

SWOT

Future technology : Wide Swath Altimetry

SWOT

(Surface Water Ocean Topography)

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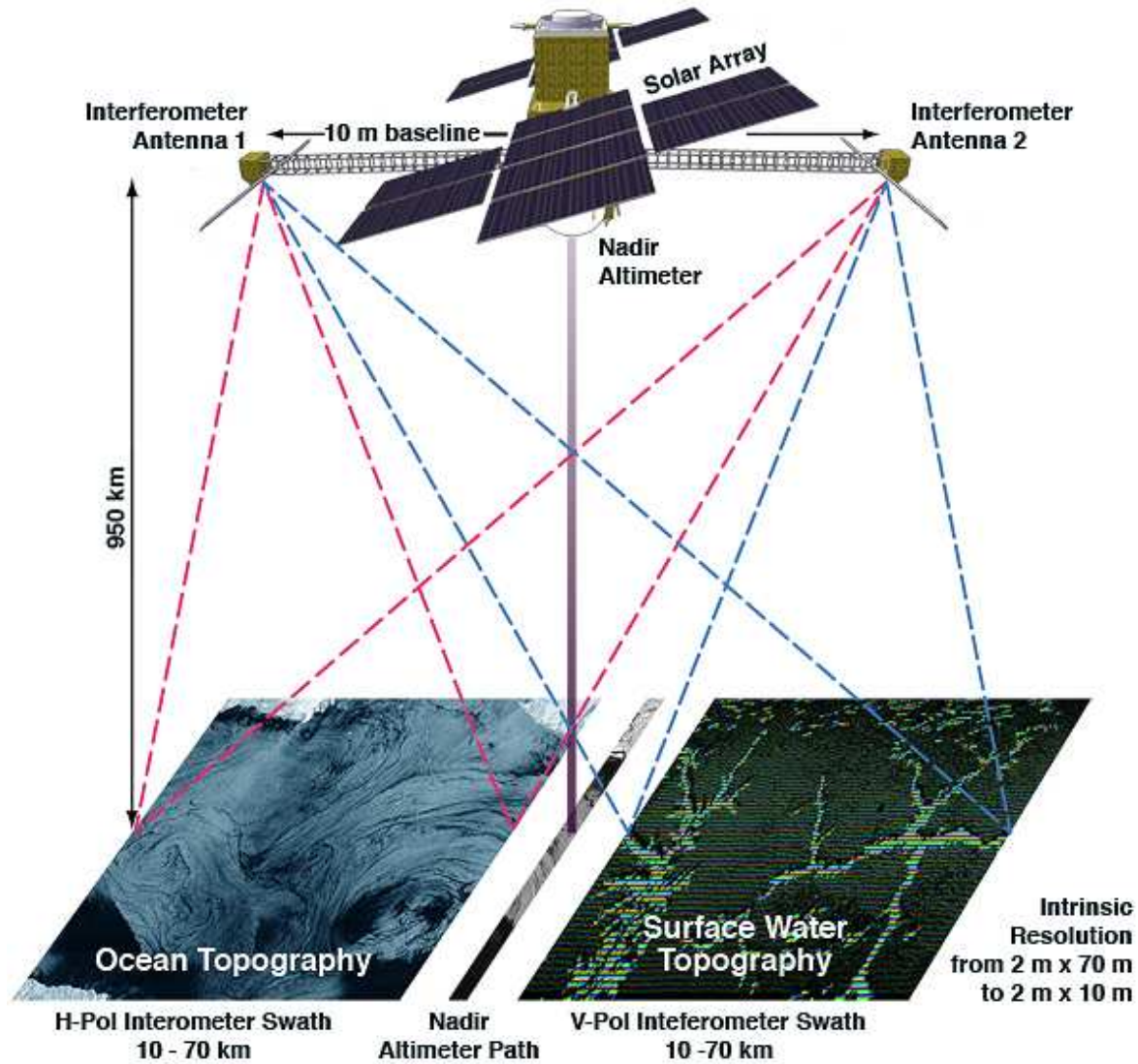
1-1



August, 2008

SWOT

Schematic of the SWOT measurement system



SWOT

SWOT

Fine spatial and temporal observations of :

Surface Water
Ocean Topography



1-3



August, 2008

Surface water storage includes rivers, lakes, artificial reservoirs, swamps, and flood plains.

Due to climate variability, climate change and direct anthropogenic forcing, land water storage fluctuates both in time and space.

However due to limited and heterogeneous in situ networks, the spatio-temporal distribution of surface waters remains poorly known.

=> Satellite observations allow monitoring of the continental water volume and discharge.

Satellite altimetry measures water level fluctuations of :

- continental lakes,
- major rivers
- flood plains

Satellite gravimetry (eg GRACE) measures fluctuations in :

- water storage in soils,
- underground reservoirs
- snow pack

<http://www.legos.obs-mip.fr/en/soa/hydrologie/hydroweb/>

Altimetric data base at CTOH / LEGOS contains time series over water levels of **large rivers, lakes** and **wetlands** around the world.

- based on altimetry data from Topex/Poseidon for rivers, ERS-1 & 2, Envisat, Jason-1 and GFO data are also used for lakes.



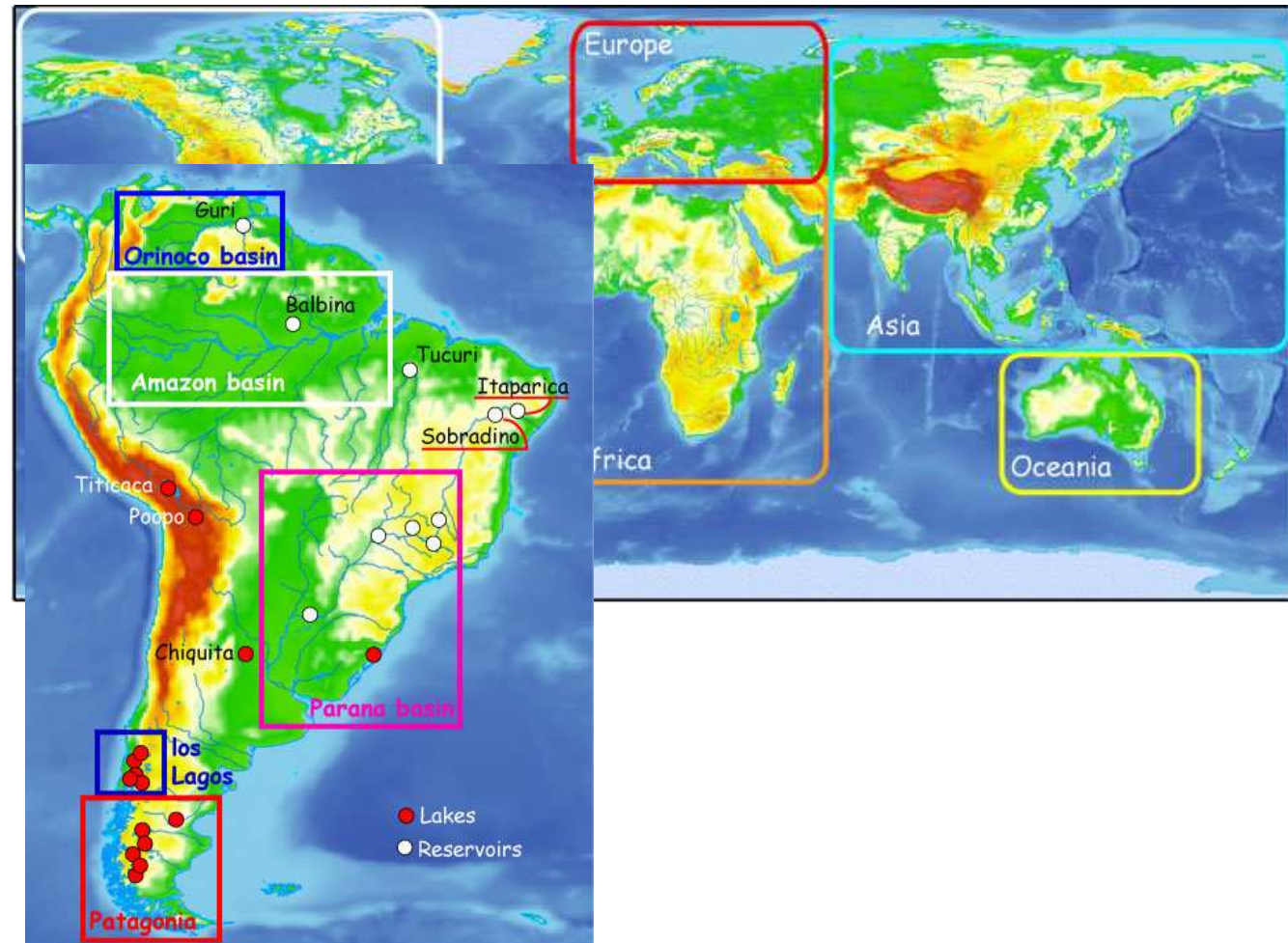
100 lakes are available



250 « virtual stations » along rivers

South America:

- **Lakes**
 - [Argentino](#)
 - [Chiquita](#)
 - [Huapi](#)
 - [Lagoa do Patos](#)
 - [Lagoa Gen Carrera](#)
 - [Llanquihue](#)
 - [Nahuelhuapi](#)
 - [Poopo](#)
 - [Ranco](#)
 - [Rinihue](#)
 - [San Martin](#)
 - [Strobel](#)
 - [Titicaca](#)
 - [Todos los Santos](#)
 - [Viedma](#)
- **Reservoirs**
 - [Balbina](#)
 - [Furnas](#)
 - [Guri](#)
 - [Ilha Solteira](#)
 - [Itaparica](#)
 - [Novaponte](#)
 - [Posadas](#)
 - [Sobradino](#)
 - [Tres Marias](#)
 - [Tucuruí](#)
- **River basins**
 - [Amazon](#)
 - [Orinocco](#)
 - [Parana](#)



Lake levels are based on merged Topex/Poseidon, Jason, ENVISAT and GFO data provided by ESA, NASA and CNES data centers.

Processing :

1Hz (1 sec) altimeter range measurements are used.

Classical corrections are applied (orbit, ionospheric and tropospheric corrections, polar and solid Earth tides and sea state bias).

Geoid corrections are not accurate over land lakes : so the mean lake surface is calculated from the repeat altimetry data.

Multi-satellite data is computed in a 3-step process:

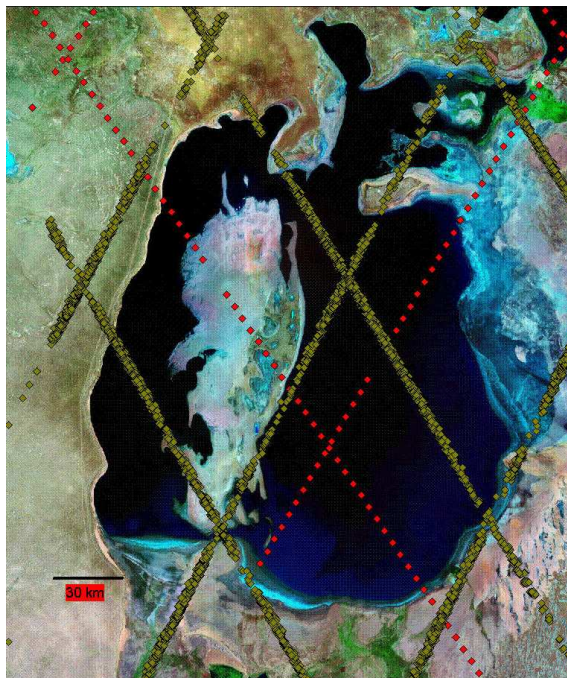
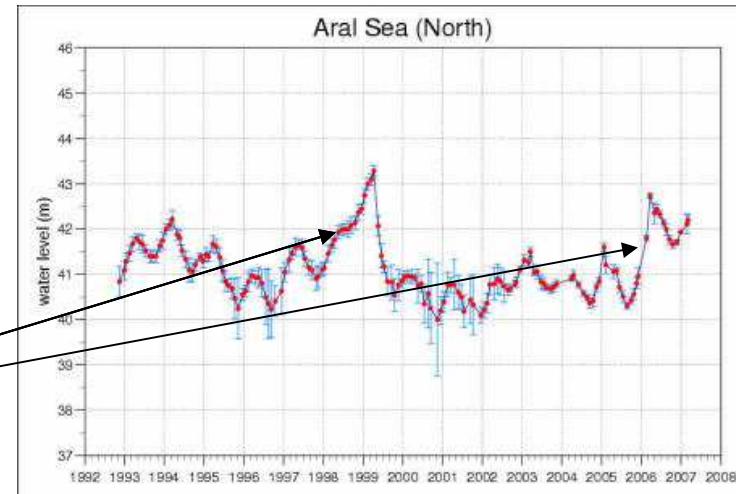
1. Each satellite data set are processed independently.
2. Biases between different satellites are removed using T/P data as reference.
3. Lake levels from the different satellites are merged on a monthly basis (recall that the orbital cycles vary from 10 days for T/P and Jason, to 17 days for GFO and 35 days for ERS and Envisat).

We generally observe an increased precision of lake levels when multi-satellite processing is applied.

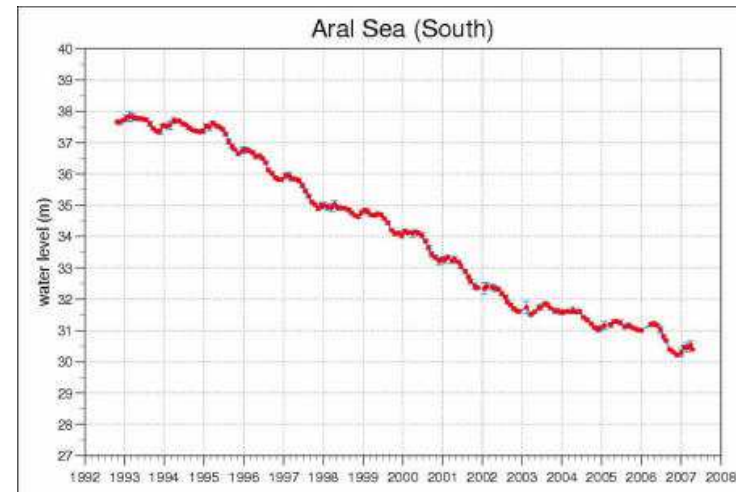


Northern
Aral Sea

Barrage



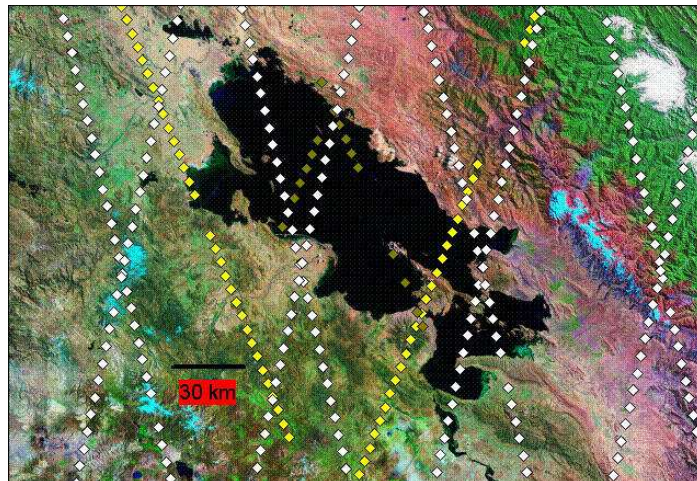
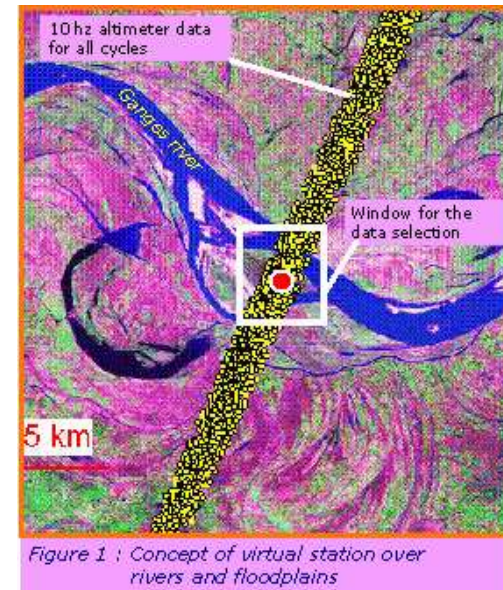
Southern
Aral Sea



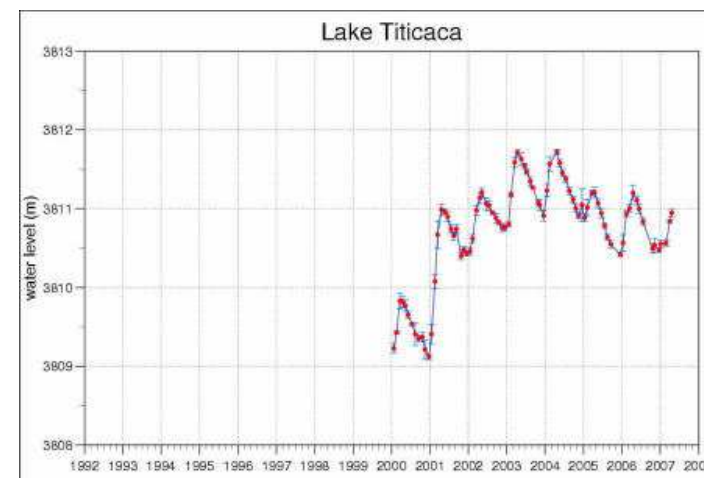
Virtual stations are defined at the intersection of the T/P track with the river.

All available 10Hz (1/10 sec) data within a rectangular «window» are used.

For each 10-day T/P cycle, 10Hz data are geographically averaged to give mean river height.

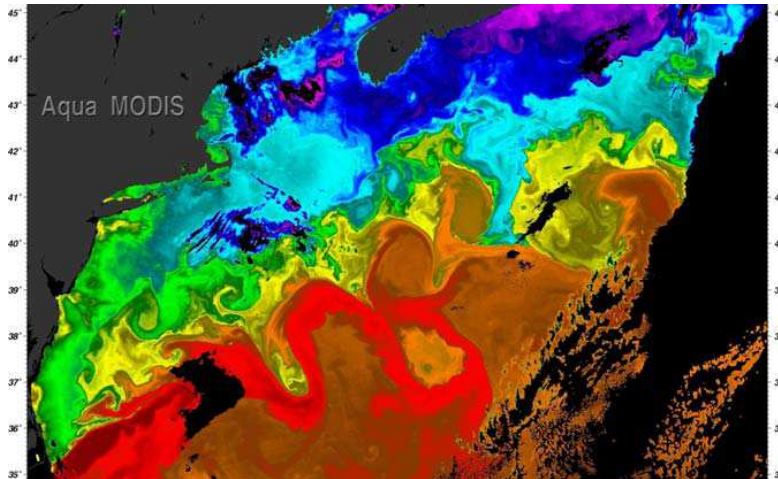


Lake Titicaca



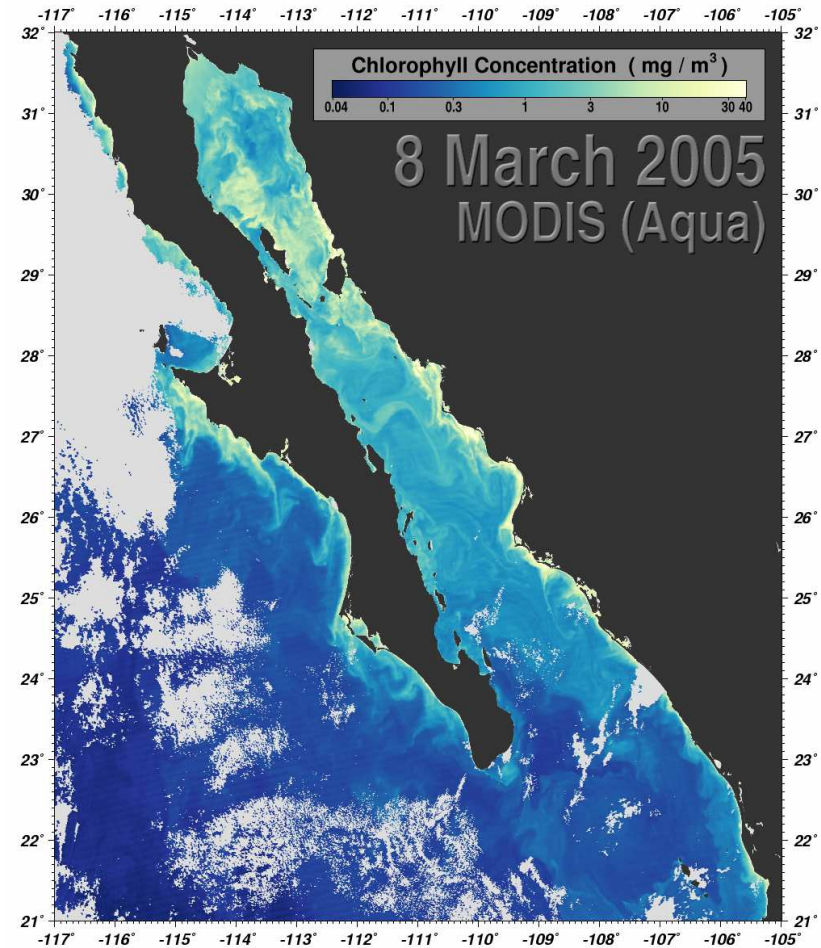
- Water movement in wetlands and flood plains is essentially unmeasured
- Number of in-situ river gauges decreasing
- Monitoring with traditional altimetry is limited :
 - Wave forms optimised for open ocean surfaces
 - Inter-track distances miss small water bodies, and undersample rivers and lakes

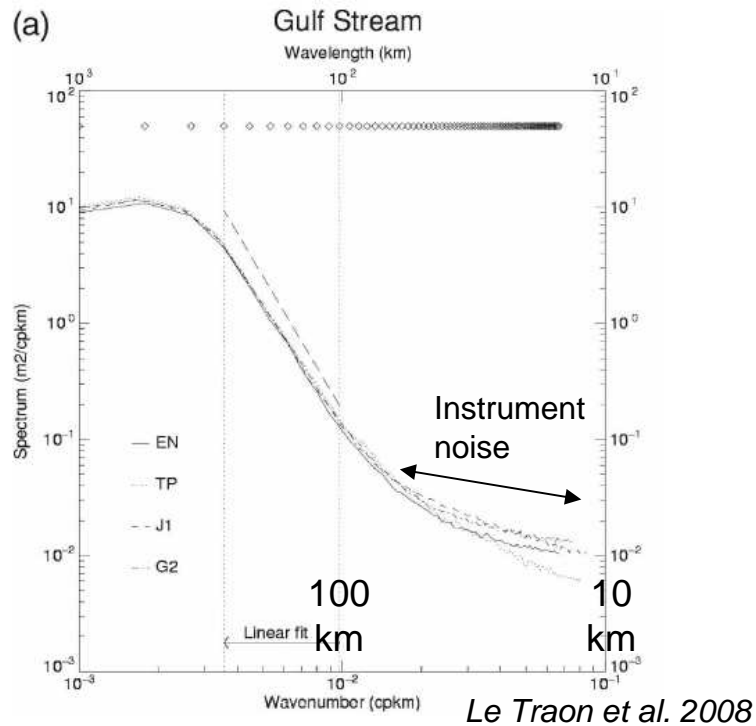
=> Need high resolution measurements over the continental water surfaces



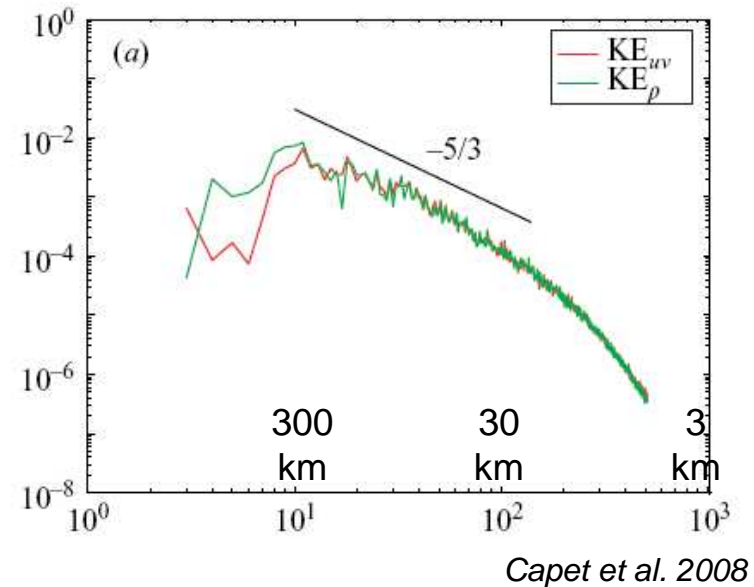
Ocean fronts and filaments studied with SST and ocean colour

- Scales 1-50 km, 2-30 days
- Important for vertical exchange of heat, carbon, nutrients
- Important regions of mixing





Altimetric sea level spectra
White noise < 50-70 km



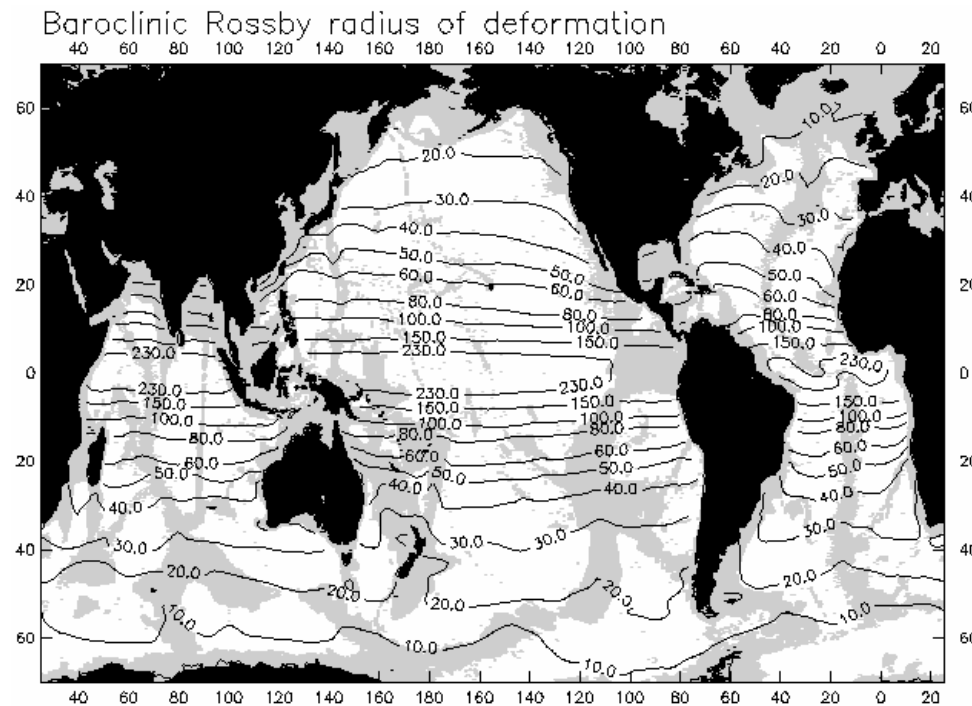
$1/54^\circ$ Model velocity spectra :
spectrum remains red

Ocean remains energetic at small space scales < 100 km

Traditional altimetry cannot resolve these small space and time scales

- 100 km spatial resolution of multi-mission altimetry : missing part of the mesoscale eddy energy (Rossby radius) at mid to high latitudes, and higher-order modes

Eddy scales
= $2\pi R$

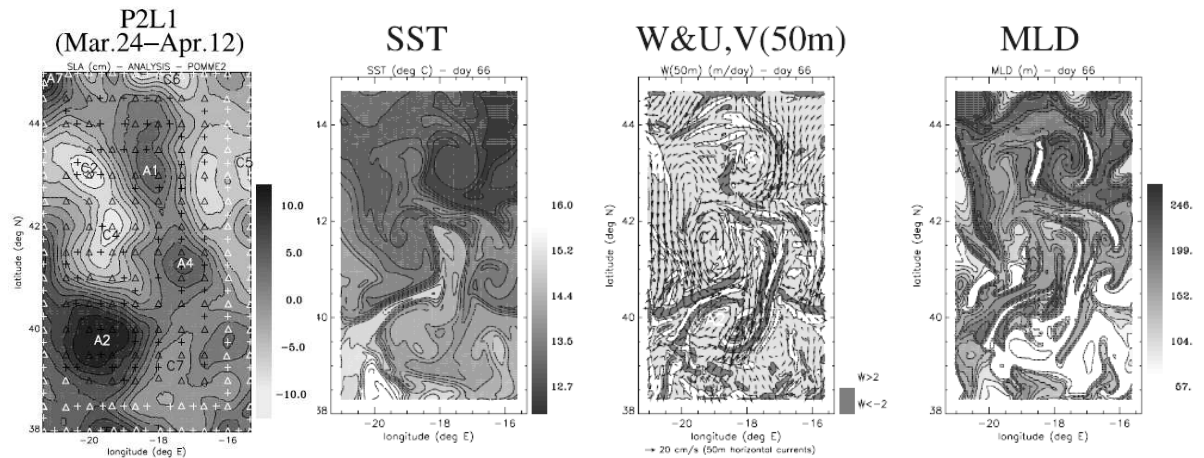


Chelton et al. 1998

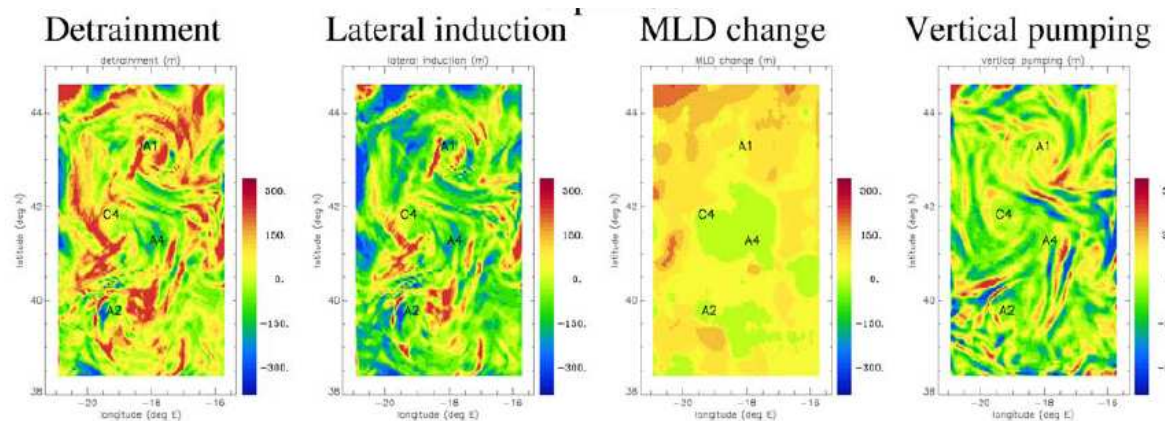
Small-scale filaments (10-20 km) surrounding mesoscale eddies important for :

- advecting tracers (SST, chlorophyll)
- inducing strong **vertical velocities** > 25 m/day
- MLD changes - exchange of nutrient-rich deeper water with surface layer
- formation of mode and intermediate water masses
- Biogeochemical cycles

Very high resolution regional ocean models ($1/20^\circ$ or 5 km with 69 vertical levels)



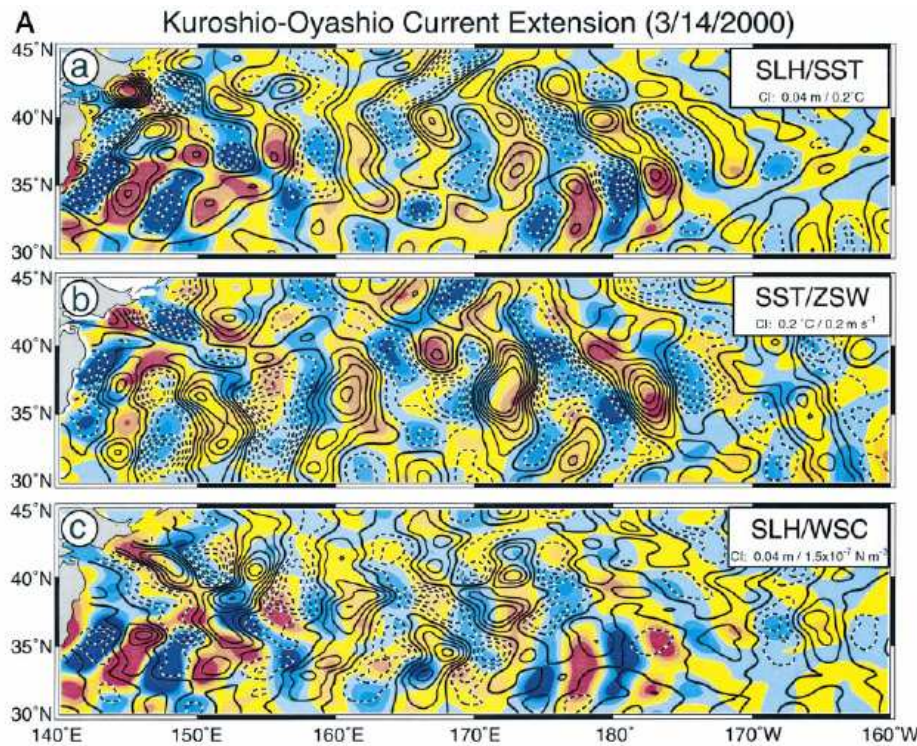
1-day snapshots 7 Mar 2001 - spring



3-month cumulative averages over spring

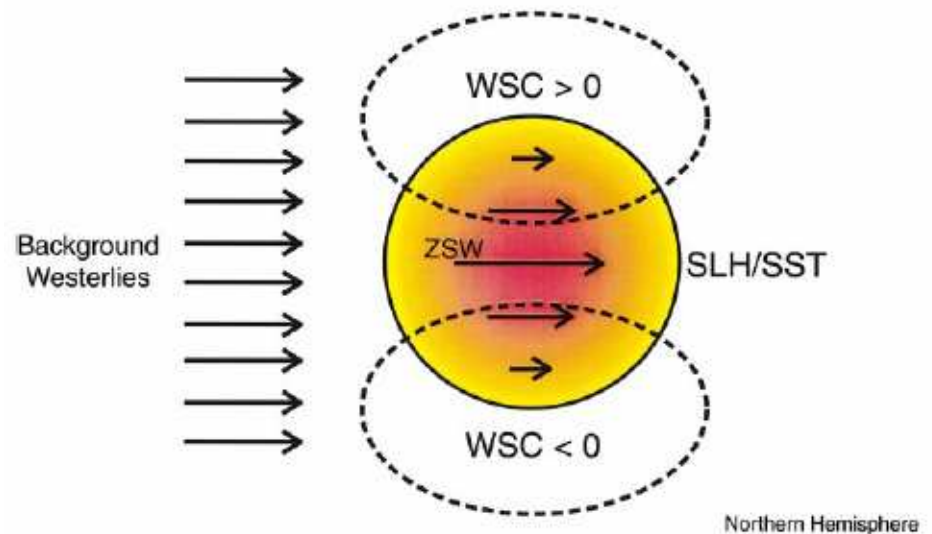
Mesoscale ocean eddies can also impact back onto the atmospheric circulation.

Mesoscale instabilities can be generated by wind anomalies (*Stammer and Wunsch, 1999; Ducet et al., 2000; Eden and Boening, 2003*).



White and Annis, JPO, 2003

Schematic of Mesoscale Eddy in Background Westerly Wind Field



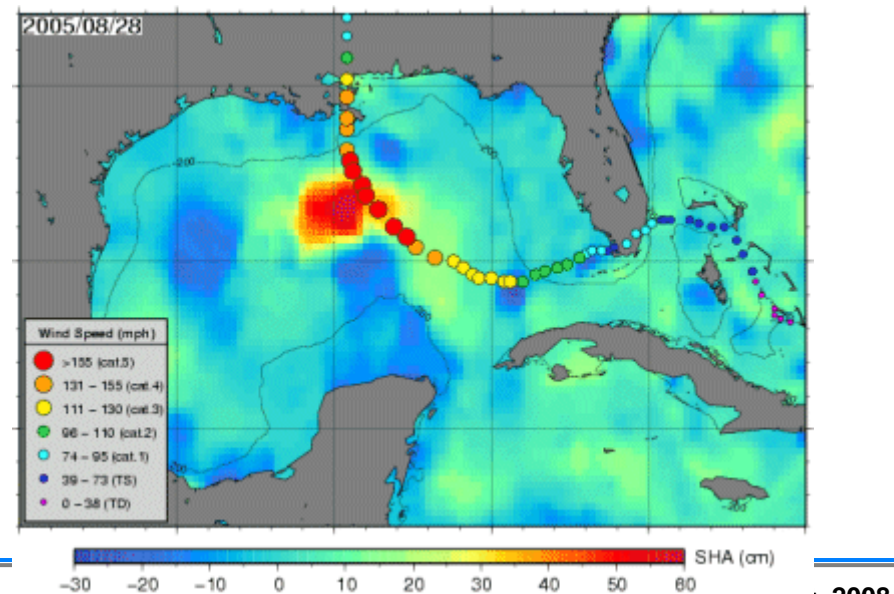
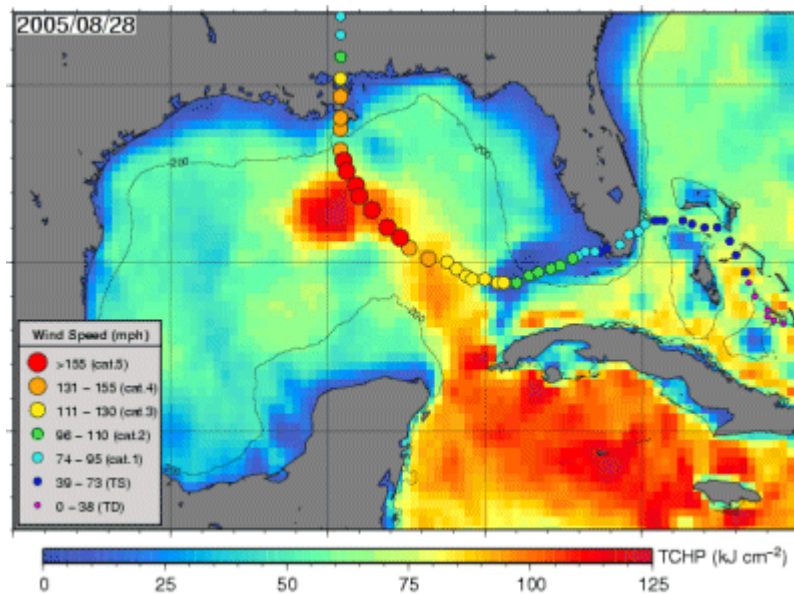
- Feedback onto the ocean
- Displace warm eddies equatorward

SWOT

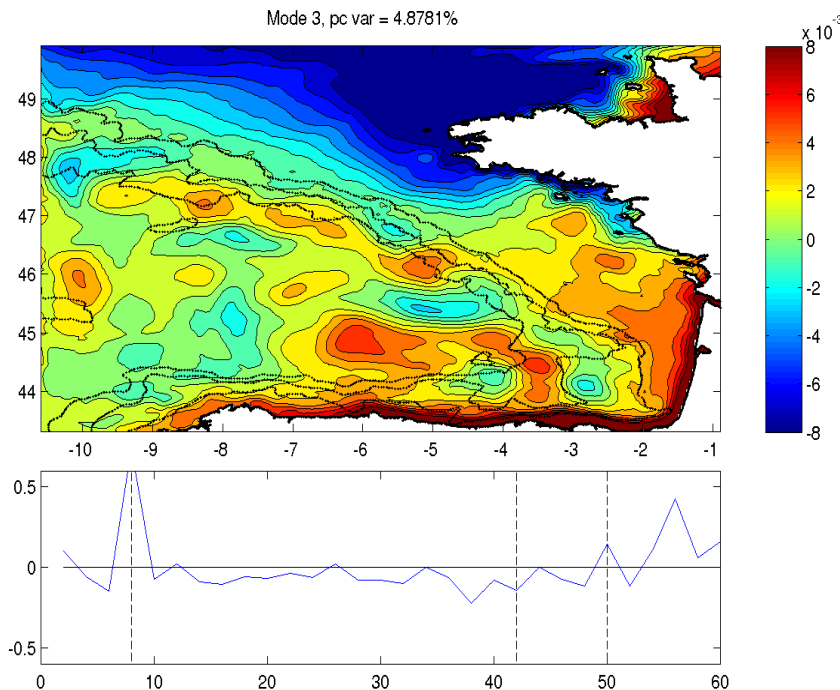
4. Hurricane Predictions



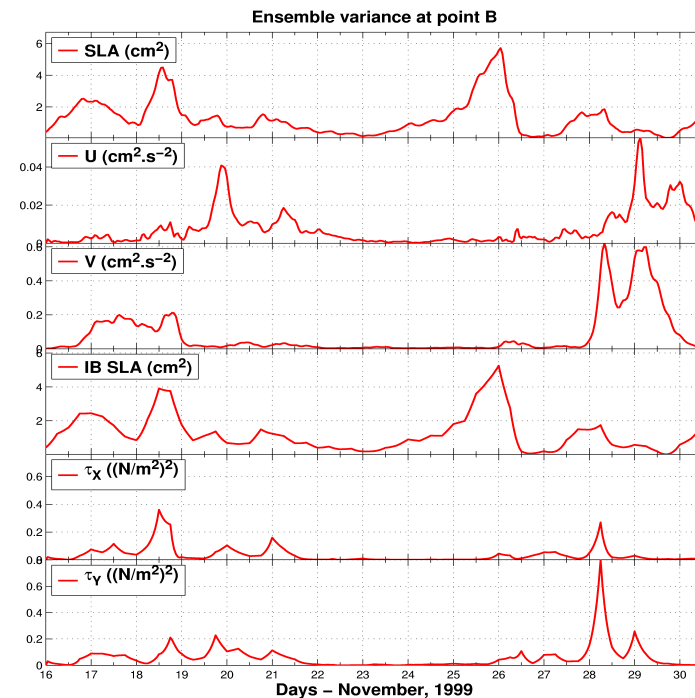
- Higher-resolution altimetry could improve our knowledge of upper ocean heat content and circulation, for improving prediction of extreme events
- These images from US Naval Research Lab. illustrate altimetry combined with sea surface temperature and a two-layer model to monitor ocean heat content for Hurricane Katrina in August 2005.



Spatial structure of a high-resolution coastal model response to atmospheric forcing perturbations



Temporal response of different model parameters to atmospheric forcing perturbations

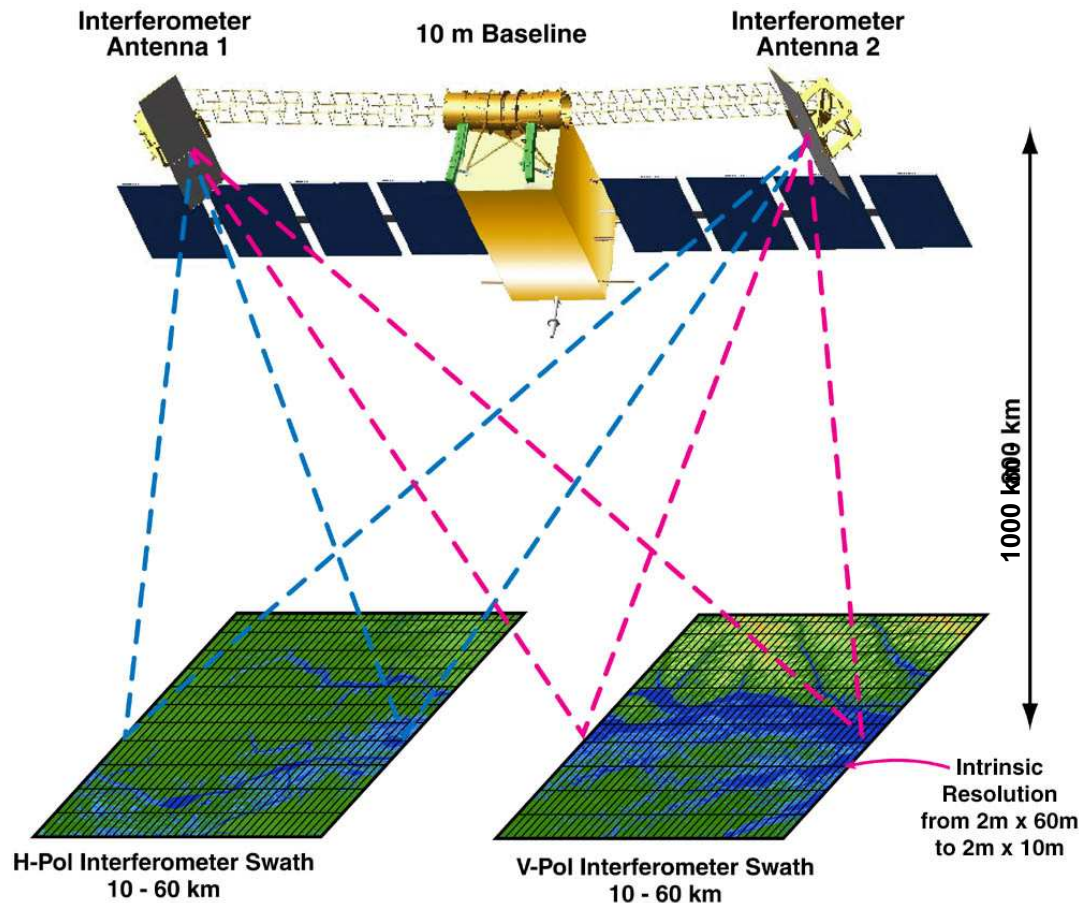


For predicting extreme events in the coastal zone, we need fine spatial resolution altimetry right up to the coast, and at high temporal resolution (~1 day)

15 years of precise satellite altimetry has revealed many new features of the mesoscale ocean circulation, but :

- **Missing** part of the **mesoscale eddy energy** (10-100 km) at mid to high latitudes
- Cannot measure the **sub-mesoscale signals** accurately – important for vertical exchange of heat, carbon, nutrients, biomass
- **Atmosphere** « feels » the ocean's mesoscale eddy field – and perhaps the sub-mesoscale?
- **Coastal regions** are poorly sampled by present altimeters, yet have small cross-shelf space scales and respond over rapid time scales

=> Need for high resolution altimetric measurements in open ocean and coastal zones



Ka band interferometer penetrates clouds and relies on small canopy openings to penetrate to underlying water surfaces (openings of only 20% are sufficient).

Two Ka-band SAR antennae at opposite ends of a **10 m boom**
Both antennae transmit and receive the emitted radar pulses along both sides of the orbital track.

Look angles are limited to 4.5° to reduce the baseline roll-error.

Provides a **120 km wide swath**.

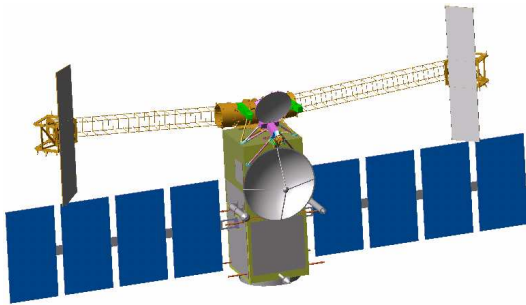
Specular scatter : water bodies scatter more strongly than land.

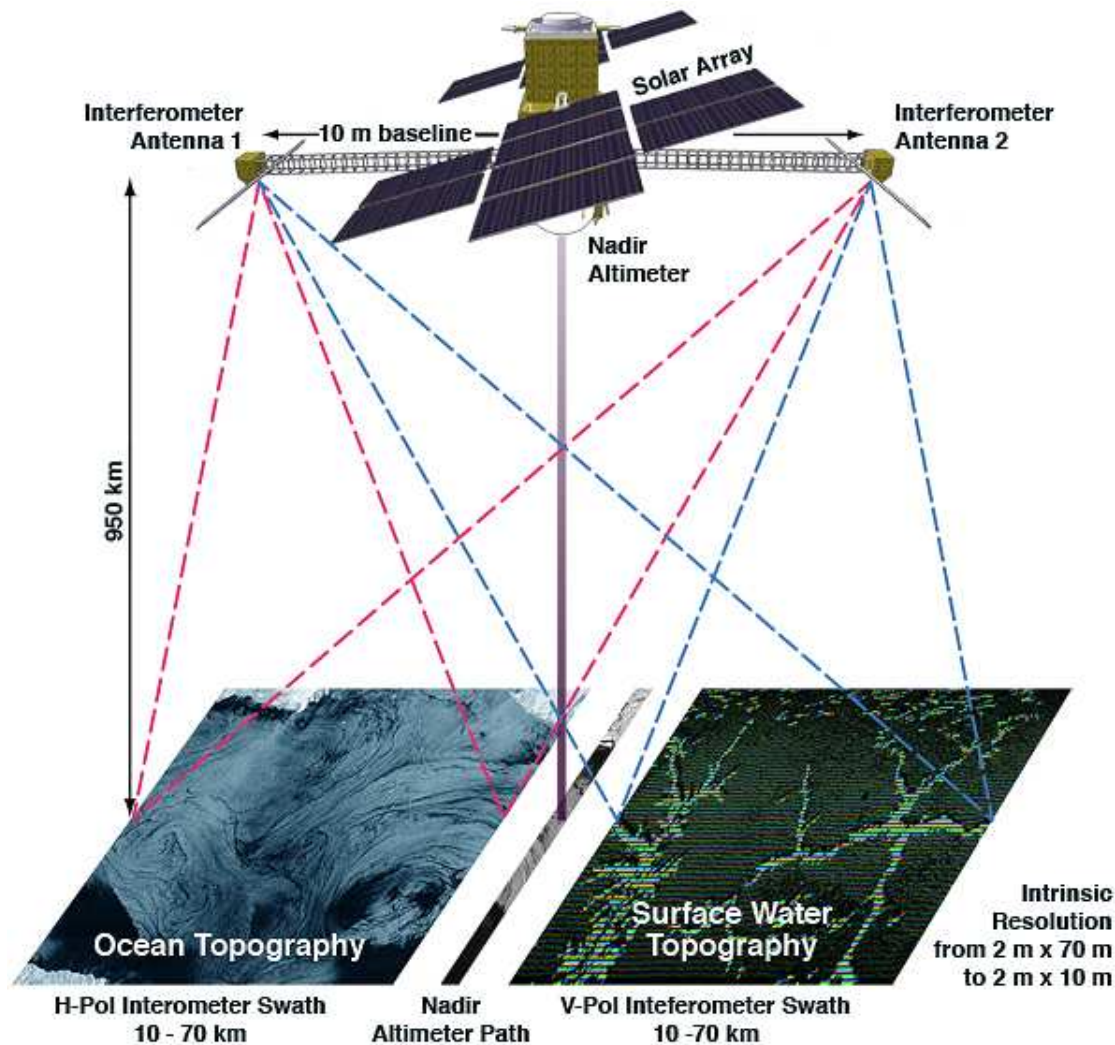
Interferometric SAR processing of the returned pulses yields a **5 m alongtrack resolution**

Crosstrack resolution is 10 m in far swath to 70 m in near swath,
Elevation precision is ± 50 cm.

Polynomial based averaging over areas less than 1 km^2 increases the **height precision** to less than ± 1 cm.

- **Nadir altimeters:** TOPEX/Poseidon & Jason & ERS & EnviSAT Altimeters: sea surface topography (1992 to present)
 - **Heritage:** Error budget for propagation delays, algorithms for range corrections, water reflectivity near nadir
- **Radar Interferometers:**
 - **TOPSAR/AIRSAR:** early 1990's-present: C-band airborne platform
 - **Star3I:** X-band airborne radar interferometer (1990's-present)
 - **GeoSAR:** X-band P-band radar interferometer
 - **Europe:** DLR airborne IFSSAR, DLR X-band spaceborne interferometer with SRTM
 - **Shuttle Radar Topography Mission:** spaceborne land topography (2000)
 - Also imaged rivers and the ocean
 - **Wide-Swath Ocean Altimeter :** centimeter level precision concept funded by NASA past design reviews, but deferred due to budget problems
 - **Cryosat:** forthcoming studies will investigate the use of cryosat interferometric/altimeter modes for surface water.
 - **SWOT:** phase A supported by CNES and NASA in 2008.
 - **Heritage:** error budget verification, instrument design and manufacture, processing algorithms, ground system, mission management, calibration and validation
- **High-frequency Radars:**
 - **CloudSat:** the proposed instrument uses technology and lessons learned from the high-frequency CloudSat mission (EIK, High Voltage Power Supply)





Ka-band SAR interferometric system with 2 swaths, 60 km each

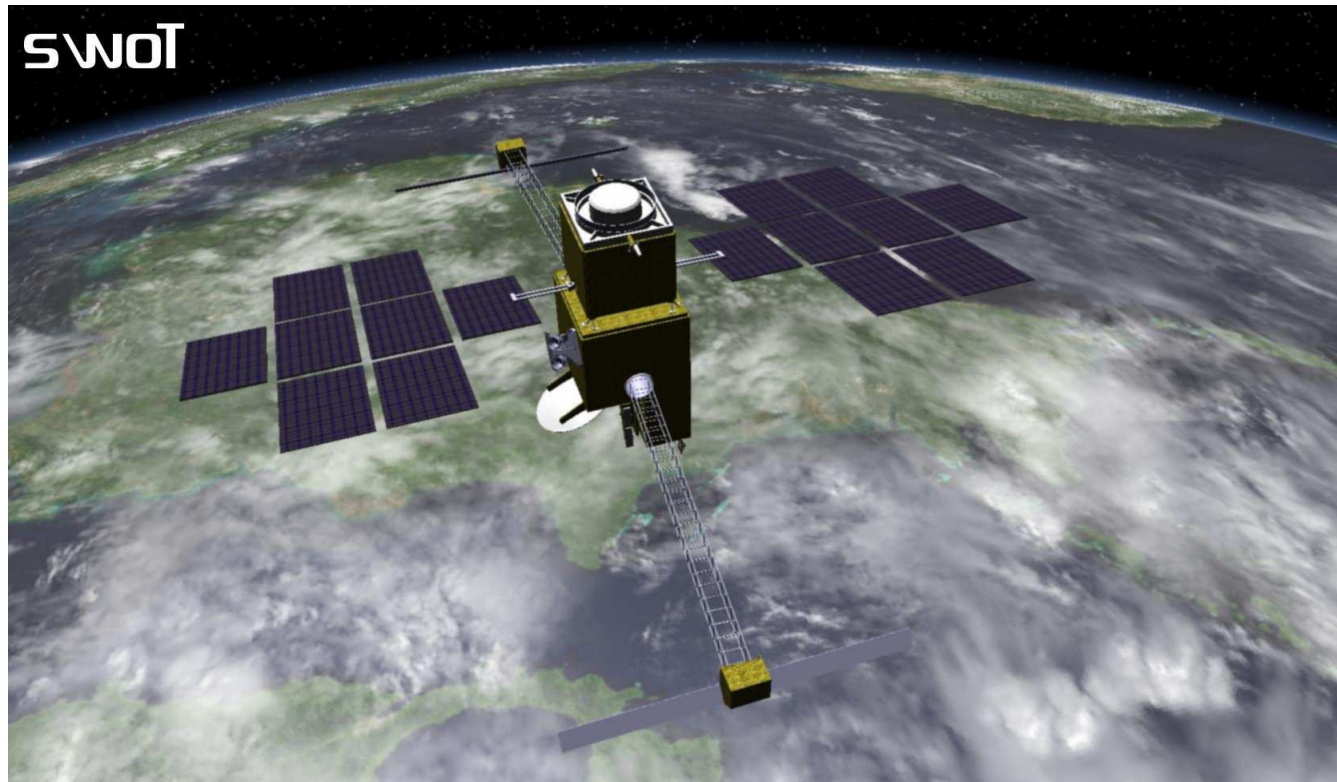
Produces heights and co-registered all-weather imagery *required by both communities*

Additional instruments:

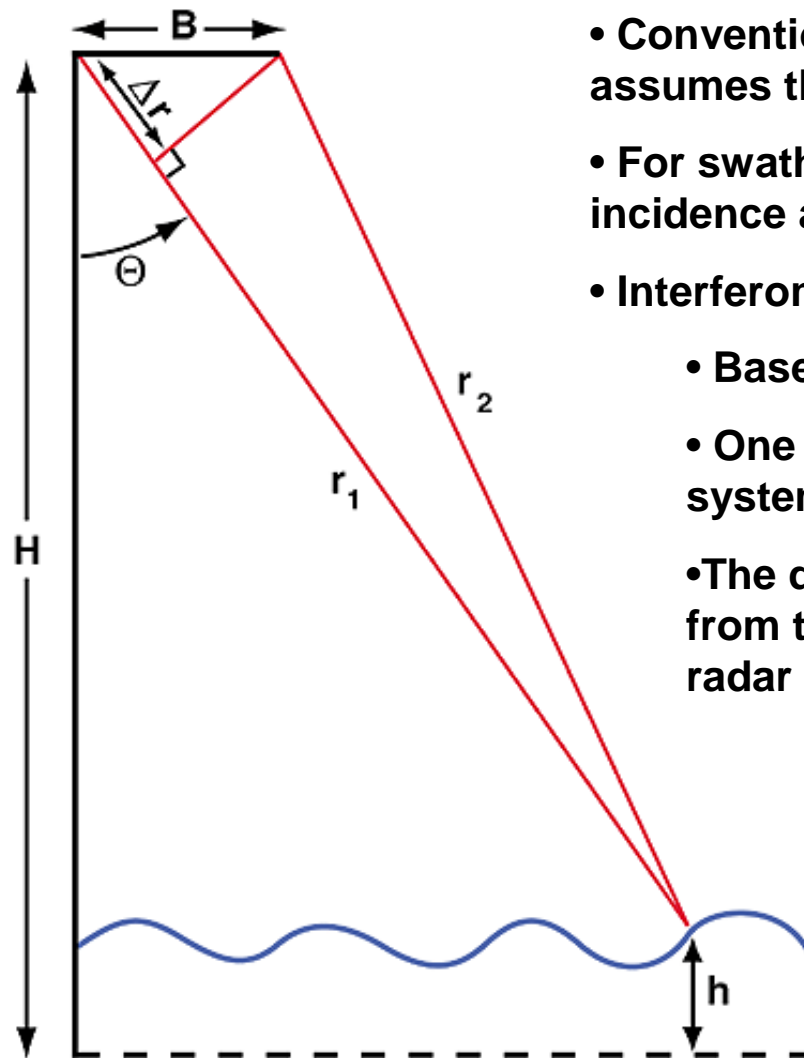
- conventional Jason-class altimeter for nadir coverage
- AMR-class radiometer (with possible high frequency band augmentation) to correct for wet-tropospheric delay

No land data compression onboard (50m resolution)

Onboard data compression over the ocean (1km resolution)



CNES conceptual drawing



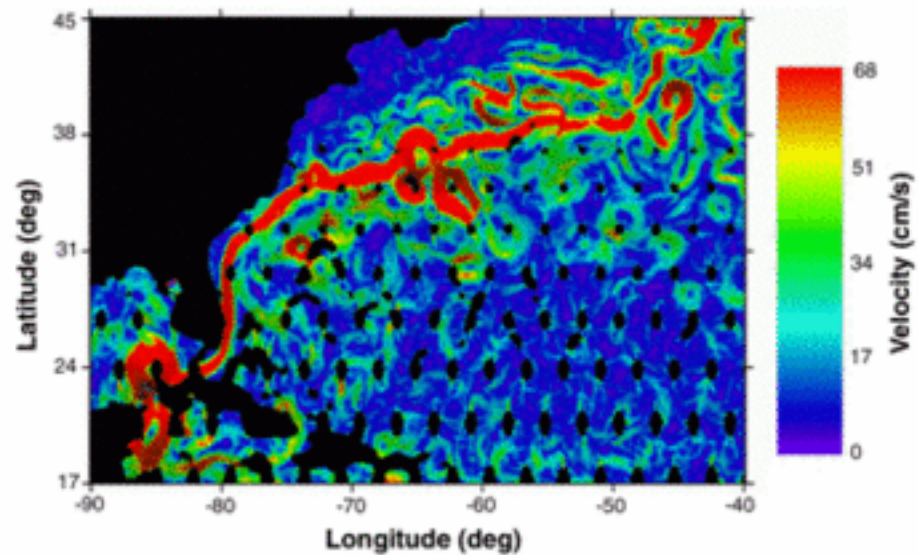
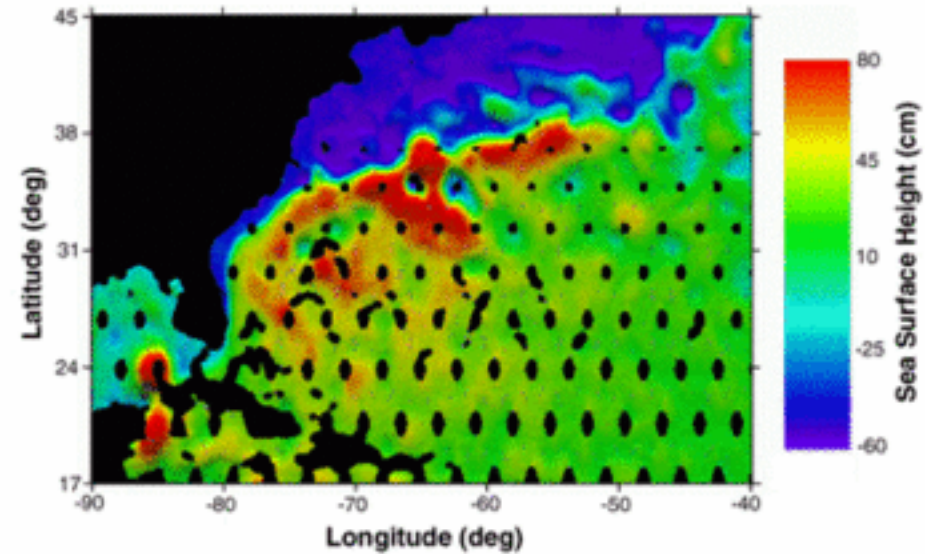
- Conventional altimetry measures a single range and assumes the return is from the nadir point
- For swath coverage, additional information about the incidence angle is required to geolocate
- Interferometry is basically triangulation
 - Baseline B forms base (mechanically stable)
 - One side, the range r_1 , is determined by the system timing accuracy
 - The difference between two sides (Δr) is obtained from the phase difference (Φ) between the two radar channels.

$$\Phi = 2\pi \Delta r / \lambda = 2\pi B \sin \Theta / \lambda$$

$$h = H - r_1 \sin \Theta$$

Typical spatial coverage from a 10-day Jason-1 orbit

Mid-to-high latitude, « crossover » points covered many times in 10 days.

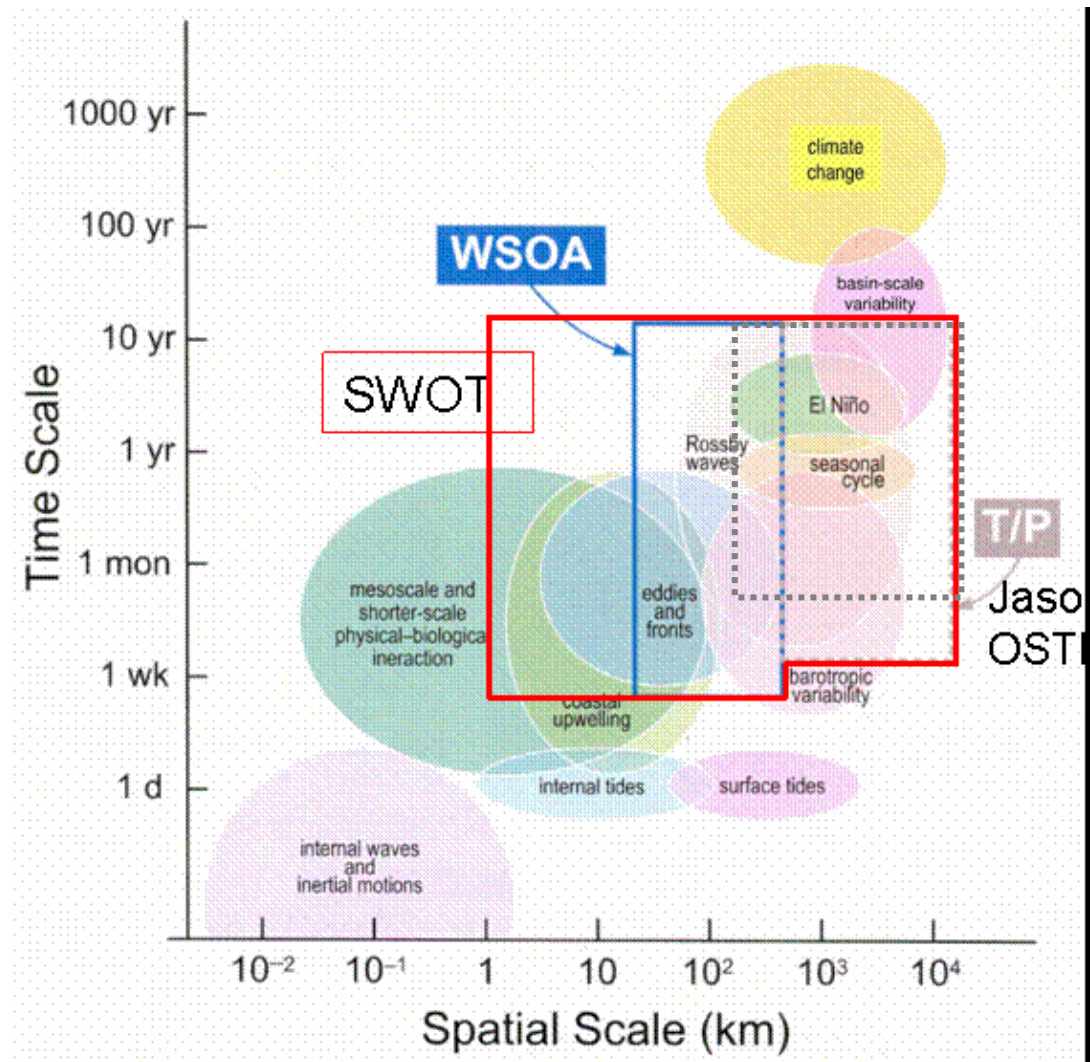


Original WSOA technology could resolve to 15-20 km resolution.

SWOT resolves to 1 km resolution

Good for sub-mesoscale ocean dynamics

Excellent for monitoring small lakes and river levels



$$\delta h = \delta H -$$

Orbit Error

$$\cos(\theta)\delta r +$$

Media Delay Error (Iono, Tropo, EMB)

$$r \sin \theta \frac{\partial \theta}{\partial \Phi} \delta \Phi +$$

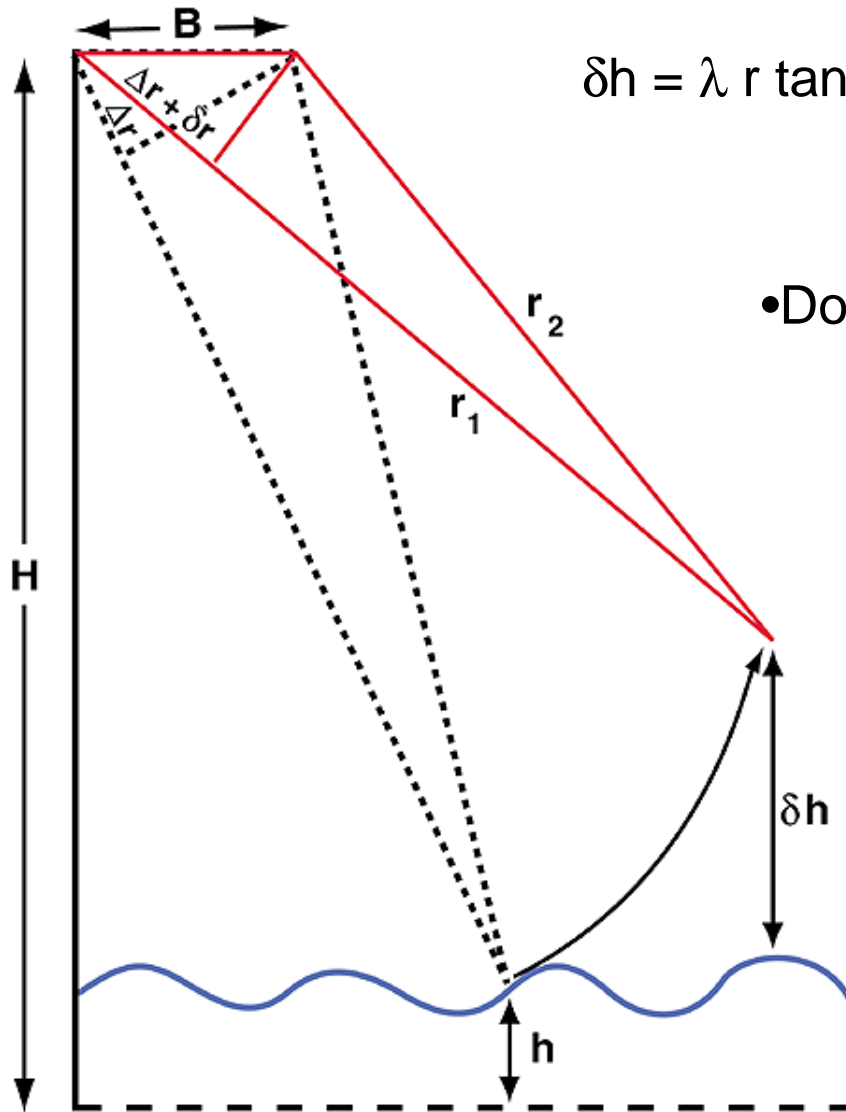
Phase Error

$$r \sin \theta \delta \theta$$

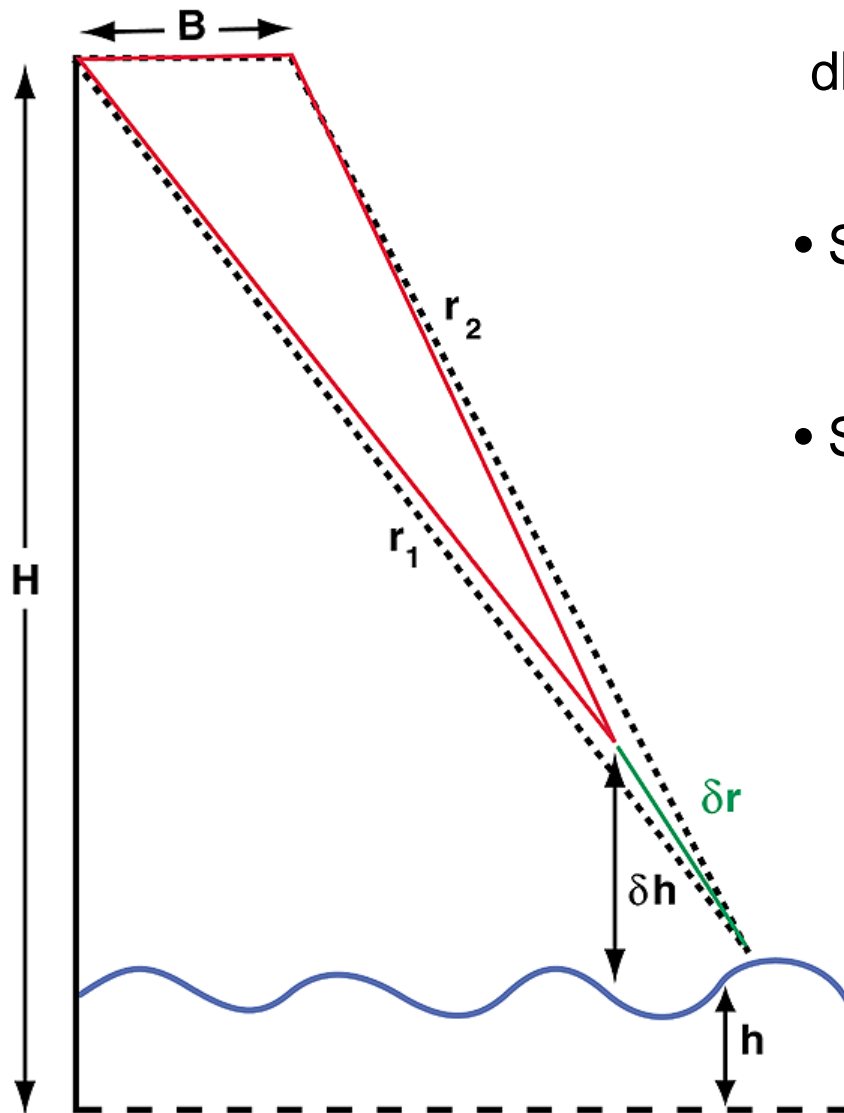
Baseline Roll Error

Other error sources (e.g., baseline length, yaw errors) can be controlled so that errors are smaller by an order of magnitude, or more.

- Errors can be divided into spatially correlated and uncorrelated
 - Uncorrelated: thermal/speckle noise. Precision improves linearly with the area
 - Correlated: geophysical, orbit. Precision does not improve significantly with averaging
- Slope (\sim velocity) is affected differently than height by spatially correlated errors
 - Relatively large height errors can result in relatively small slope errors
- For ocean, geostrophic velocity \sim slope. Sea-level rise \sim height. Heat content \sim height
- For hydrology, velocity (discharge) \sim slope (or assimilated height). Storage \sim height

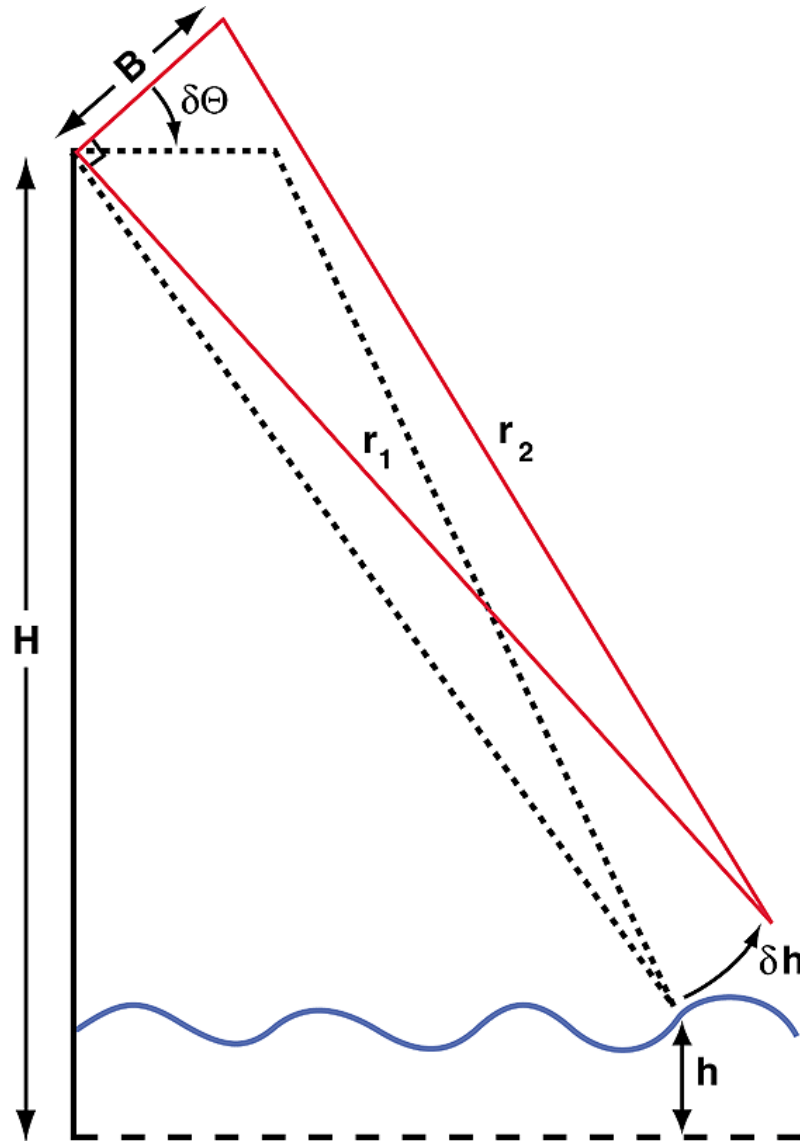


- Dominant sources of phase noise:
 - Thermal noise in radar signal (random)
 - Decorrelation of the two returns due to speckle decorrelation of scattered fields (random)
 - Phase imbalance between the two interferometric channels:
 - Temperature driven (slow change)
 - Can be calibrated using calibration loop.



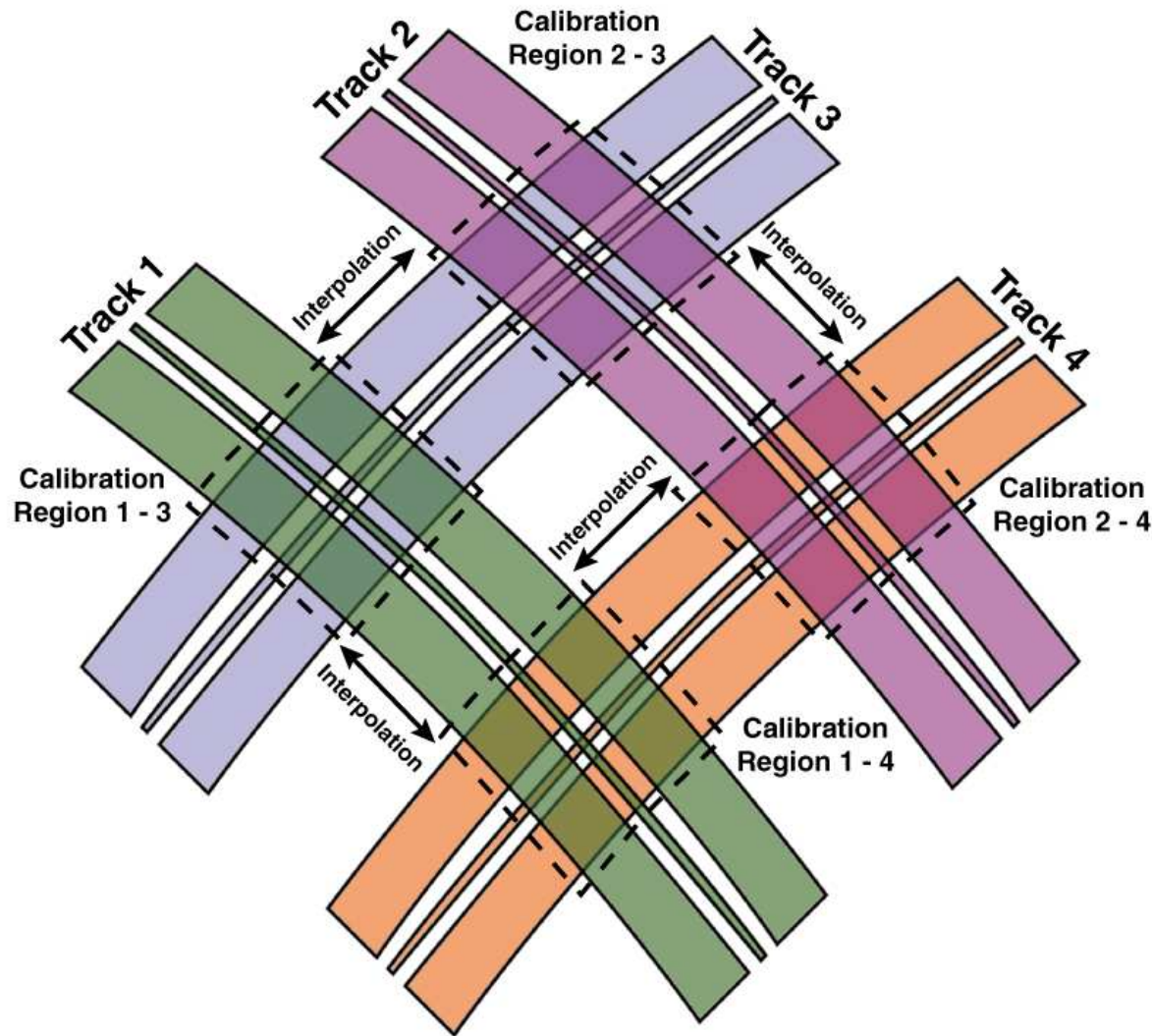
$$dh = -\delta r \cos \Theta$$

- Similar to nadir altimeter range errors
- Sources of range error:
 - Ionospheric delay
 - Dry and wet tropo delays
 - EM Bias



$$\delta h = r \sin \Theta \delta \Theta$$

- An error in the baseline roll angle tilts the surface by the same angle.
 - This is equivalent to introducing a constant geostrophic current in the along-track direction
- As an order of magnitude, a 0.1arcsec roll error results in a 4.5cm height error at 100km from the nadir point
- Roll knowledge error sources:
 - Errors in spacecraft roll estimate
 - Mechanical distortion of the baseline (can be made negligible if the baseline is rigid enough)



- Roll errors must be removed by calibration

- Assume the ocean does not change significantly between crossover visits (< 10 days)

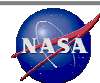
- For each cross-over, estimate the baseline roll and roll rate for each of the passes using altimeter-interferometer and interferometer-interferometer cross-over differences, which define an over-constrained linear system.

- Interpolate along-track baseline parameters between calibration regions by using smooth interpolating function (e.g, cubic spline.)

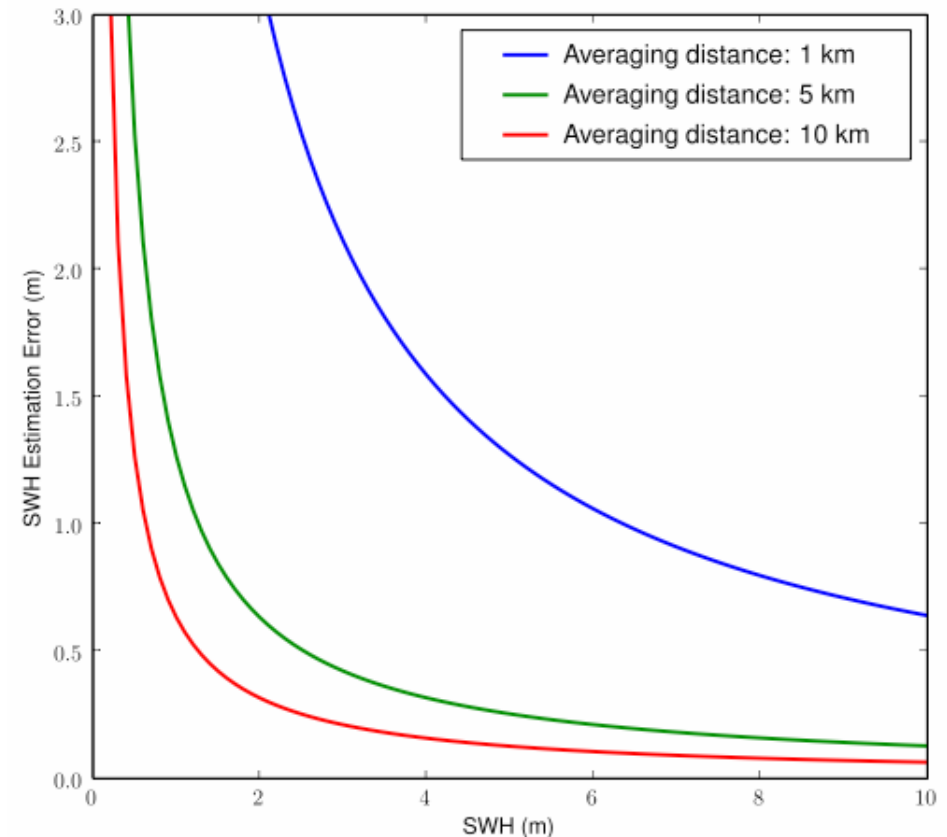
	Height Error	Typical wavelength	Typical geophysical magnitude	Slope error
Phase error	1cm @ 1km ² area	uncorrelated	NA	0.7μrad @ 5km 0.15μrad @ 10km
Roll errors	2cm (max) at swath edge	200km	Depends on platform dynamics	0.1μrad (max)
Wet Troposphere	1cm (Ocean) <5cm (Land)	>100km	3-6cm	0.1μrad
Dry Troposphere	<1cm	>1000km		0.01μrad
Ionosphere	1cm	>900km	50cm	0.2μrad
EM Bias	1-2cm	100km	<5cm	0.1μrad
Orbit	1cm	>8000km		

SWOT

Other geophysical parameters



- The effect of waves is to increase the observed height variance
- This is a small effect on the height precision (on a single pixel, random noise $\sim 2\text{m}$ SWH)
- SWH can be estimated by estimating the excess variance relative to the predicted variance
- To make a meaningful measurement, a large area must be used for averaging
- The area required is not that different from the altimeter area used for SWH



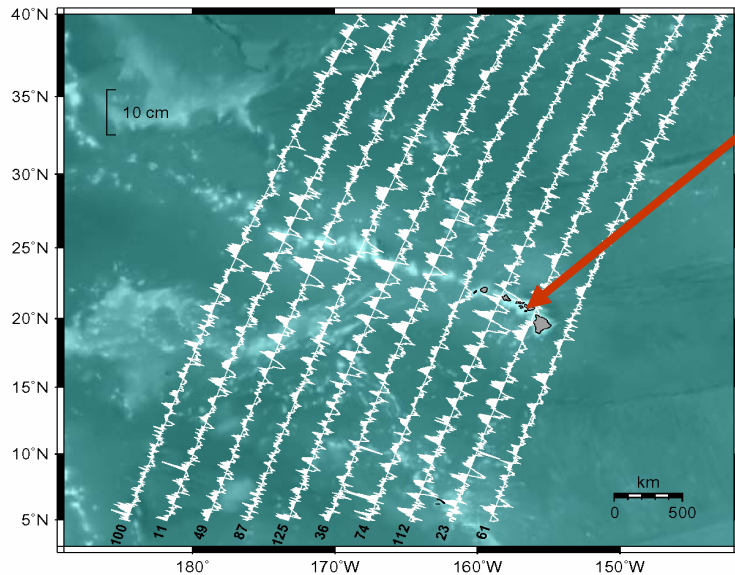
- SWOT will measure radar sigma0 at 1km resolution
- Sigma0 can be converted to wind speed (without direction)
- Similar to traditional altimetric data

The SWOT orbit has been chosen to resolve the major tidal frequencies

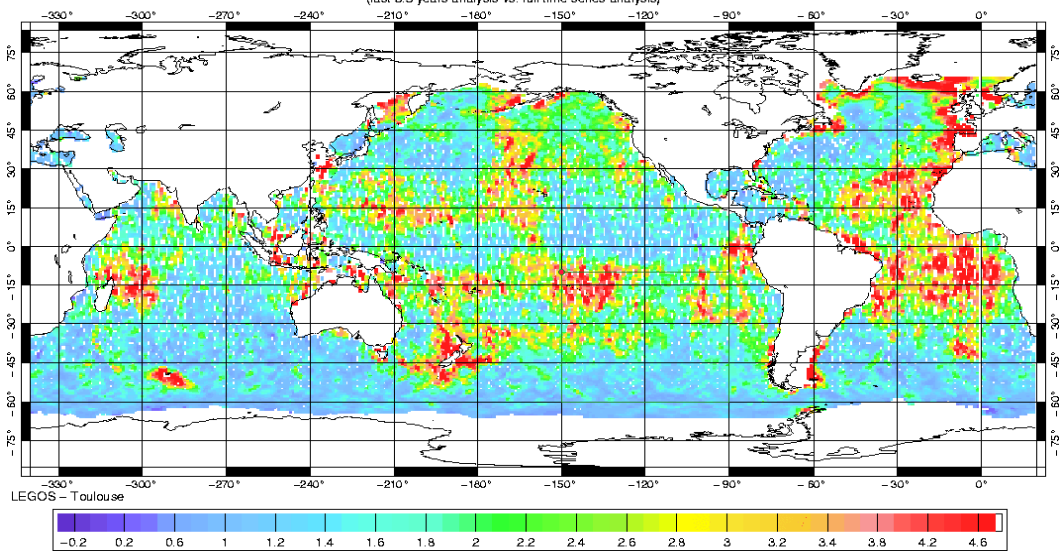
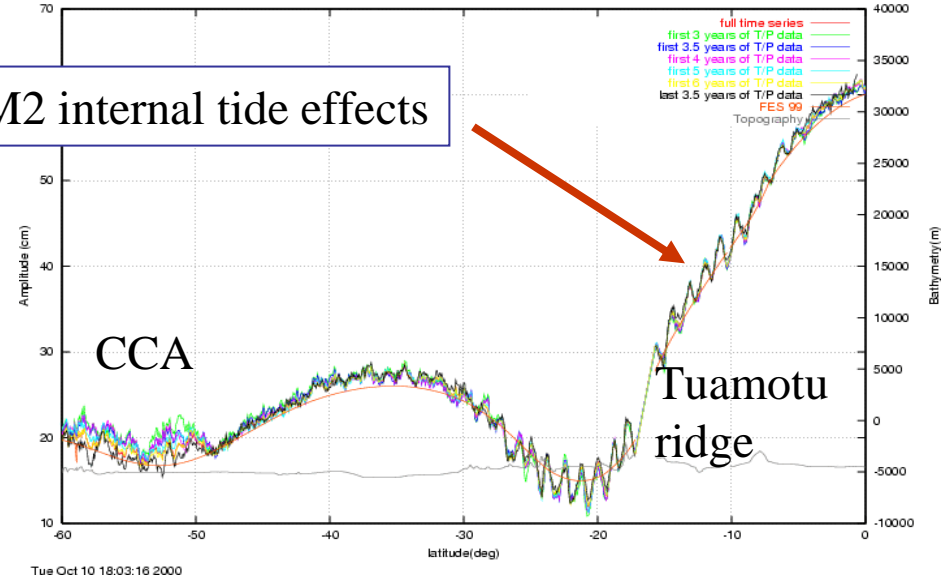
After a few years :

- Fine resolution repeat observations will greatly improve tidal models, especially in coastal zones.
- Possibility of resolving internal tide amplitudes

Initial tidal corrections for SWOT will be imprecise in regions not sampled by traditional altimetry (between groundtracks, coastal zones)



M2 internal tide effects

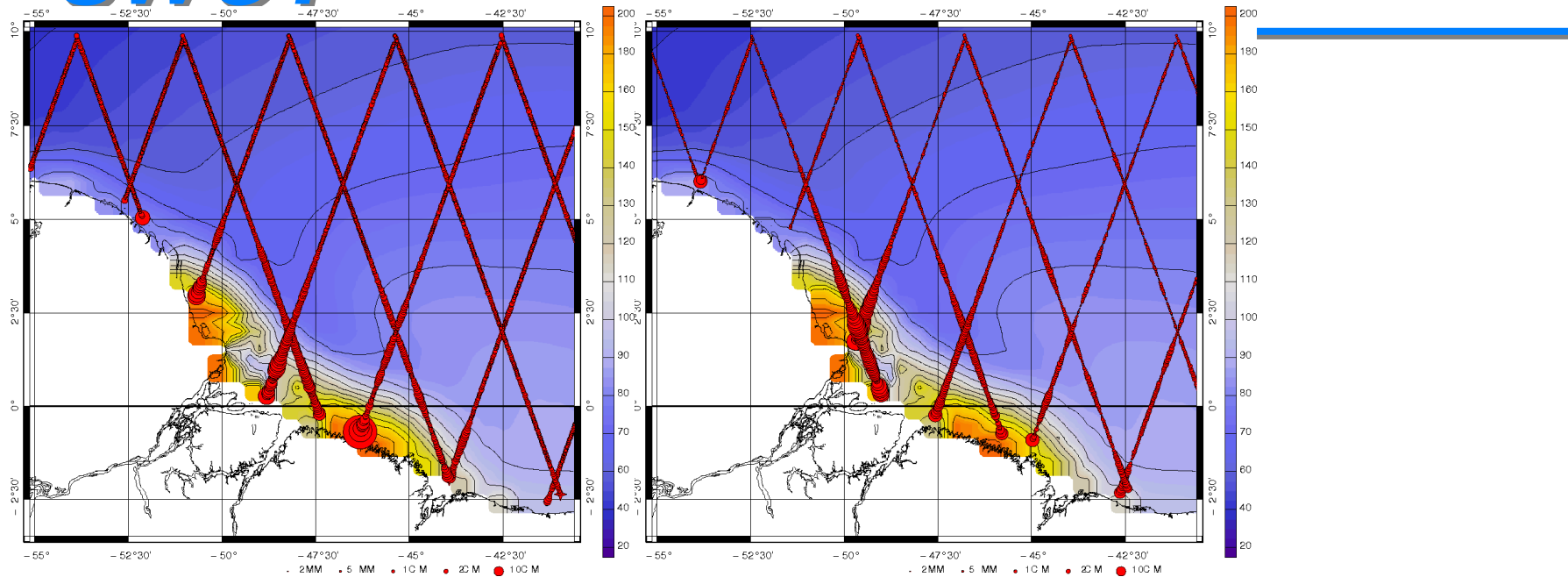


Internal tide signal (2-3 cm) can be resolved from traditional alongtrack data with long time series

SWOT will provide data between groundtracks – finer resolution around bathymetry

SWOT

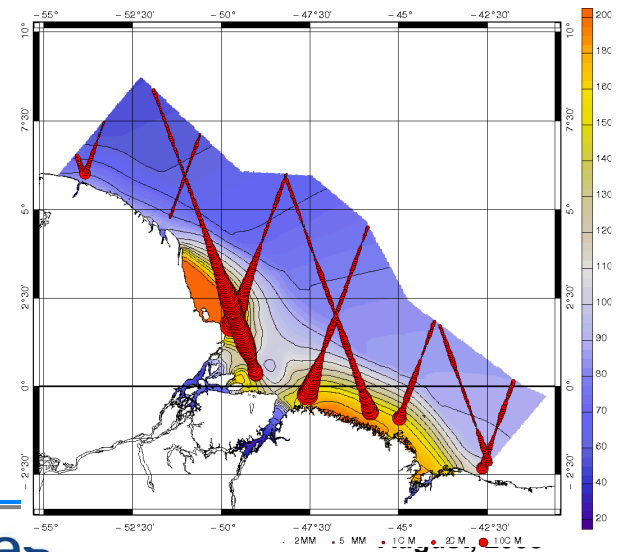
Improving Shelf/coastal tides



GOT00 M2 tide (amplitude in cm) against T/P (right) and T/P interlaced (left)

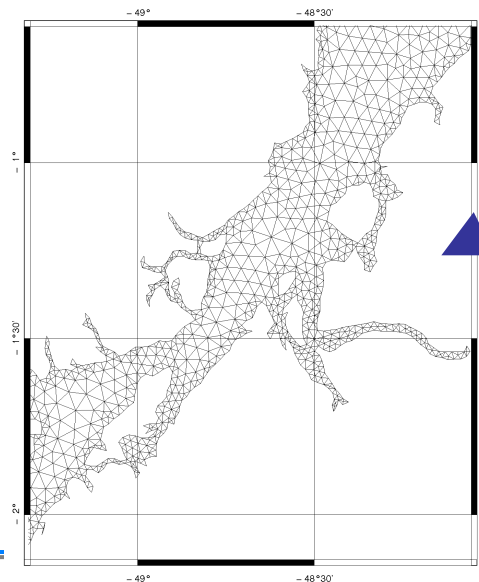
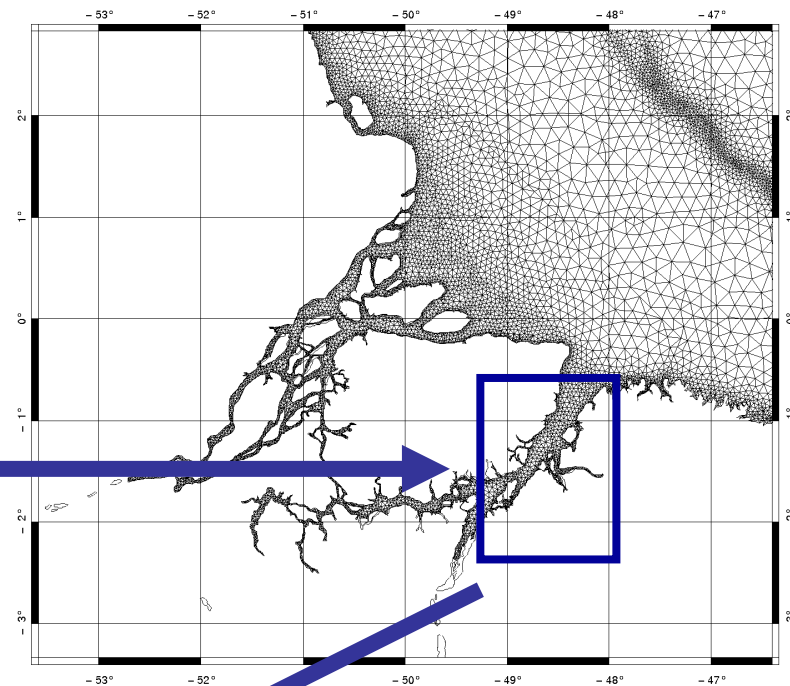
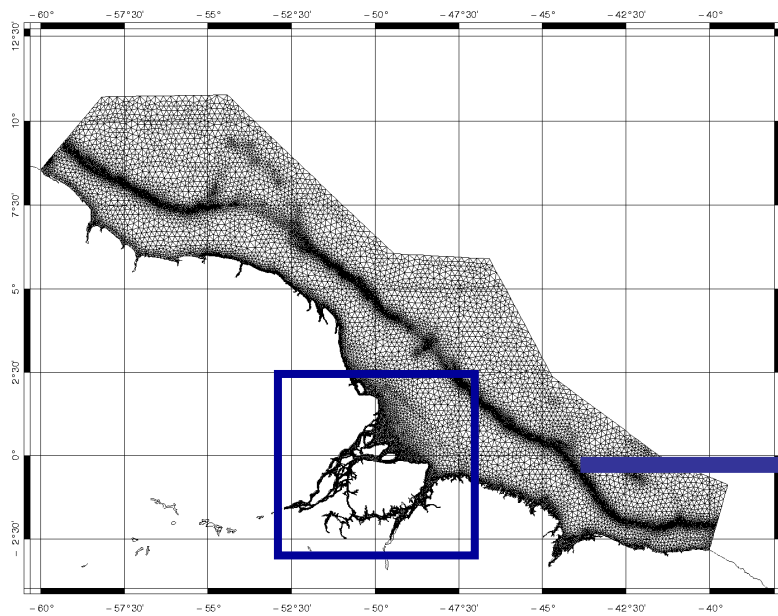
- Loss of data in coastal regions induce an increased error budget in tidal analysis, thus in tidal models
- Harmonic analysis accuracy is limited by local effects

We need to improve the shelf/coastal tidal atlases using as regional models-> then SWOT



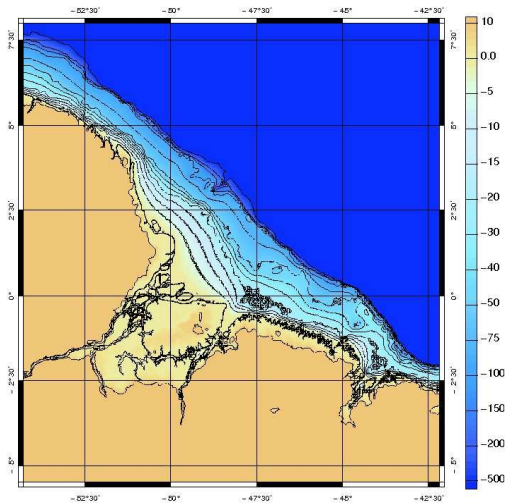
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Regional modelling

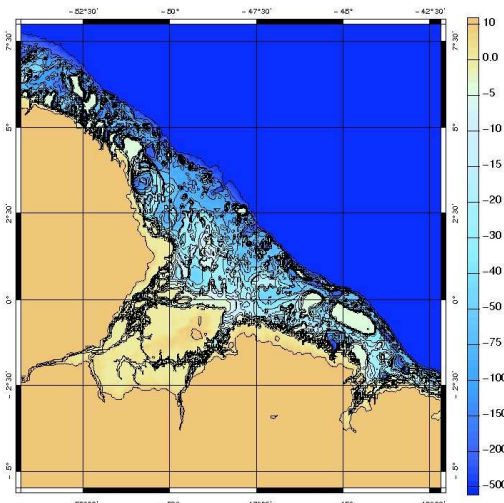


FE mesh useful to properly describe the regional, estuary and river channel details

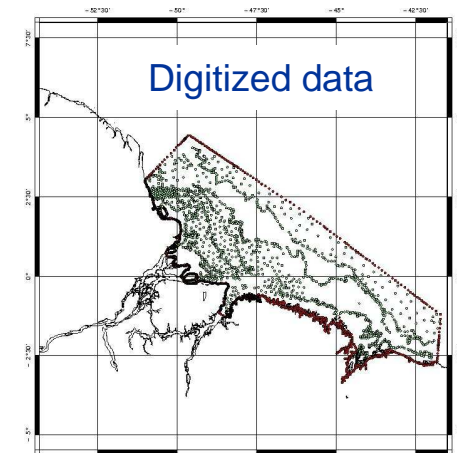
Amazon river shelf



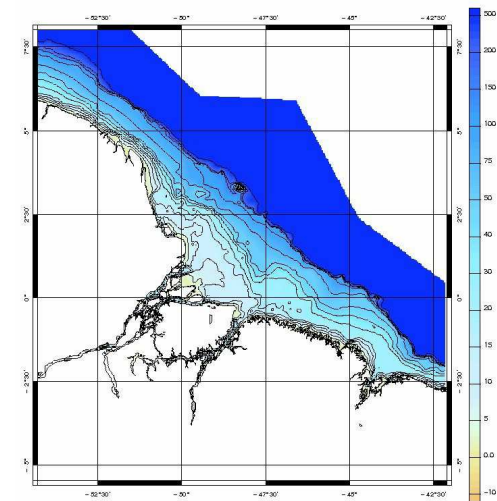
GEBCO (gridone) : lack of details



Smith & Sandwell : noisy and unrealistic



© Pôle d'océanographie Côtière (POC) - Toulouse

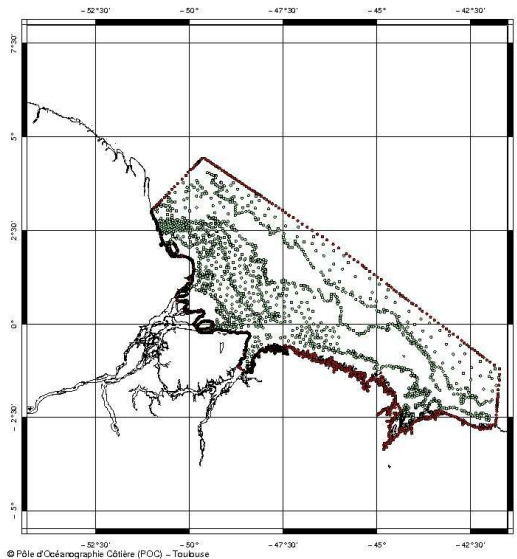


Large bathymetry errors are found in available databases. Effects on tidal simulation are dramatic.

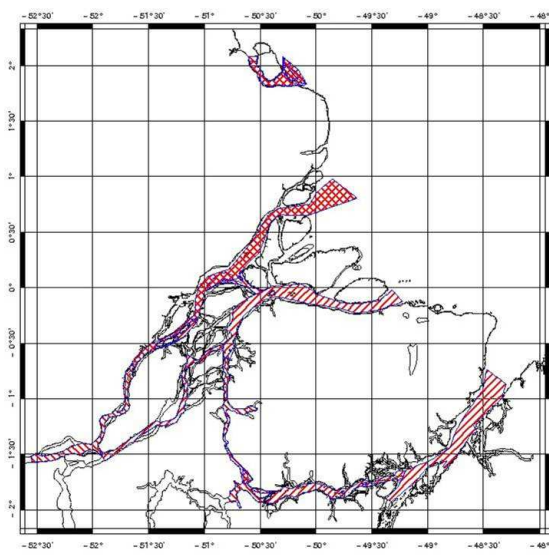
Scan and digitize bathymetric and nautical charts.
Mapped and merged into global database

Urgent need to improve regional bathymetries...

improved model bathymetry

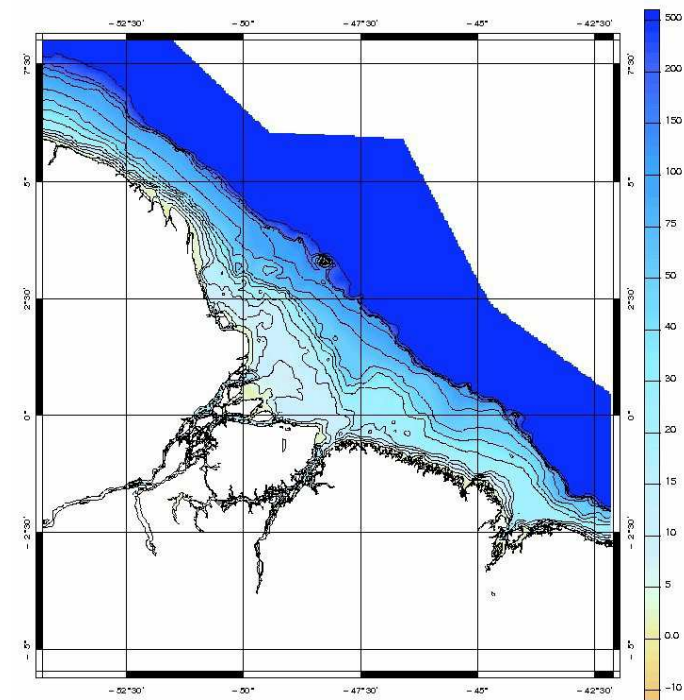


Digitized data



Re-worked channel path

improved model bathymetry



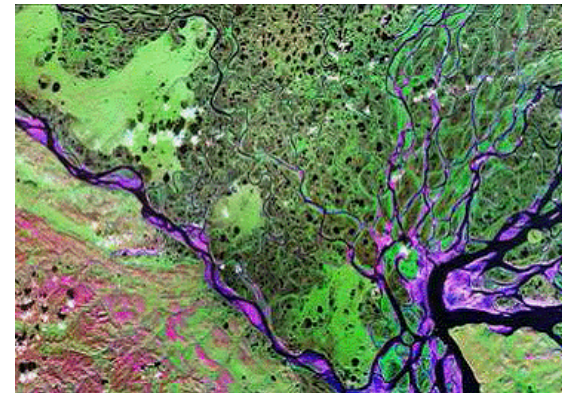
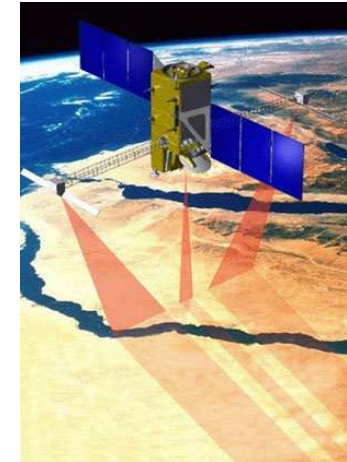
Scan and digitize bathymetric and nautical charts

Special care to estuarine channels

Mapped and merged into global database

SWOT will be the **first satellite mission dedicated to surface water land hydrology**. SWOT will measure water storage changes in all wetlands, lakes, and reservoirs and estimate river discharge.

Global, high-spatial resolution measurements of ocean surface topography will monitor the **full range of ocean currents and eddy fields**, help determine their role in the **air-sea exchange**, and finely map the **coastal transport**. For the first time, we will observe the sea level variations driving **fine-scale filament structures**, which play a decisive role in the vertical exchange of heat, carbon and nutrients.



SWOT

Extras

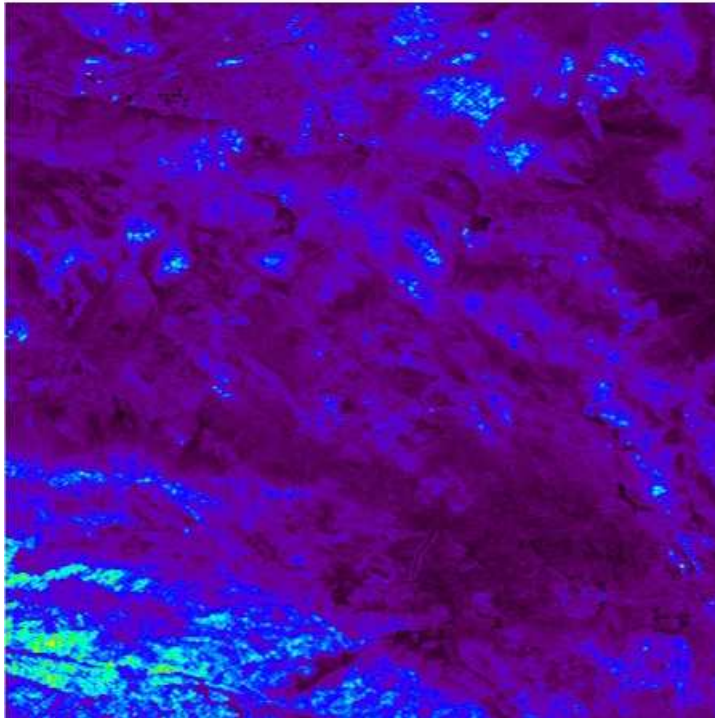


1-44

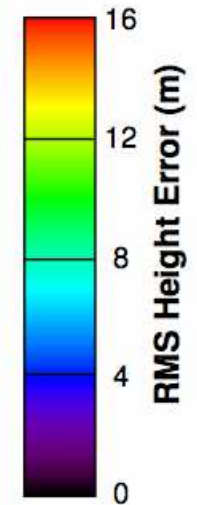
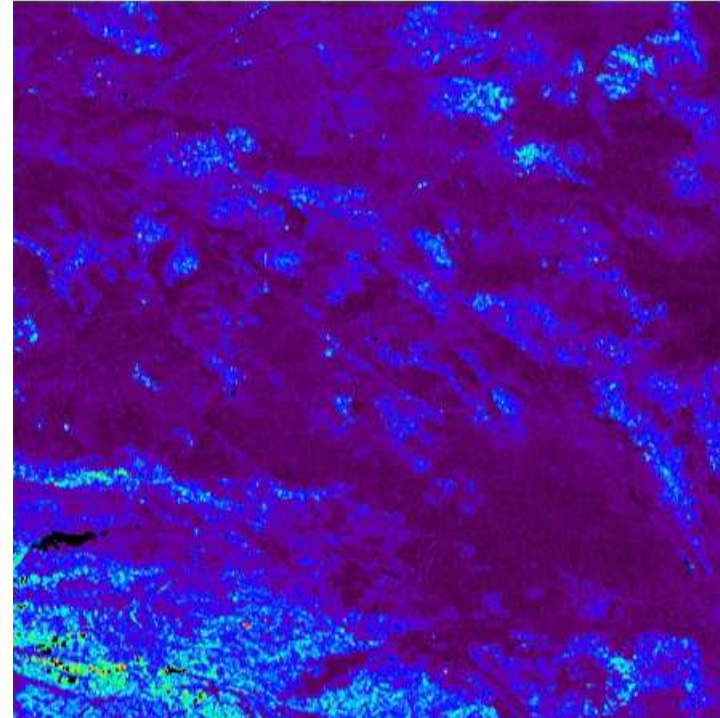


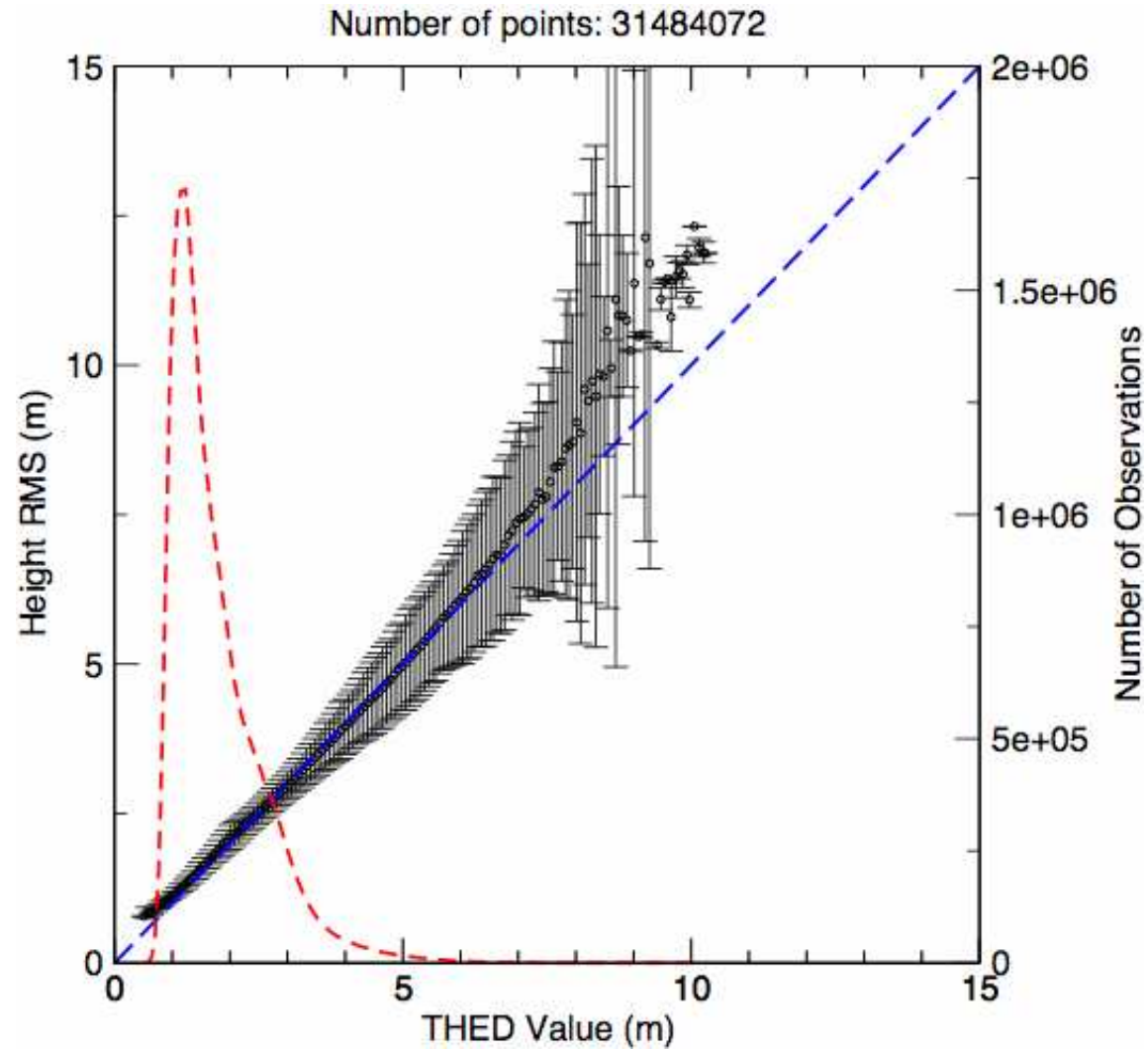
August, 2008

29 Palms THED Values

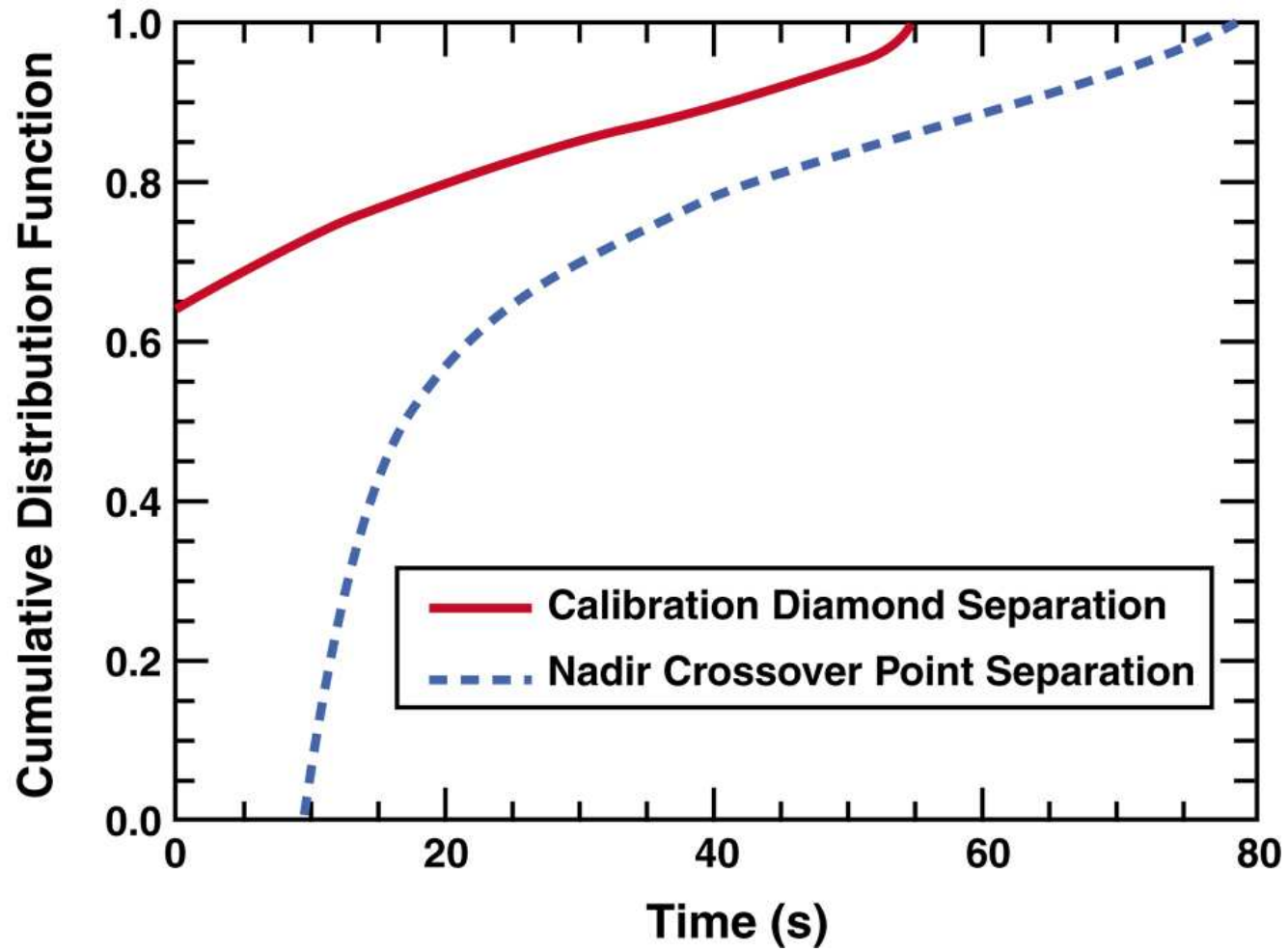


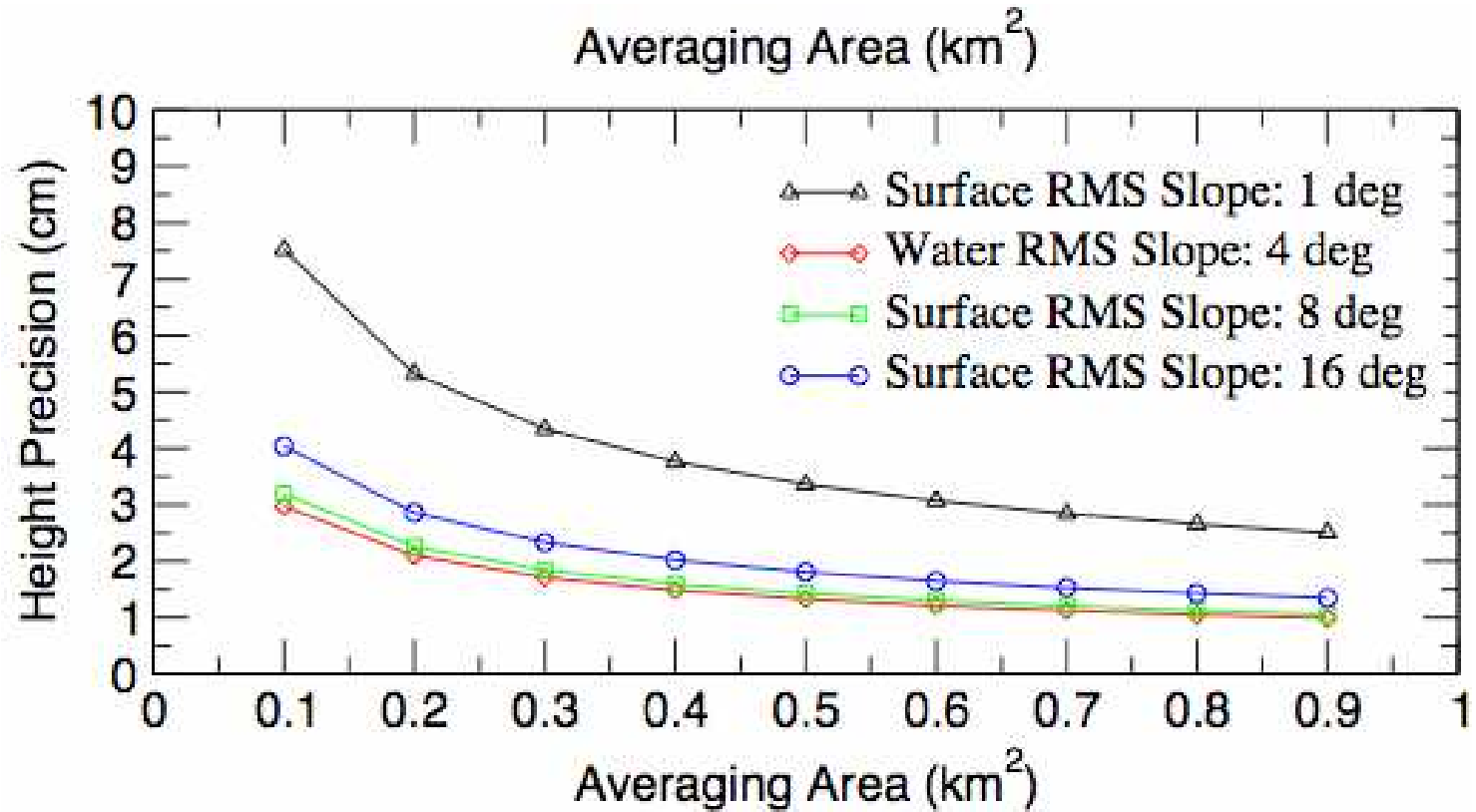
29 Palms Random Height Error

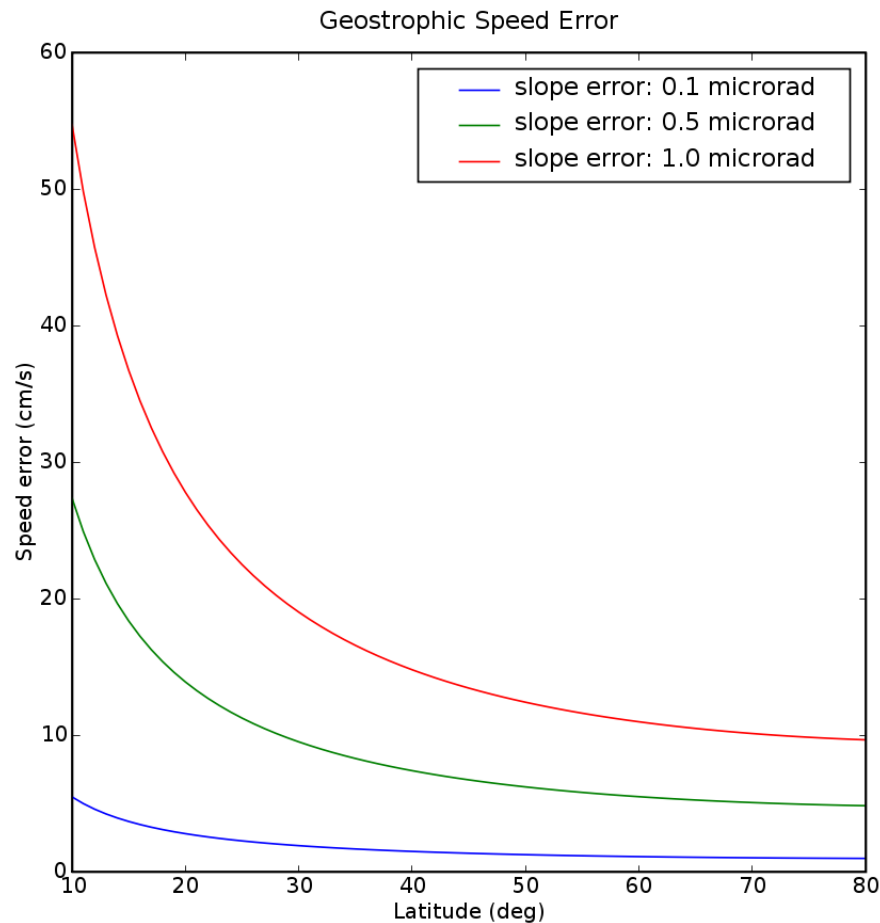




Temporal Separation Scales







$$\sigma_s \approx \sqrt{\frac{12}{N}} \frac{\sigma_h}{L}$$

$$\sigma_V = \sqrt{2} \frac{g}{f} \sigma_s$$

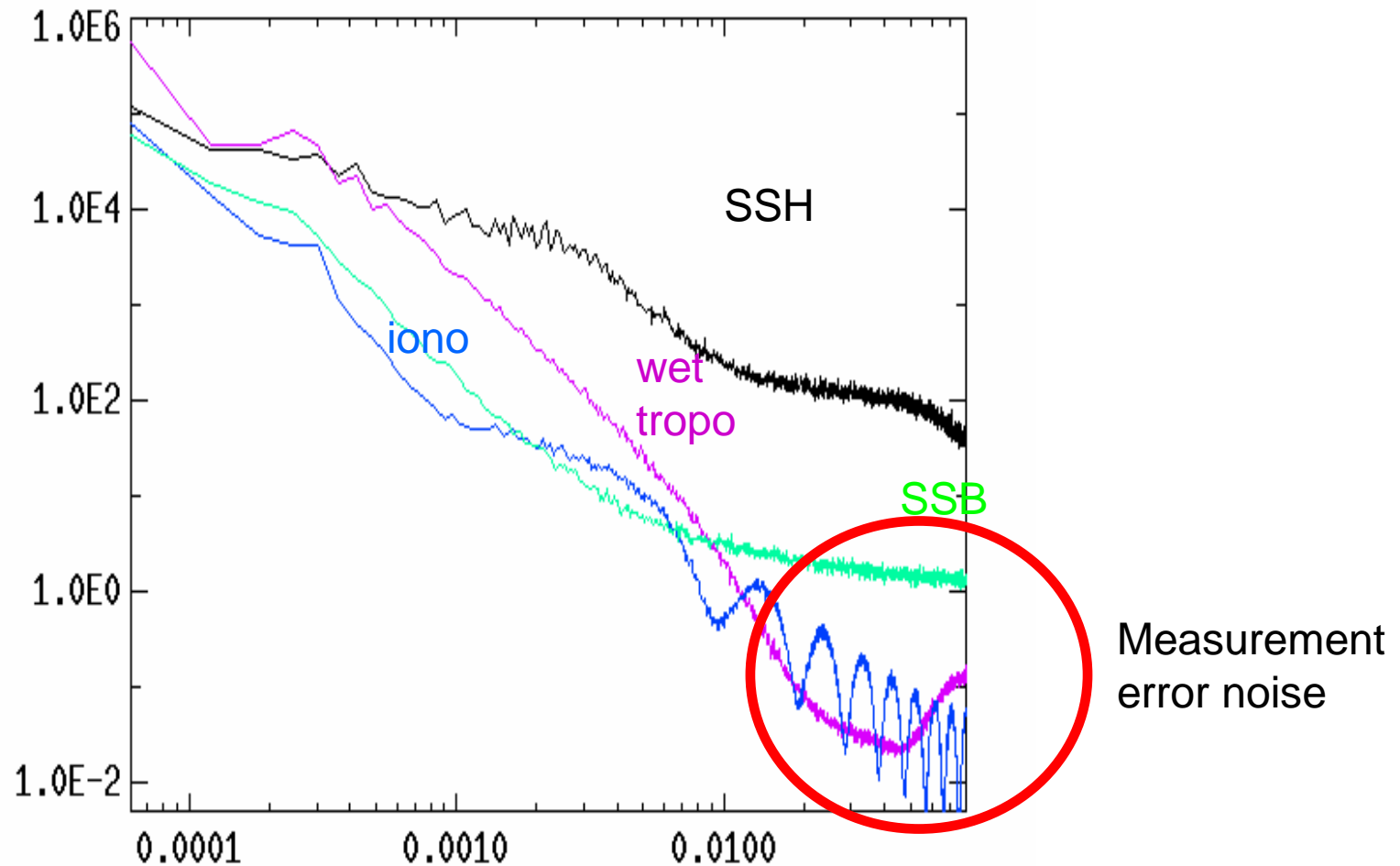
σ_s : Slope std

σ_h : Height std

L : Distance for slope calculation

Δx : Sample spacing

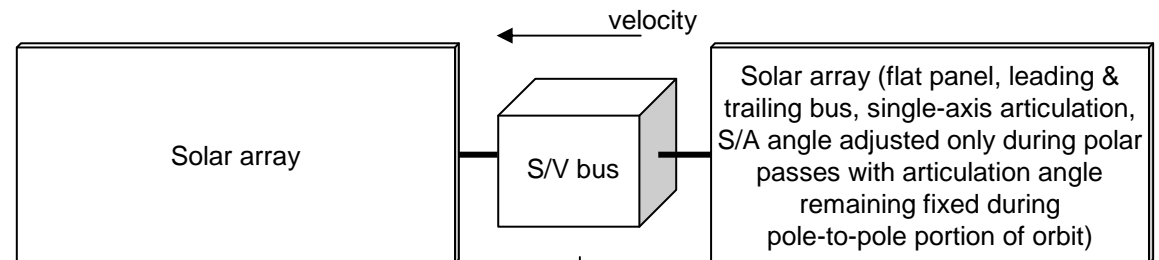
N : Total number of samples



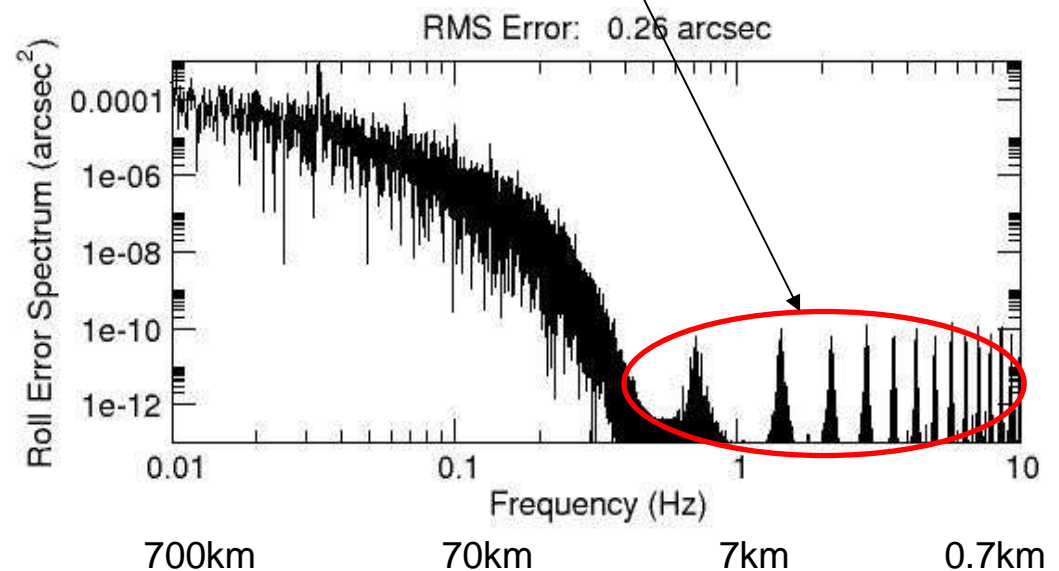
Media errors and sea-state errors have scales larger than 100 km and not affecting submesoscale SSH measurement.

Minimizing high-frequency motion errors can be achieved with an appropriate architecture (e.g., Grace has no moving panels)

Both CNES and JPL have determined that a feasible architecture exists where no high-frequency spacecraft component motion will occur during data collection



Need to minimize these



- The slope accuracy and spatial resolution are compatible with Abyss mission requirements, for even 1 repeat cycle (not taking into account ocean mesoscale contamination)
- Using compromise orbit and expanded swath (120km -> 140km), there are no holes in the coverage

