

SAR imagery: an asset to the sea ice altimetry communauty

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Foreword on sea ice backscattering from single-pol SAR images (L-C-X bands)

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Sea ice varies according to many parameters:

- EM parameters such as the dielectric properties of brine inclusion and of ice depending on salinity

- Geo-physical parameters such as the sea ice thickness, its top surface roughness, its porosity, the size/shape of scatterers (brine inclusions, air particles...).

- Potential snow cover (dielectric constant of 1.5 for dry snow – varying also with density, frequency and wetness)

⇒ Large number of parameters impacting sea ice backscattering, potential ambiguities between open water and thin ice in leads with single-pol SAR imagery



..... Freshwater ice

Attenuation by sea ice depending on the frequency of EM microwaves (from Fingas and Brown)

Surface vs volume backscattering at 20-50°

- Average thickness of FYI about 1.5 meter (MYI: 2.5 m)
- FYI: Limited volume backscattering from C-band and higher frequencies

MYI: Limited volume backscattering from L-band and higher frequencies



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Benefit of polarization diversity to characterize sea ice and leads

Assessment of sea surface roughness in leads

Few words on Sentinel-1

On the use of Wave mode specific acquisition mode to complement altimetry iceberg detection





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Backscattering from leads in SAR imagery

Thin ice lead

- Open water rapidly freezes (directly exposed to the cold air temperatures) -> thin ice skim hence formed followed by black ice

- Backscattering close to NESZ of current available sensors

\Rightarrow Potential ambiguity with open water in leads (limited fetch -> very calm sea surface)



Simulations and measurements of EM backscattering at C-band (Left) and L-band (Right) from thin lead ice (from Nghiem et al. 1995)





Automatic lead detection with single-pol X-band data

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Benefit of dual-pol HH and VV C-band

New thin ice:

Slushy layer of the order of millimetre thick and composed of ice and brine with salinity as high as 100% exists on the top of new ice

High VV/HH Polarization Ratio (PR) first observed by Nghiem et al. 1995





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New thin ice detection

- Ratio between measured and simulated PR (Mouche et al. 2007) assuming open water
 - If very high, potential new thin ice

⇒ Thin ice generates high PR values for a large set of ice/radar configurations ?

Nghiem S.V, Kwok R., Yueh S.H., Drinkwater M.R., "Polarimetric signatures of sea ice 2. Experimental observations", Journal of Geophysical Research, Vol. 100, no. C7, pp. 13,681-13,698, 1995



In (Nghiem et al. 1995), about VV-HH coherency over sea ice -> Valid for long EM waves (penetrating Lband data, and to a lesser extend C band)

- MYI: high correlation between HH/VV with spherical air inclusions
- FYI: Decorrelation due to ellipsoidal shape of brine inclusion, and preferred vertical orientation
- Thin Lead ice: same as FYI, with additional decorrelation due to noise

What about higher frequency data ?.... Below X-band VV/HH TSX data



High Polarisation Ratio over leads with possibly thin ice (open water otherwise)

High co-pol coherency over homogenuous FYI ice

Deformed ice / fragment of MYI induces lower coherency at X-band



NRCS VV [dB] 20100314_154242

400

500

1000

500

000

500

100

200

300 wspd:7.5 inc:35.9

Dual-pol HH/VV analysis in X-band: study case



Observed ice thickness with respect to co-pol coherence with TSX X-band data (from Kim et al. 2012)



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• Wind stress: primary forcing mechanism on the variability and magnitude of small-scale windwaves

 Perturbation of wind roughened open sea surface in ice to wavelengths that are resonant to C-band SARs is driven by upwind fetch and wind speed.

Morphology of leads (size, orientation) with respect to wind contributes to the availability of fetch (i.e distance from ice edge)



Measured wavenumber in melt pond with mean size [10-30

Variable	Pearson's Correlation to σ° (HH, VV) a			
	20°	30°	40 °	50°
Wind Speed	.298	.346	.327	.317
	(.012)	(.003)	(.006)	(.007)
Pond Fetch	.479	.500	.508	.502
	(.000)	(.000)	(.000)	(.000)
Pond Depth	.452	.454	.469	.477
	(.000)	(.000)	(.000)	(.000)

a. Correlation does not change as a function of polarization.

Wind speed, upwind fetch length, and depth contribute to C-band σ° from a FYI sea ice melt ponded surface

Even at low wind speeds (e.g., 3 m/s): decametric ponds are reactive to Bragg scattering at C-band.

* Fetch: Waves are formed by wind blowing along the water's surface. Surface roughness is dependent on wind speed, fetch length and duration of time the wind blows consistently over the fetch. Wind "fetch" is the distance the wind blows over water with similar speed and direction.



Measurement of fetch effect in lead with open water

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Automatic fetch assessment

o Automatic lead detection then followed by lead identification

Lead is rotated wrt to upwind direction, and NRCS analyzed upon distance to the ice edge (wrt wind direction)





Measurement of fetch effect in lead with open water (#1)



For this specific case, wind-wave equilibrium with about 3.8 km fetch distance

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Measurement of fetch effect in lead with open water (#2)

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Measurement of fetch effect in lead with open water (#2)





Measurement of fetch effect in lead with open water (#3)

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Presence of floating ice (here probably melting ice) prevents from the establishment of wind-wave equilibrium in the lead



Cause of the backscattering and fetch dependence



- Similar distances found with X-band data
- Possible effect of fetch/lead
 - Fetch effects on EM backscattering: wind-sea interaction is not at its equilibrium.
 - Bragg wave is instantaneous.
 - Contributing longer waves not formed yet.
 - Atmospheric effect
 - Acceleration of the wind speed over the leads as modeled by (Dare et al.) over leads and MIZ
 - Cold wind from ice (-10°C) interacts with warmer open water (-2°C) -> turbulent atmospheric layer : instability -> favor the air/water interaction
- Loss of energy due to joint fetch/atmospheric effect difficult to model



R.A. Dare and B.W. Atkinson Numerical modeling of atmospheric response to polynyas in the Southern Ocean sea ice zone



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Sentinel-1 acquisition modes



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Wave mode: a systematic acquisition mode for iceberg detection around Antarctica



20 km by 20 km vignettes, at 5 m by 5 m spatial resolution, every 100 km along the orbit, acquired alternately on two different incidence angles

- VV polarization
- Quasi-systematic default mode over open ocean when not covered by EWS/IWS
- In complement of EWS (acquisition mostly during winter time over sea ice, and summer time in Weddell sea)
- Less spatial coverage than altimetry, but high resolution with potential < 10 meter detected icebergs
- Reconstruct the whole story of icebergs (from large tabulars to small icebergs and bergy bits) ?
 - Antarctic Iceberg Tracking Database from BYU for large tabular -> IFREMER altiberg database -> ASAR/ENVISAT & S-1 Wave mode database



Detection comparison between ENVISAT WV mode and ENVISAT Altimeter

- First assessment between ENVISAT ASAR WV data and altiberg ENVISAT database during 2003
 - Coverage of ENVISAT WV mode between -45 and -66 was about 3% only of the ocean surface during one repeat 35 days cycle
 - Based on CFAR detector with fixed kernel size (below configured for icebergs no larger than 100 meters)





First assessment between ENVISAT ASAR WV data and altiberg ENVISAT database in 2003

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- Coverage of ENVISAT WV mode between -45 and -66 is about 3% only of the ocean surface for one repeat 35 days cycle
- Based on CFAR detector with fixed kernel size (below configured for icebergs no larger than 100 meters)
- No sea ice flagging for WV mode detection



Promising agreement with same detected hot spots such as the Weddell Sea and the West Ice shelf (to a lesser extent Ross sea) ... Except for the Amundsen sea



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SAR imagery of primary interest to analyze and complement altimetry data over sea ice ... and vice versa !!!