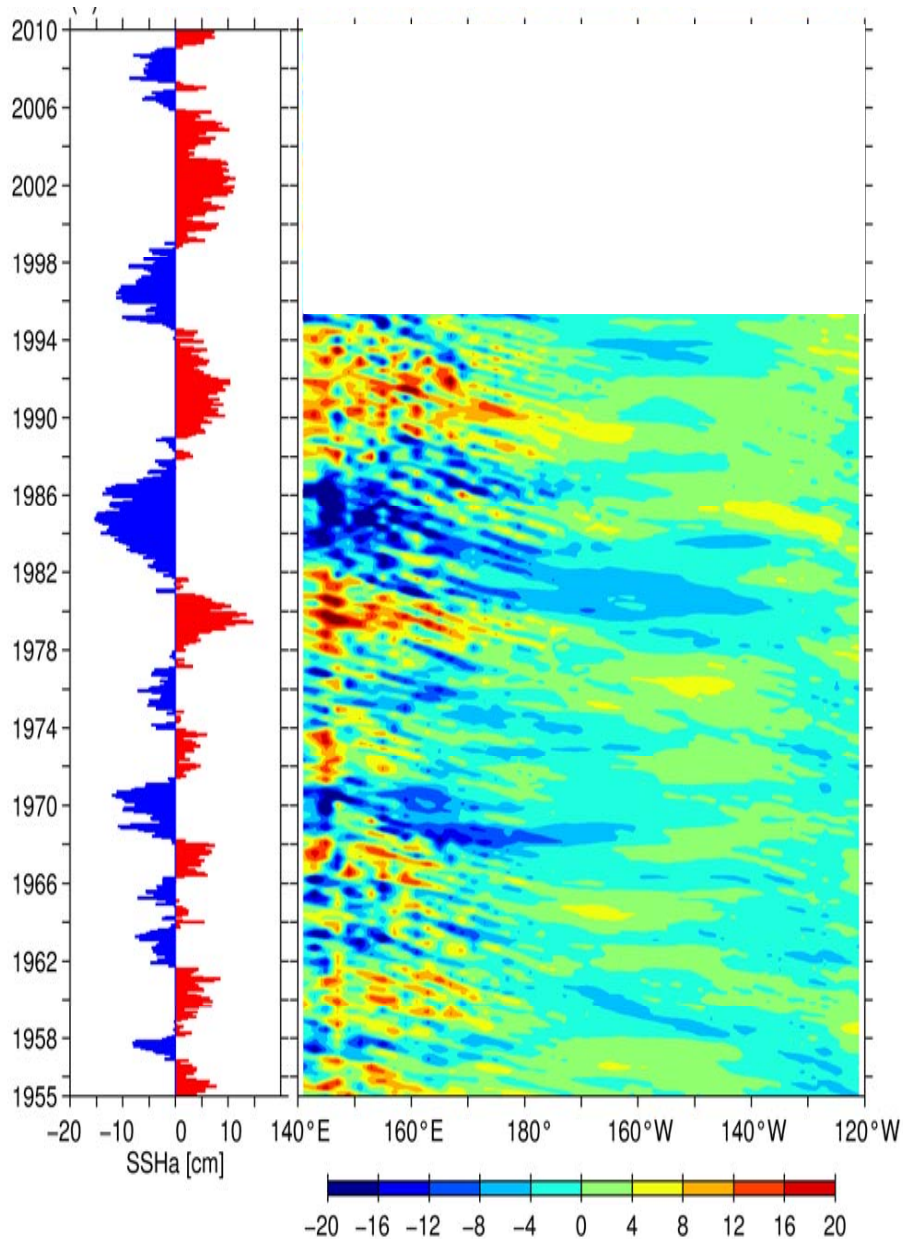


# Summary

- KE dynamic state (i.e. EKE level, path, and jet/RG strengths) is dominated by decadal variations after the 1976-77 climate shift.
- SSH anomaly signals in 31-36°N, 140-165°E provide a good proxy for the decadally-varying KE index.
- A positive KE index induces overlying-high and downstream-low pressure anomalies. This feedback favors a coupled mode with a ~10 yr timescale.  
→ Oscillatory nature of this mode enhances predictability
- Rossby wave dynamics contributes to the KE index predictability with 2~3 yr lead times.
- Inclusion of the KE-feedback wind forcing increases the predictive skill in the lead times of 4~5 yrs.



## OFES KE index and x-t SSHAs



## 1. Prediction with Rossby wave dynamics

$$h_1(x, t) = h_{\text{obs}} [ x + c_R(t - t_0), t_0 ]$$

where

$h_{\text{obs}}(x, t_0)$  : initial SSHAs

$c_R$  : Rossby wave speed

## 2. Prediction with Rossby wave dynamics + KE feedback to wind forcing

$$h_2(x, t) = h_1(x, t) +$$

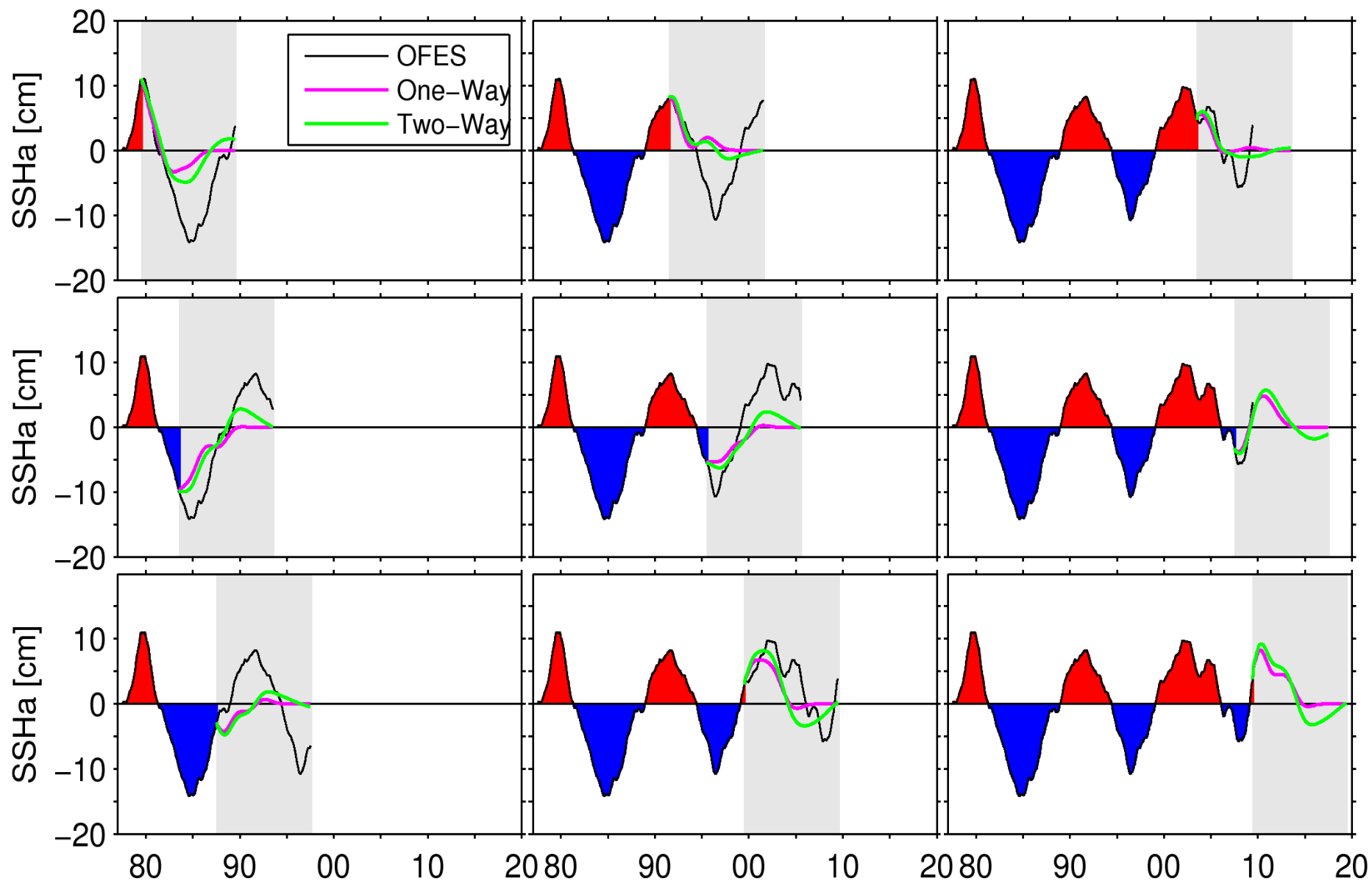
$$\int_{t_0}^t b [ x + c_R(t' - t_0) ] K(t') dt'$$

where

$b(x)$  : feedback coefficient

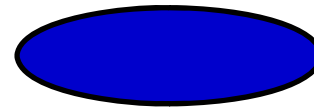
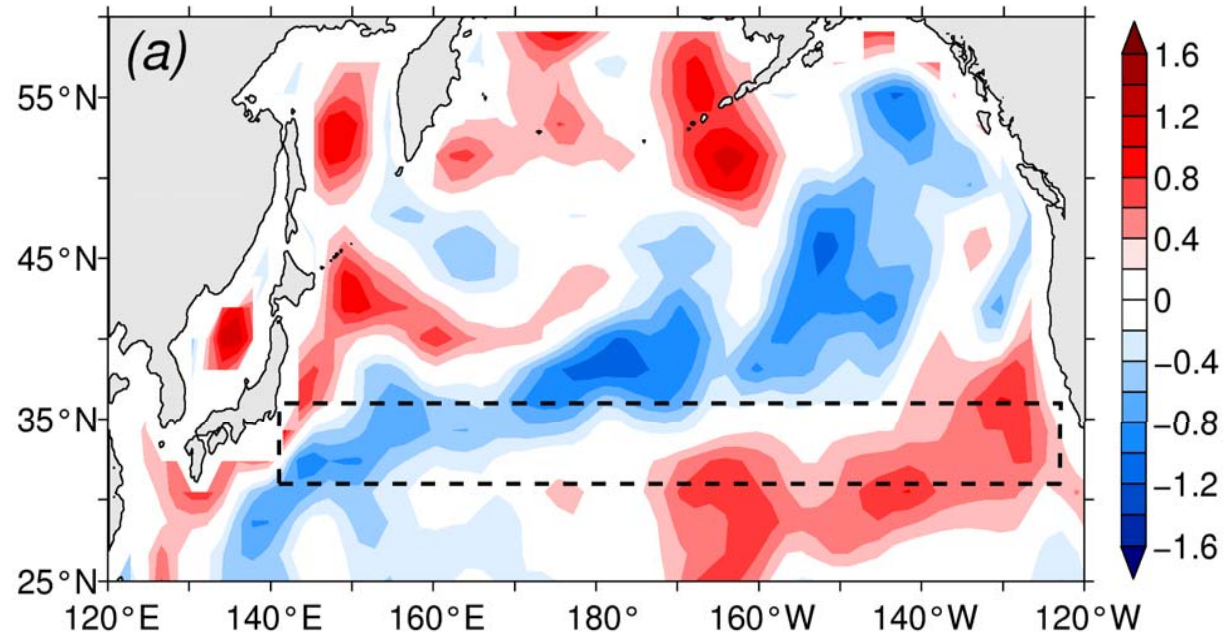
$K(t)$  : forecast KE index

# Examples of decadal KE index predictions and verifications

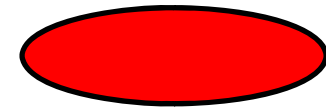


# Lagged regression between the OFES KE index and NP curl-tau field

- NCEP reanalysis data (1977-2010)
- ENSO signals (Nino 3.4) regressed out
- $b(x,y)$  with  $m = 2$  months

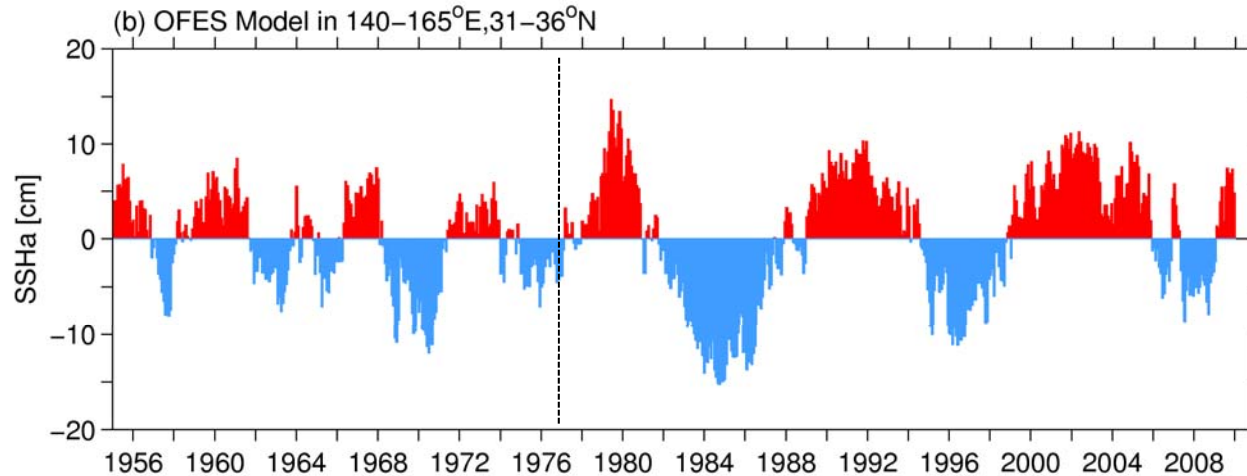


$p' > 0; \text{curl-tau} < 0$



$p' < 0; \text{curl-tau} > 0$

Our goal is to explore the predictability of **the proxy KE index** after the 1976-77 climate shift

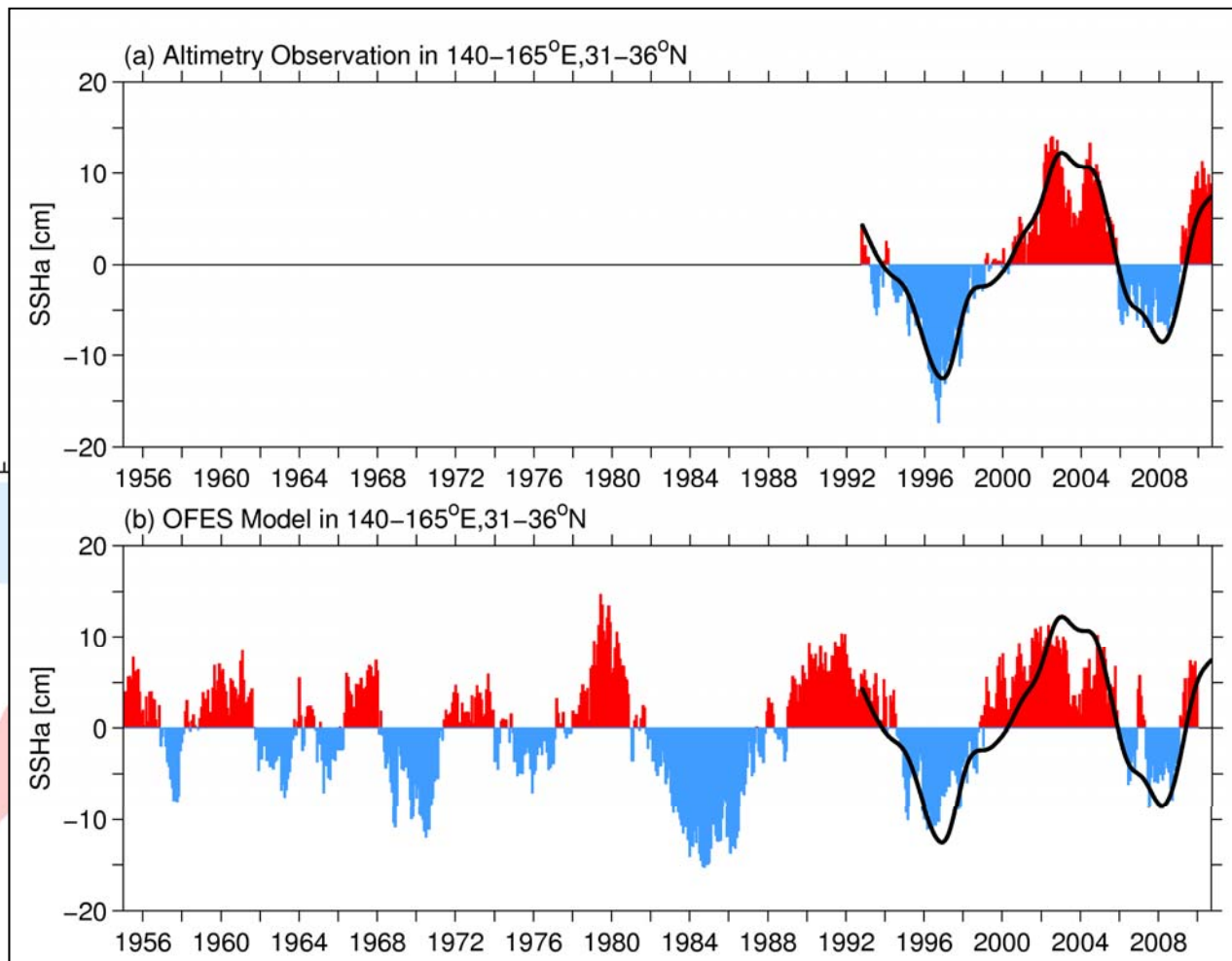
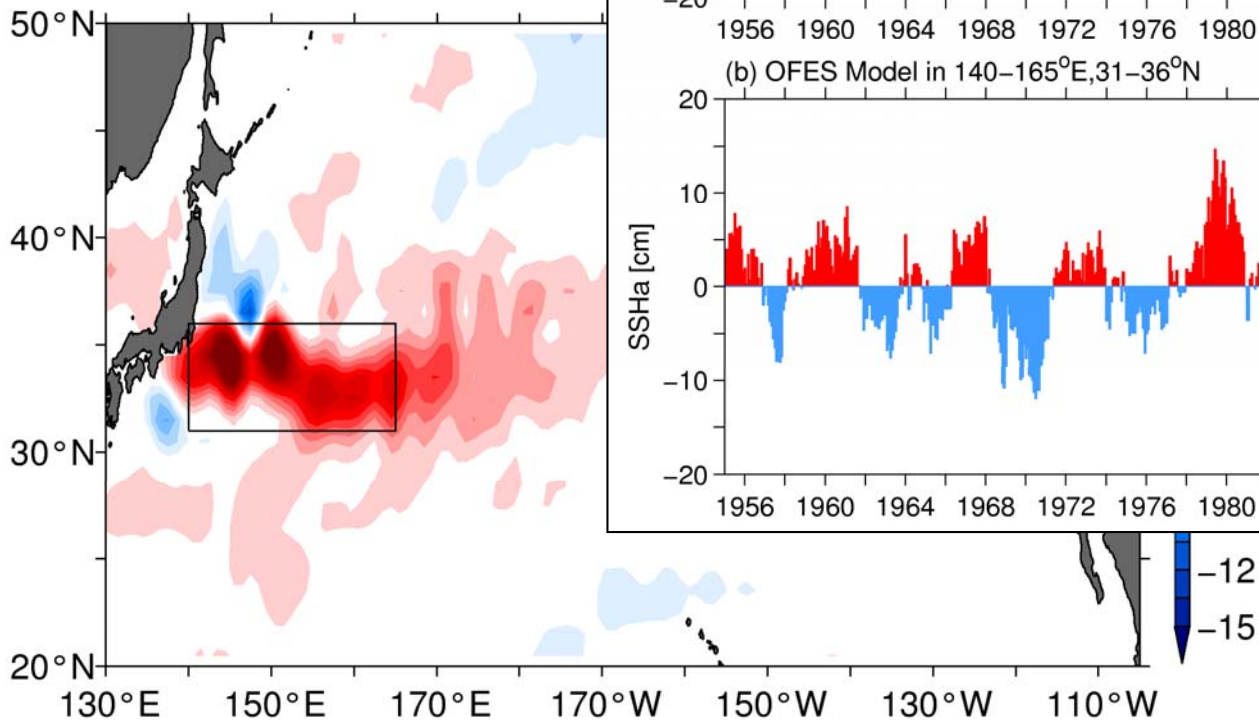


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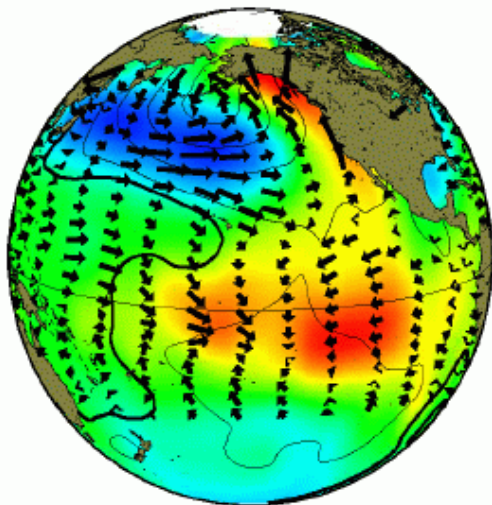
# Lengthening **the (proxy) KE index** using the OFES hindcast results

**OFES hindcast run  
simulates well the  
observed decadal KE  
variability (see  
Taguchi et al. 2007;  
Nonaka et al. 2008)**

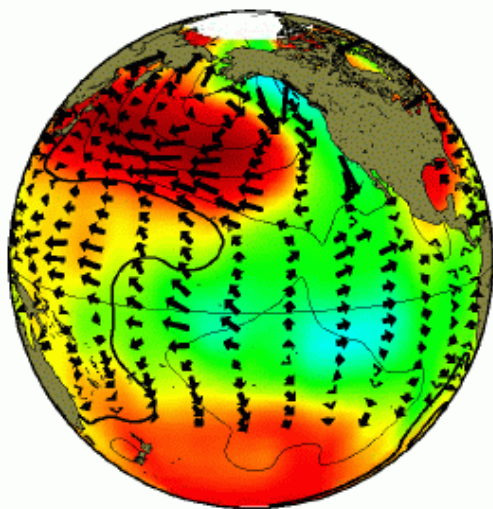


# Composite PDO patterns

positive phase

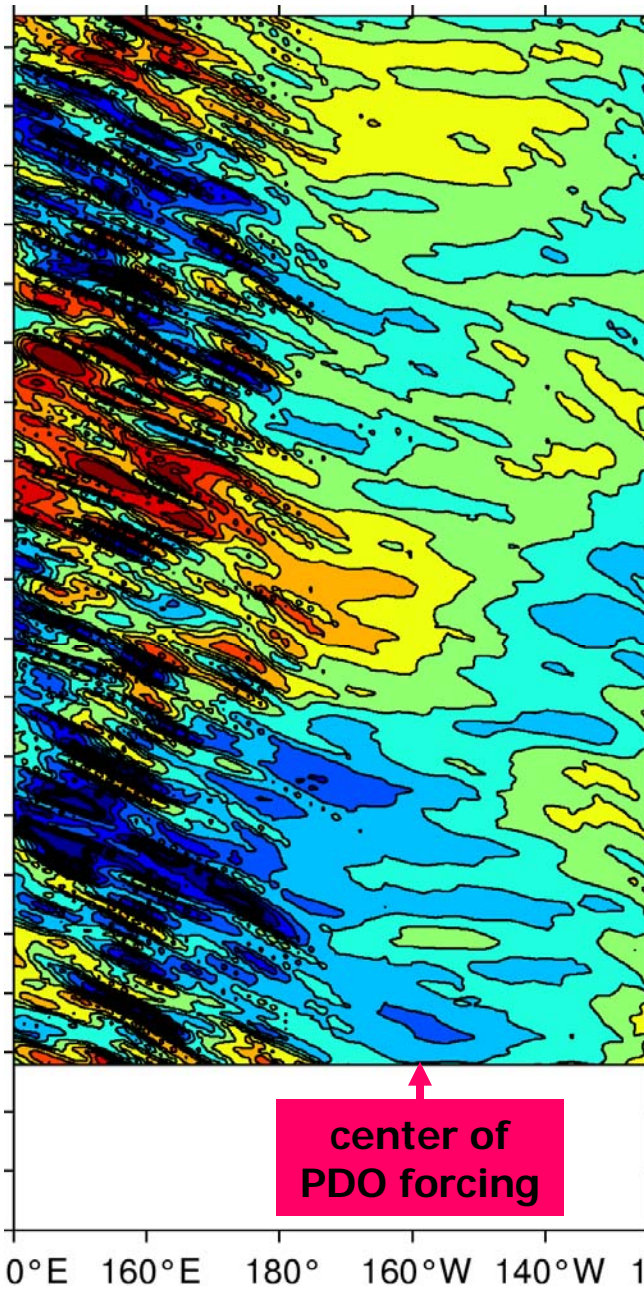


negative phase



Mantua et al. (2007, BAMS)

# SSHA along 34°N



# PDO index/AL pressure

